



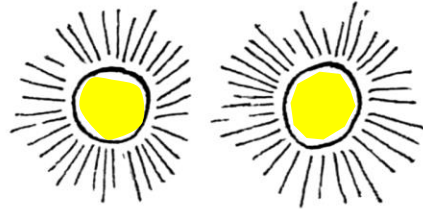
Demystifying Climate Change

Session 2

Our Goldilocks Earth: A Radiative Balancing Act

OLLI at Illinois
Spring 2021

D. H. Tracy



"It looks as if our troubles are only beginning."

New Yorker 3 Dec 2018

Course Outline



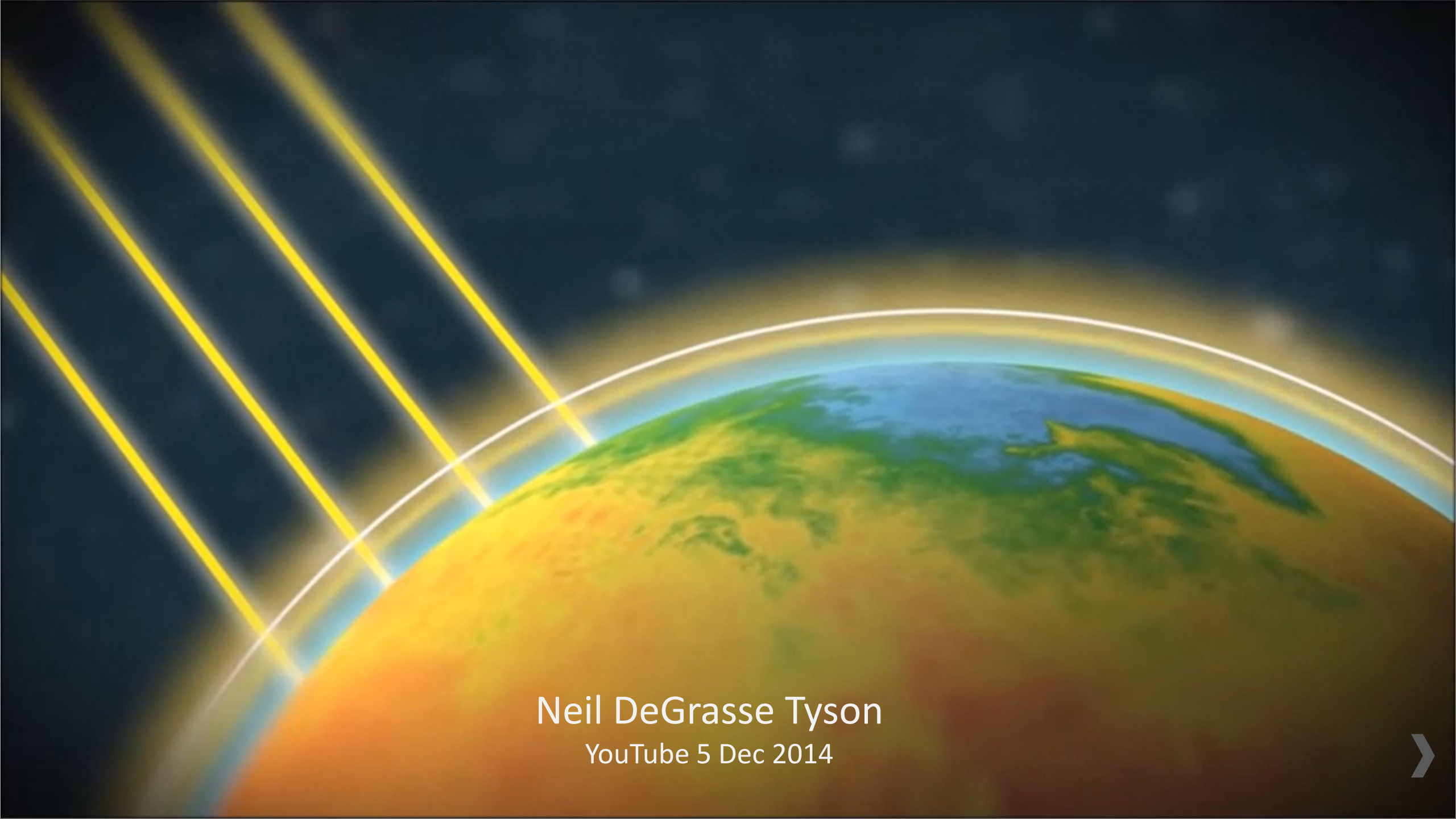
1. Building Blocks: Some important concepts
2. **Our Goldilocks Earth: a Radiative Balancing Act**
3. The Role of the Atmosphere: Greenhouse Gases & Clouds
4. Dynamics of the Earth System: Oceans, Atmosphere, Biosphere, Cryosphere, People, Plate Tectonics
5. Natural Variability of the Climate, short and long term. Ice Ages
6. Carbon Dioxide and other Greenhouse Gases: where do they come from, where do they go, how are they regulated?
7. Impacts and Future Projections for Global Warming -- Uncertainties
8. Adaptation and Amelioration Strategies. The Climate debate. Policy options.



Today's Discussion:
50 second Version

Neil DeGrasse Tyson
YouTube 5 Dec 2014





Neil DeGrasse Tyson
YouTube 5 Dec 2014



Our Goldilocks Earth



$$T_{\text{eff}} = 254 \text{ K } (-2^{\circ}\text{F})$$

$$T_{\text{surf}} = 288 \text{ K } (59^{\circ}\text{F}) +13\%$$

Global
Averages

Why?



Turn Off the Sun



$$T_{\text{eff}} = 254 \text{ K } (-2^{\circ}\text{F})$$

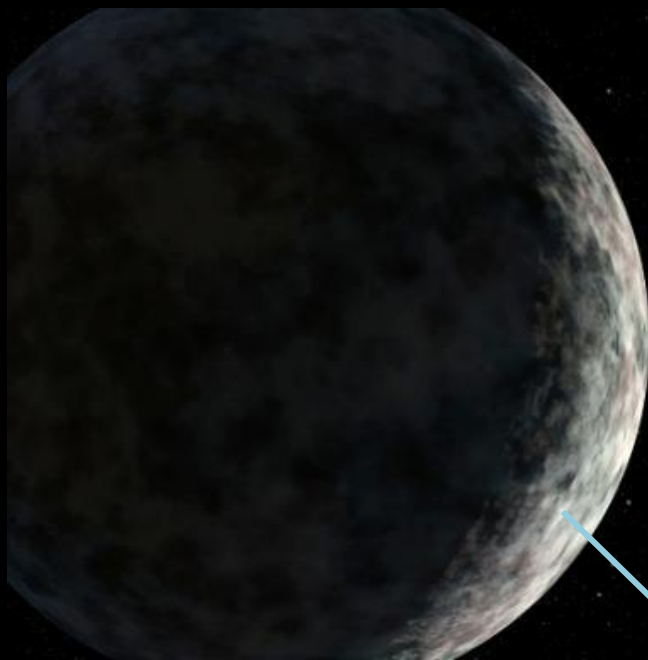
$$T_{\text{surf}} = 288 \text{ K } (59^{\circ}\text{F})$$



Turn Off the Sun



Turn Off the Sun



30 Ft thick coating of frozen air

$$T_{\text{eff}} = 36 \text{ K } (-395^\circ\text{F})$$

$$T_{\text{surf}} = 36 \text{ K } (-395^\circ\text{F})$$

Why does it only cool down to 36K? Why not colder?



Turn Off the Sun



$$T_{\text{eff}} = 36 \text{ K } (-395^\circ\text{F})$$

$$T_{\text{surf}} = 36 \text{ K } (-395^\circ\text{F})$$

30 Ft thick coating of frozen air

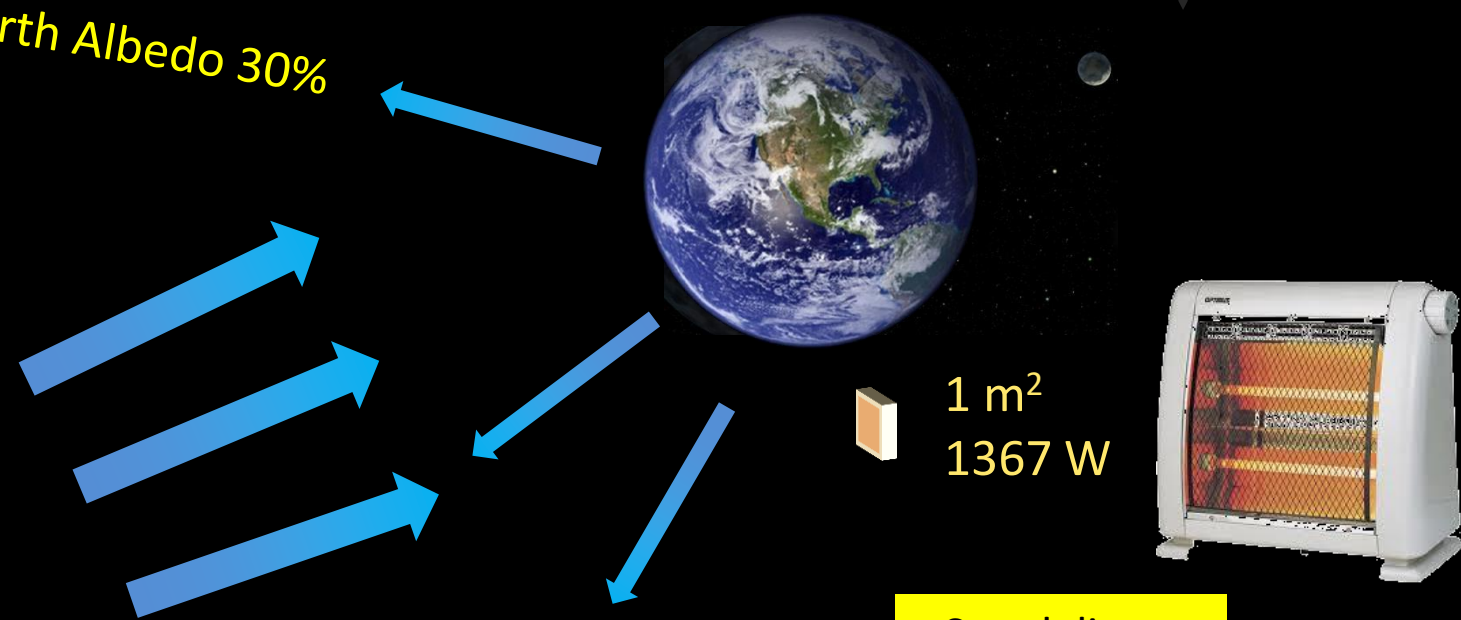
The core is still hot from earths creation, and from radioactive decay. This geothermal heat keeps the surface slightly warm



Turn the Sun Back On

30% of the sunlight hitting the earth is reflected back into space, so 70% is absorbed.

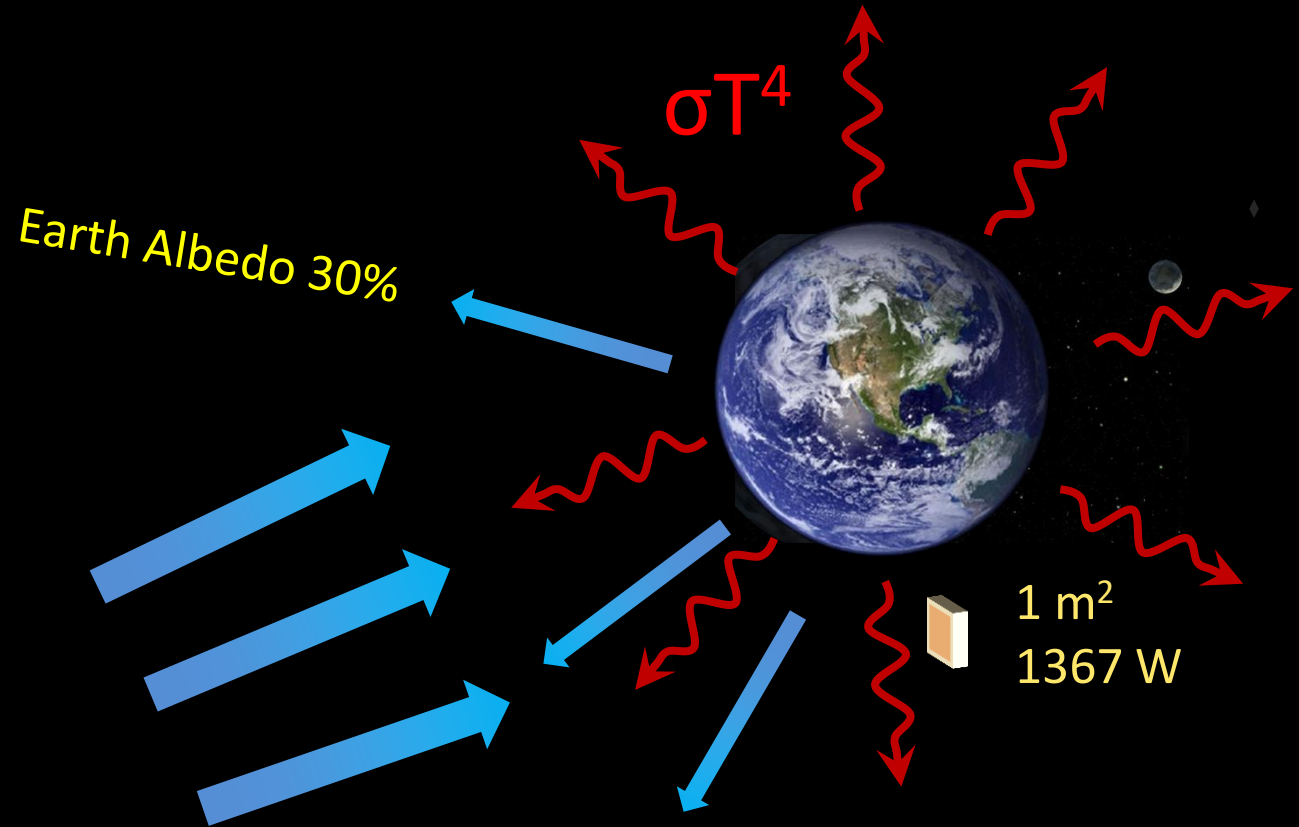
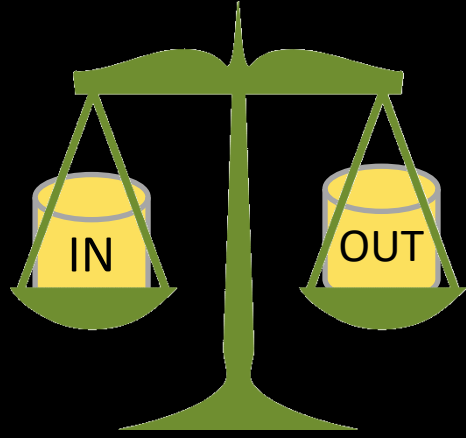
Earth Albedo 30%



1 m²
1367 W

Sun delivers 1367 watts per square meter in earth's vicinity

Sun is Back On

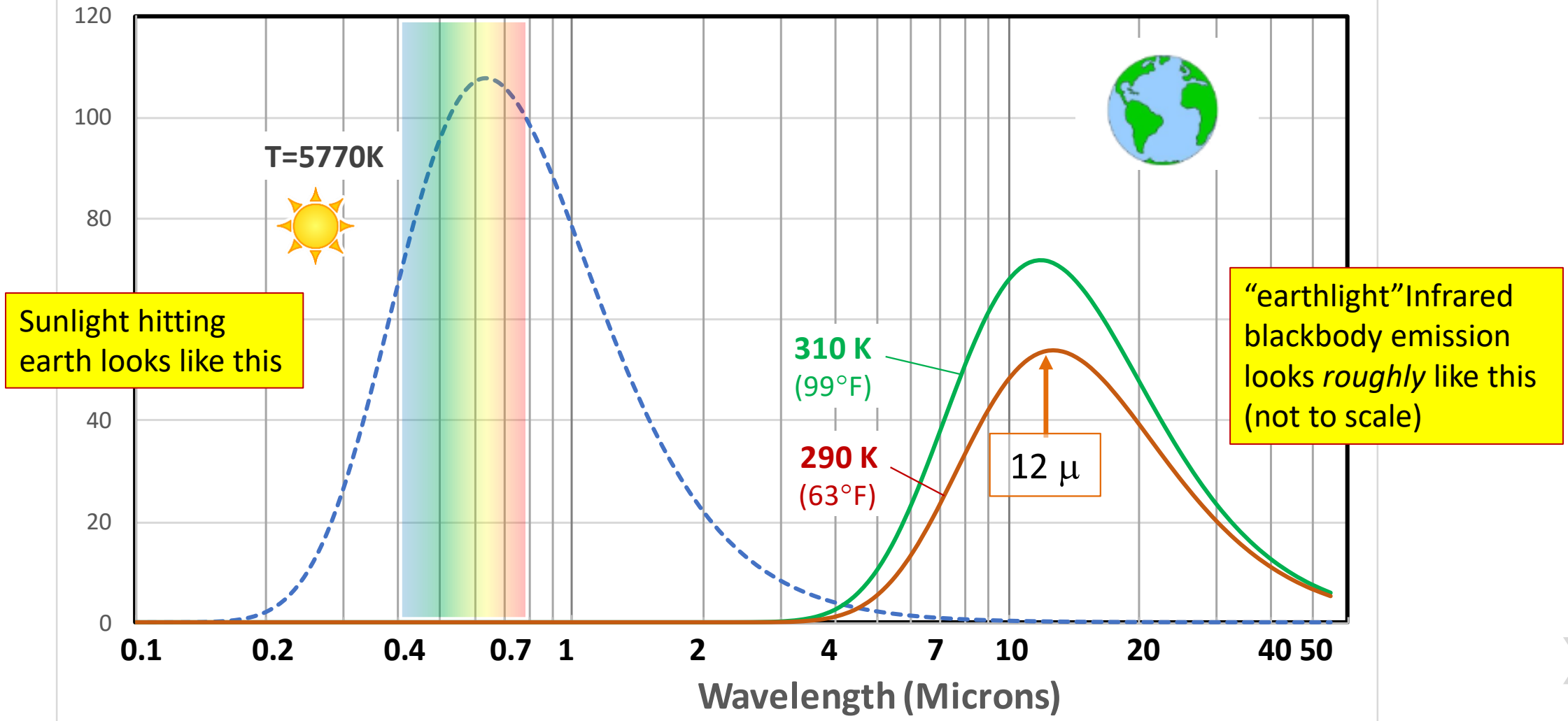


The incoming sun heat must be balanced approximately by Blackbody IR emission from the earth into space. That depends on T^4 ! So eventually earth's apparent temperature must adjust.

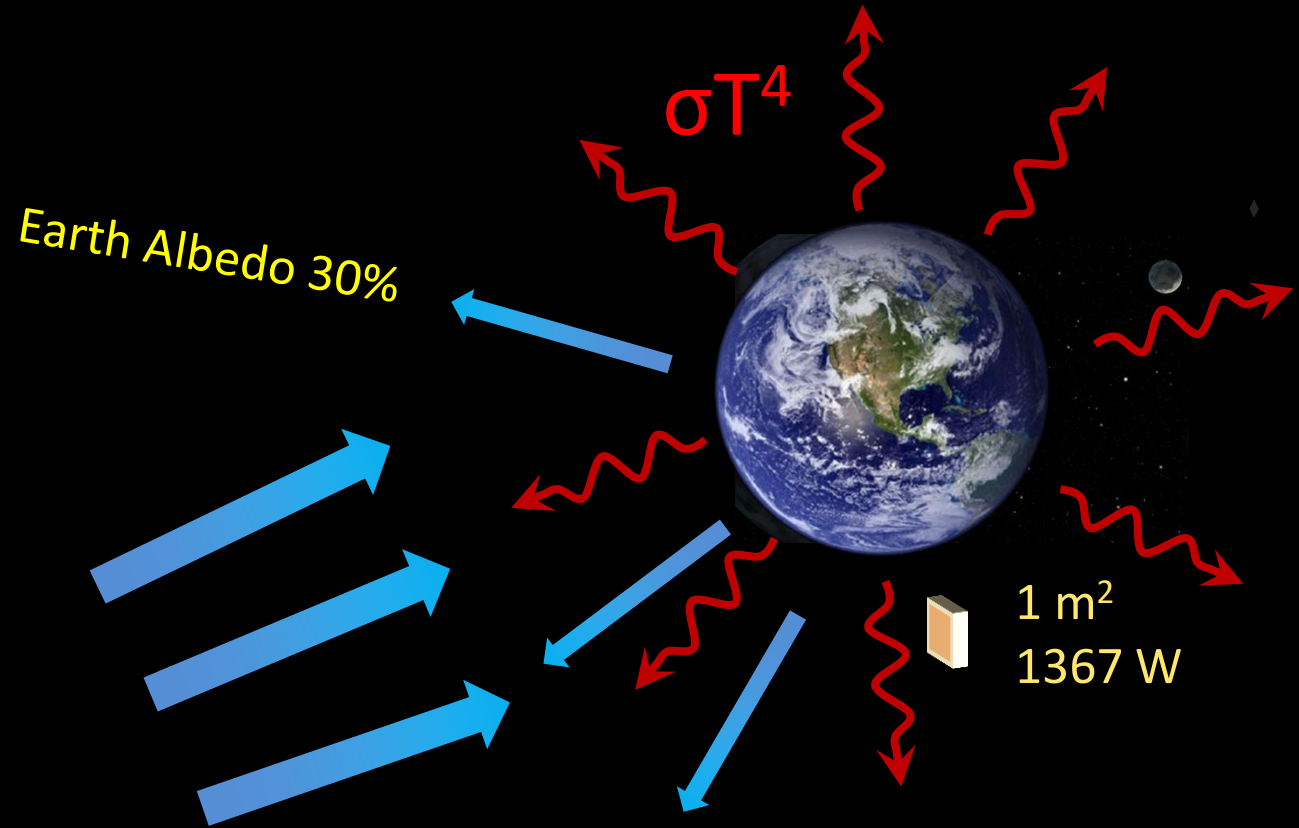
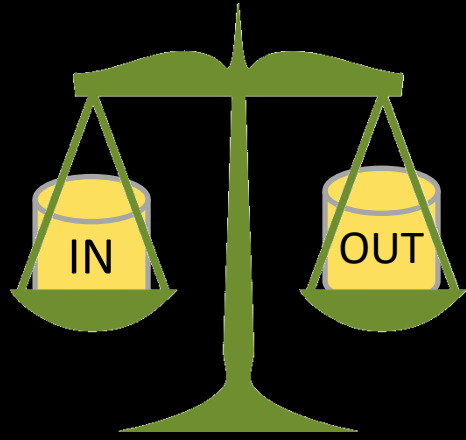


Solar Radiation vs. Earth Radiation

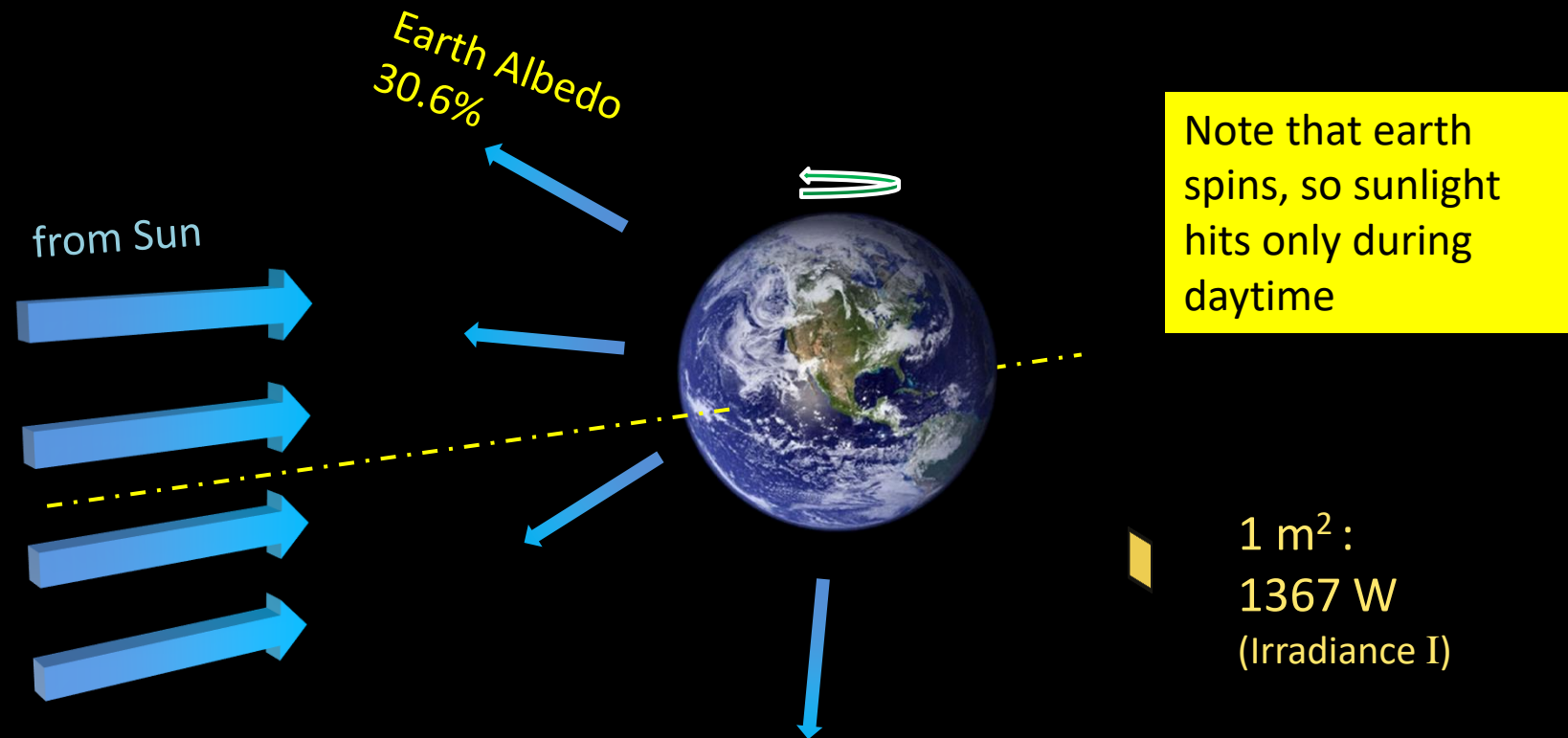
Blackbody Spectra: Sun vs Near Room Temperature



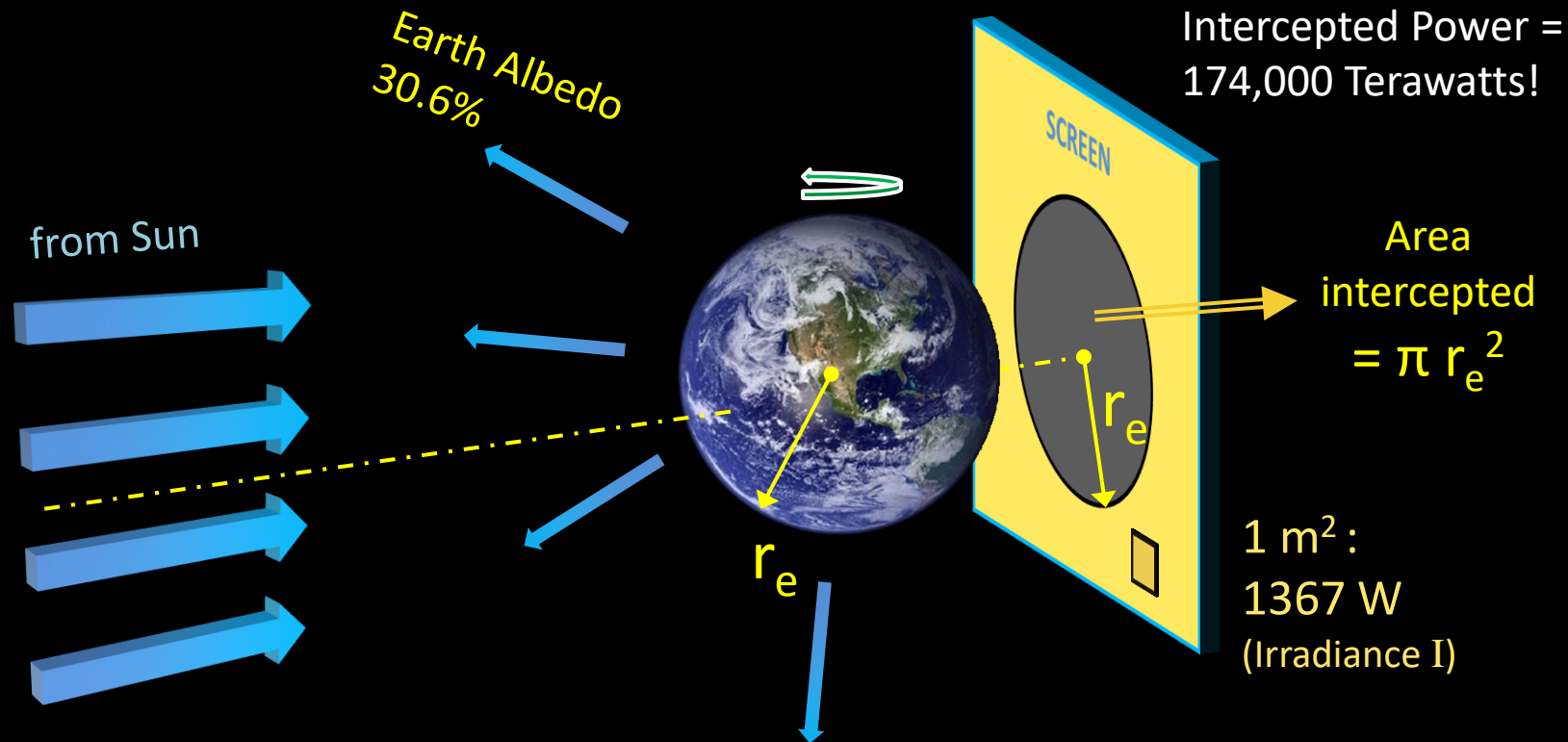
Turn the Sun Back On



How much sunlight does each square meter absorb?



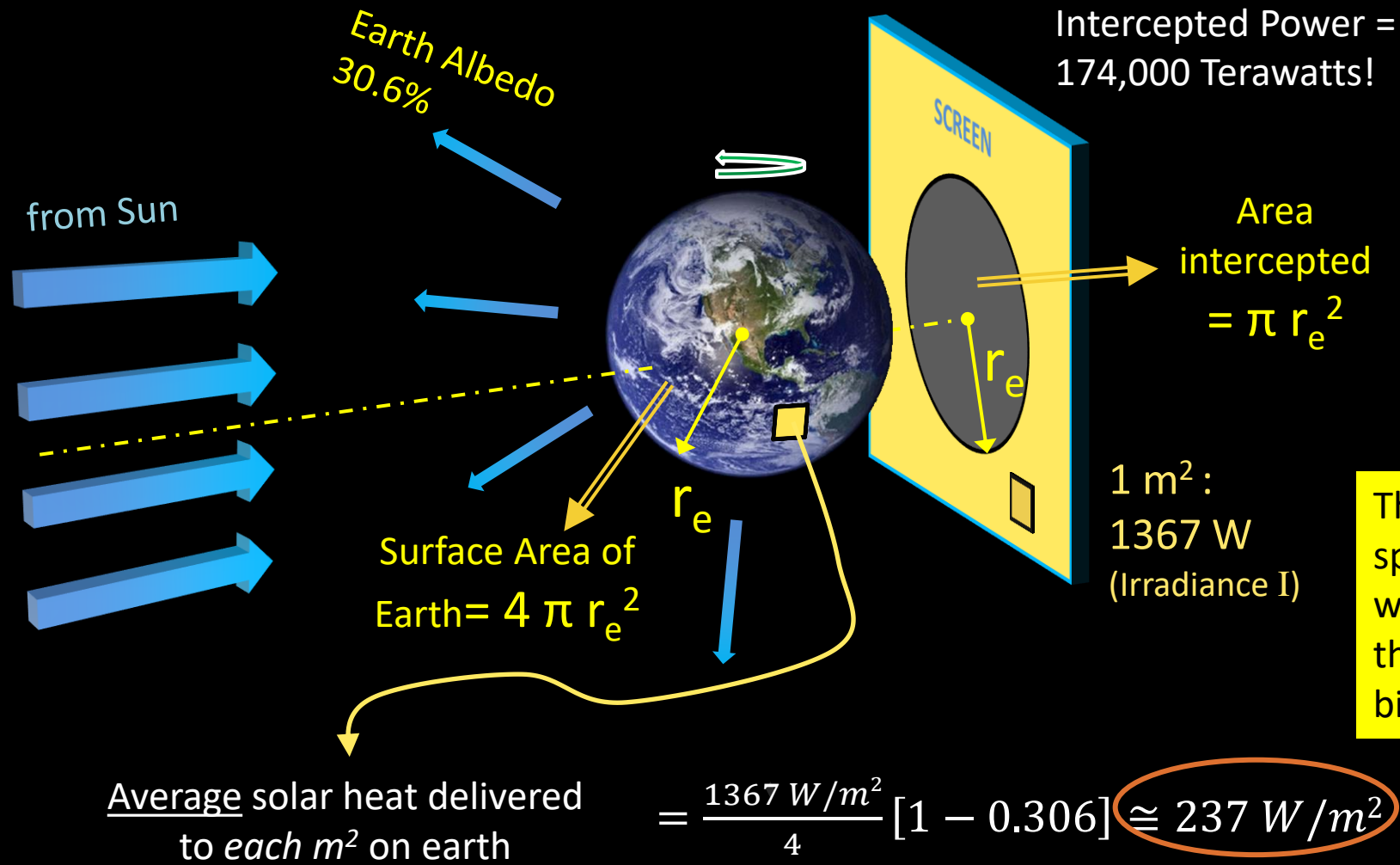
How much sunlight does each square meter absorb?



Credit for
Shadow Idea:
David Archer

An easy way to figure out how much heat is absorbed from the sunlight is to consider a screen just behind the earth. The shadow area directly gives us the amount of intercepted solar radiation. [But only 70% of this is actually absorbed]

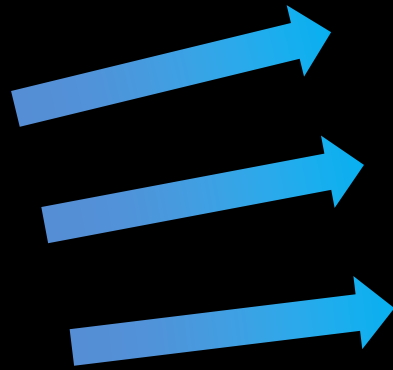
How much sunlight does each square meter absorb?



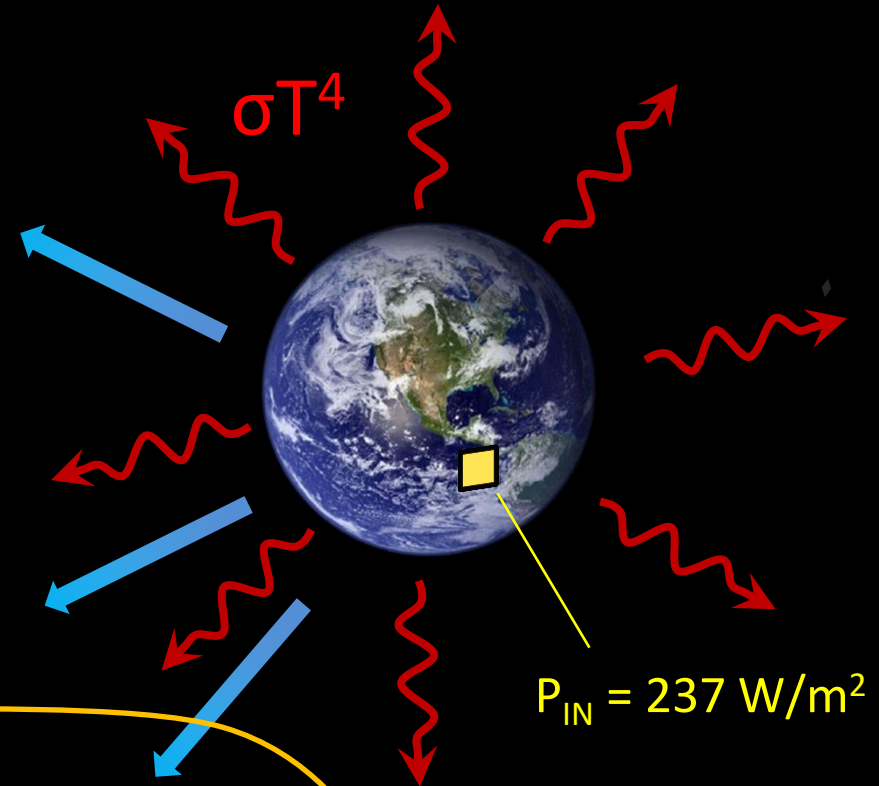
Credit for Shadow Idea: David Archer

This absorbed heat is spread out over the whole surface area of the Earth, which is 4x bigger than its shadow!

So What Effective Temperature Does This Imply?



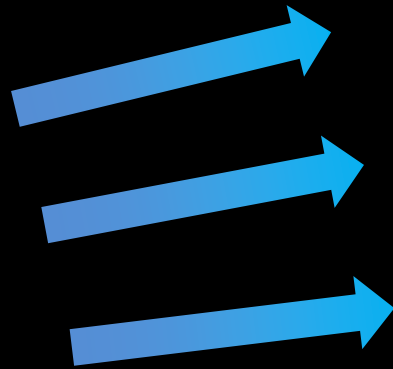
The required energy balance gives us a very simple equation for the needed effective Temperature.



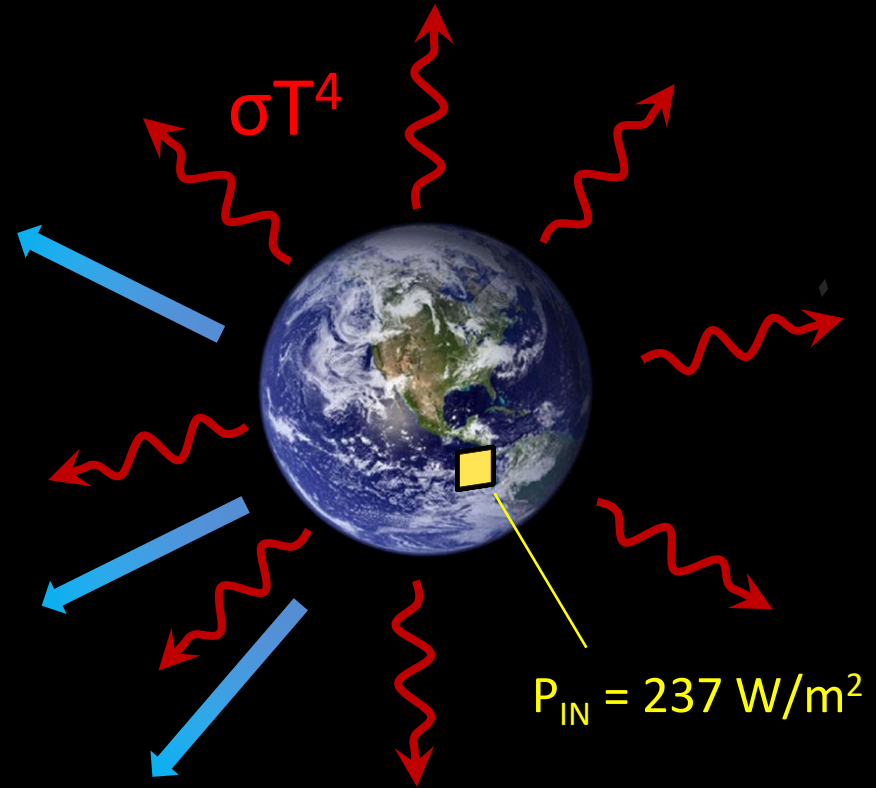
$$P_{\text{OUT}} = 237 \text{ W/m}^2 = \sigma T_{\text{EFF}}^4$$



So What Effective Temperature Does This Imply?



We can then easily calculate the effective Temperature as -2°F . This is the temperature seen from outside the earth – also known as “Skin Temperature”



$$P_{\text{IN}} = 237 \text{ W/m}^2$$

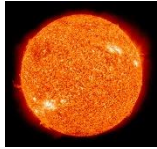
$$P_{\text{OUT}} = 237 \text{ W/m}^2 = \sigma T_{\text{EFF}}^4$$

$$T_{\text{EFF}} = \sqrt[4]{237/\sigma} = 254 \text{ }^{\circ}\text{K} \quad -2^{\circ}\text{F}$$

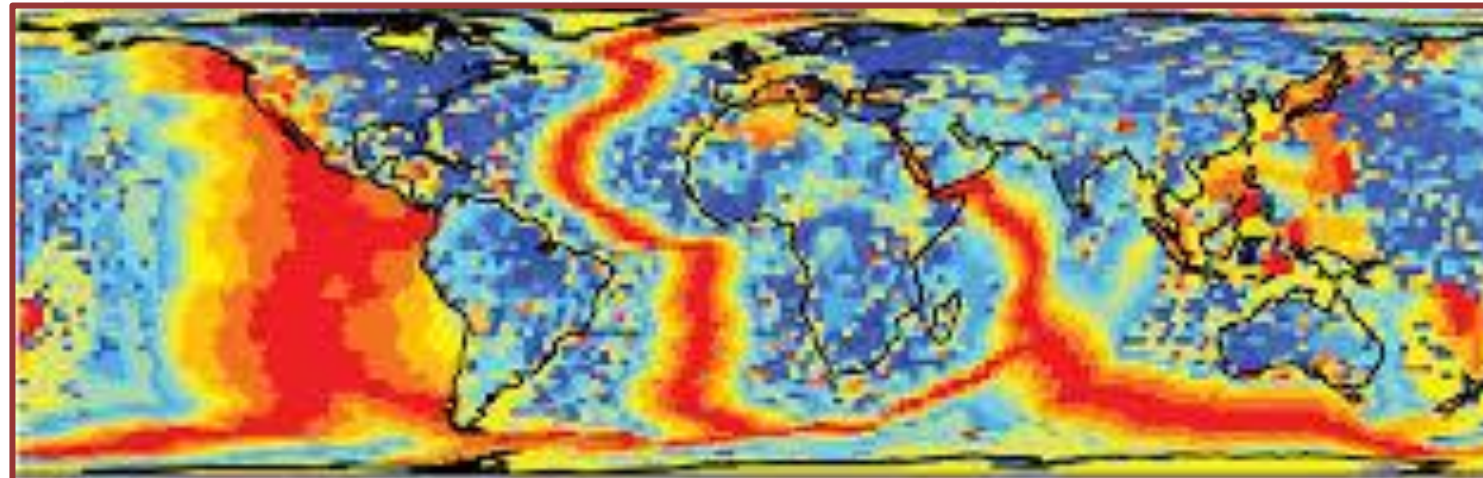
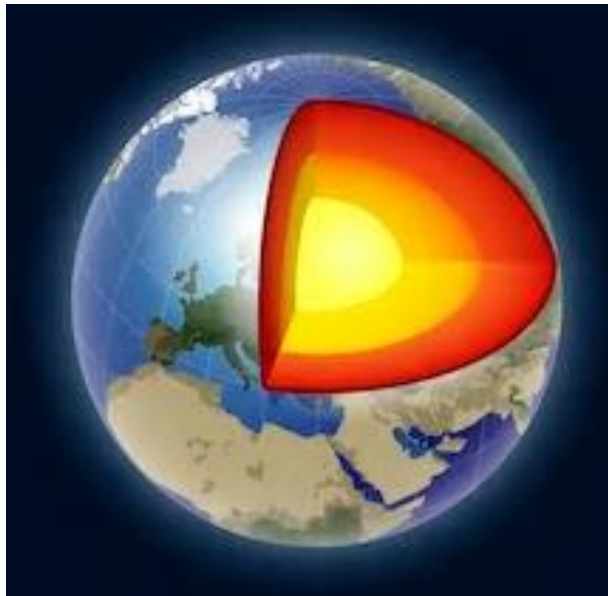
“Skin Temp.”



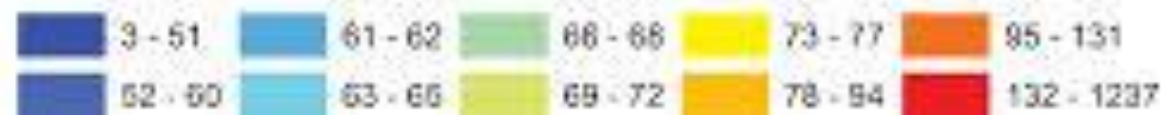
What Are the Heat Sources for Earth?



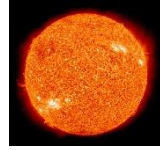
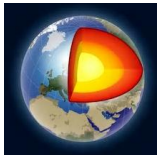
Source	Total Power	Power/sq m
	(TW)	W/m^2
Sun	120,900	237
Geothermal	45	0.088



Final Estimate of Heat Flow ($mW m^{-2}$) (Area-weighted Mean)



What Are the Heat Sources for Earth?

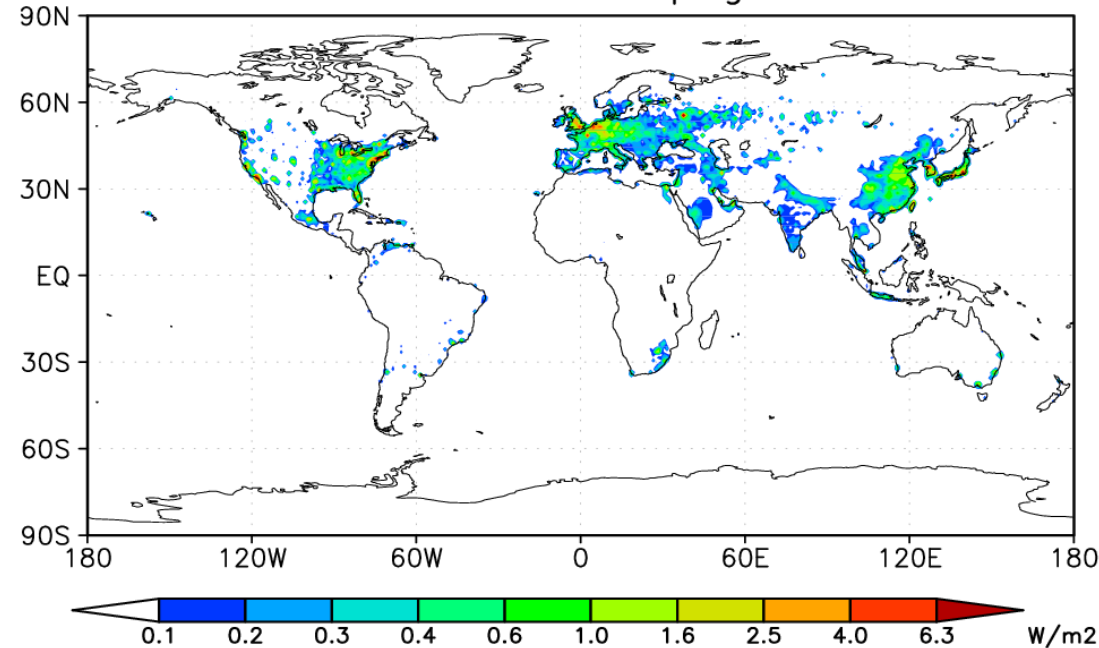


Source	Total Power (TW)	Power/sq m W/m^2
Sun	120,900	237
Geothermal	45	0.088
Human Activity	≈ 17	0.033

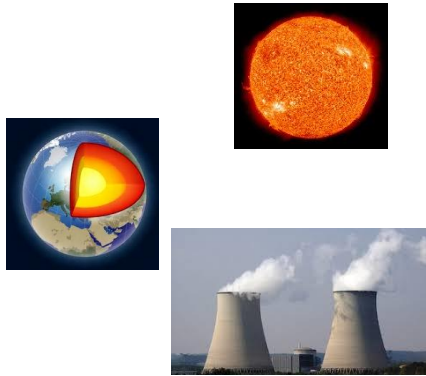
This is the *directly* generated human heat (e.g. nuclear power plants, burning fuels, etc.), **not** the indirect effect of CO2 emissions, etc.



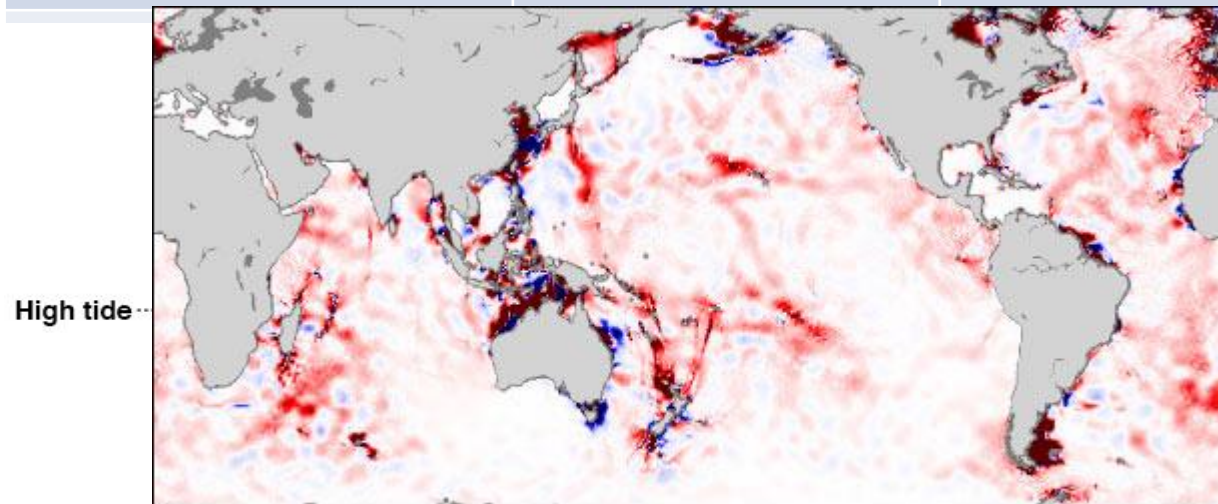
2005 Annual-Mean Anthropogenic Heat Flux



What Are the Heat Sources for Earth?



Source	Total Power	Power/sq m
	(TW)	W/m^2
Sun	120,900	237
Geothermal	45	0.088
Human Activity	≈ 17	0.033
Tides	≈ 4	0.008

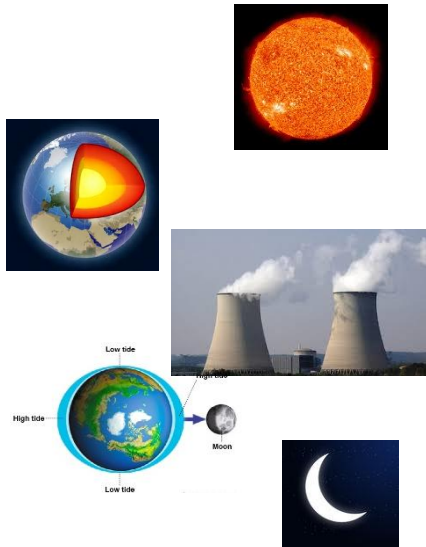


Tidal Energy Dissipation (mW/m^2)

-30 -20 -10 0 10 20 30



What Are the Heat Sources for Earth?

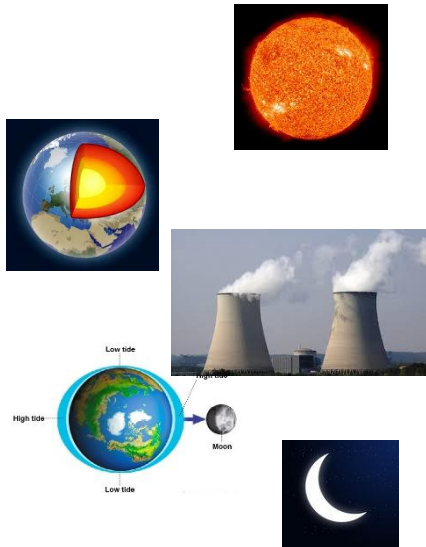


Source	Total Power	Power/sq m
	(TW)	W/m^2
Sun	120,900	237
Geothermal	45	0.088
Human Activity	≈ 17	0.033
Tides	≈ 4	0.008
Moonlight*	≈ 1.8	0.0035



* Includes IR emission from hot face!

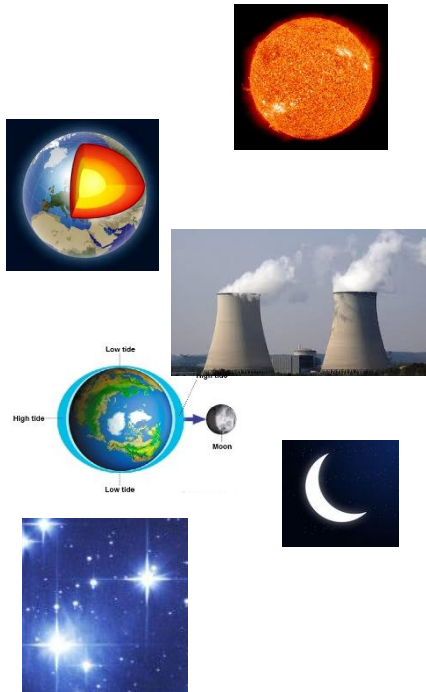
What Are the Heat Sources for Earth?



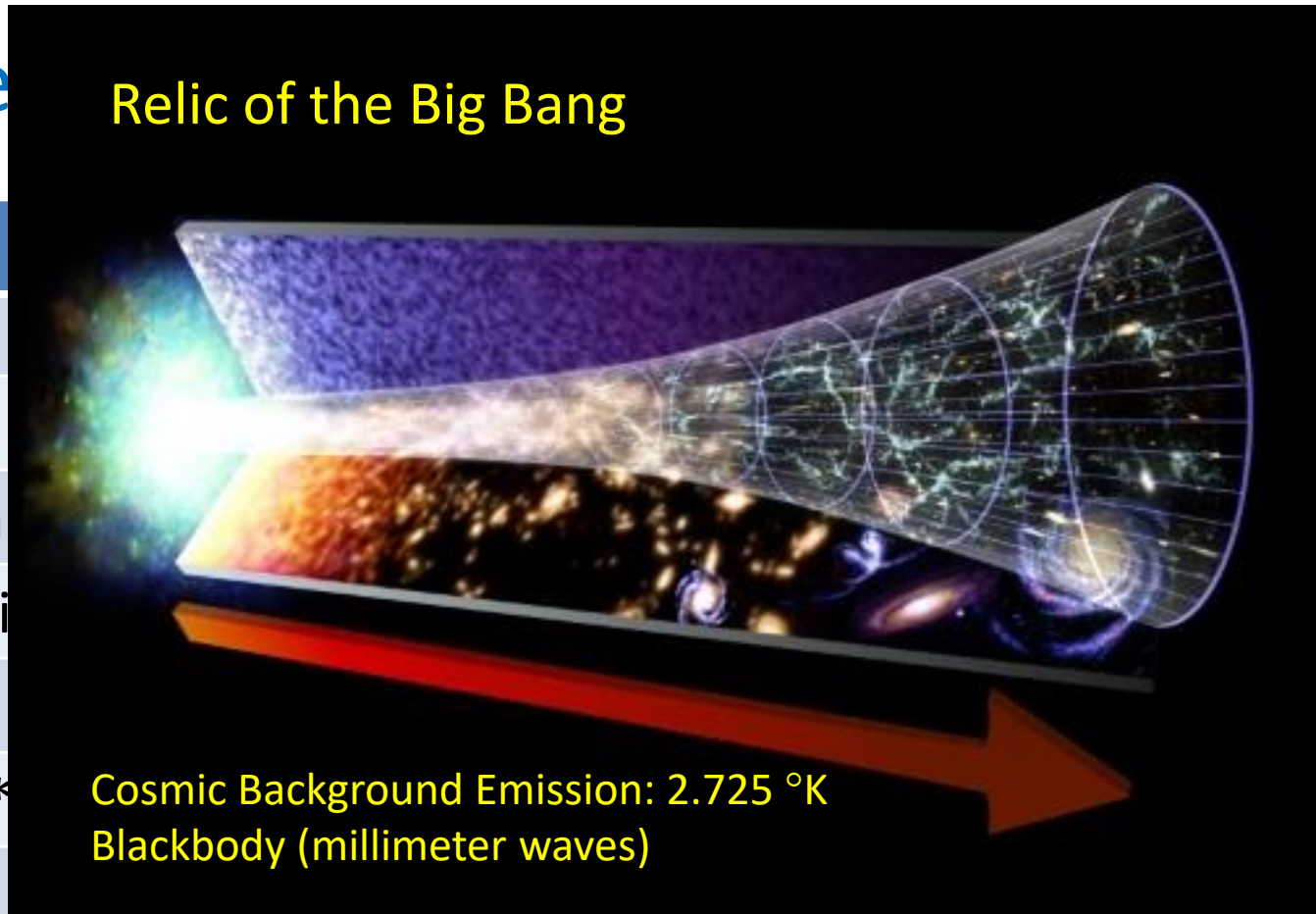
Source	Total Power	Power/sq m
	(TW)	W/m^2
Sun	120,900	237
Geothermal	45	0.088
Human Activity	≈ 17	0.033
Tides	≈ 4	0.008
Moonlight*	≈ 1.8	0.0035
Starlight	≈ 0.005	0.00001



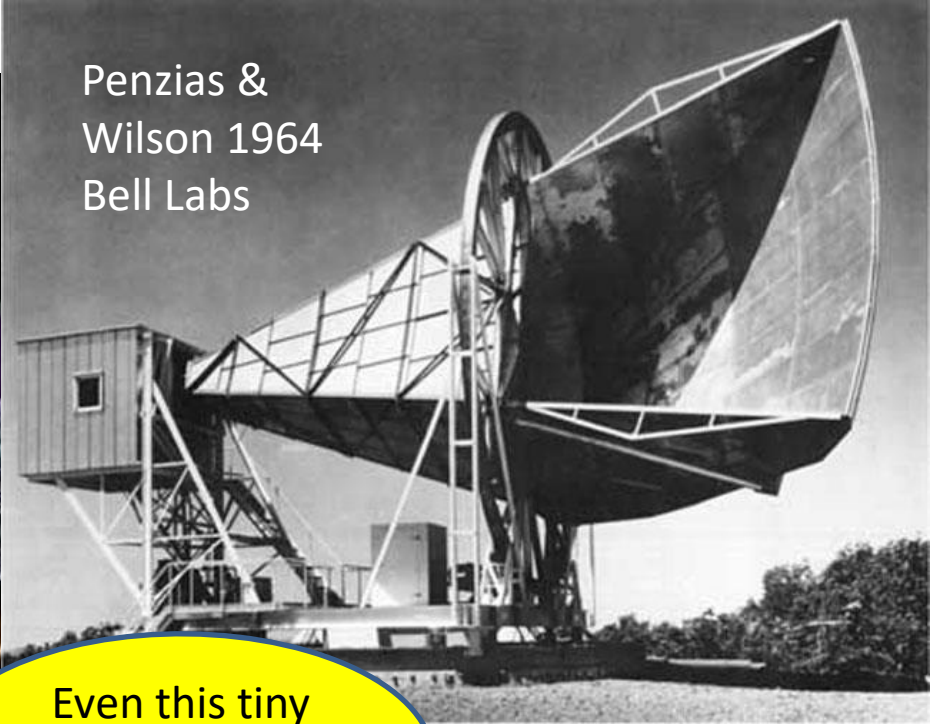
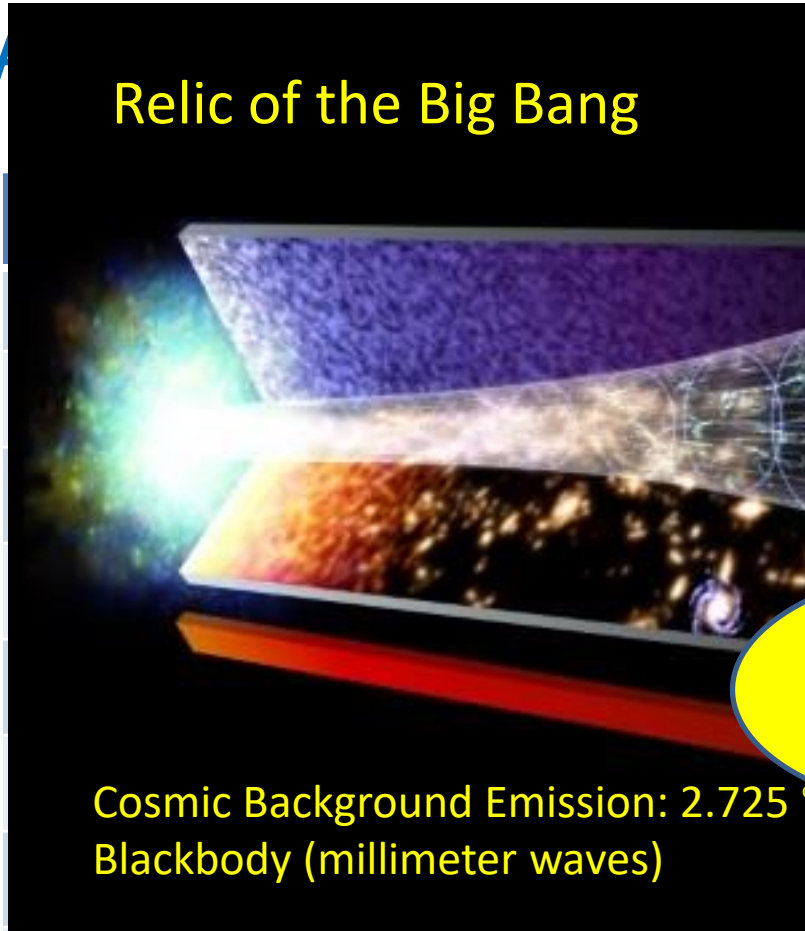
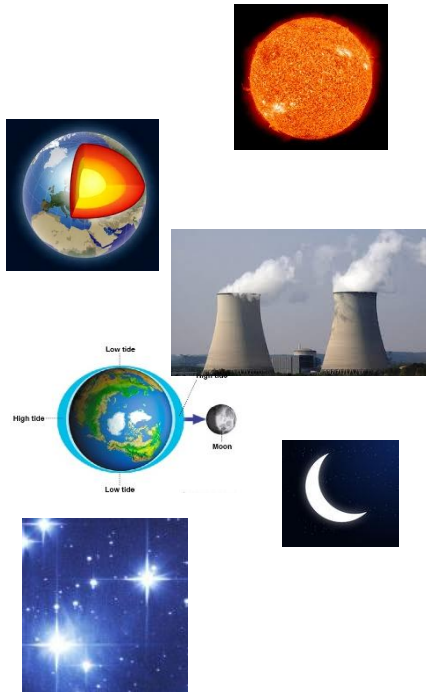
What Are the He



Source		
Sun		
Geothermal		
Human Activity		
Tides		
Moonlight*		
Starlight		
CBE	0.0016	0.0000031
Totals	121,000	237.13



What A



Even this tiny amount = 1.6 Gigawatts!

Cosmic Background Emission: 2.725 °K
Blackbody (millimeter waves)

CBE	0.0016	0.0000031
Totals	121,000	237.13

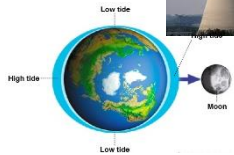
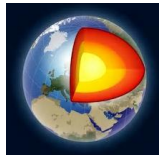
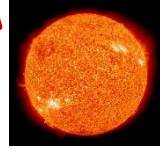


What Are the Heat Sources for Earth?



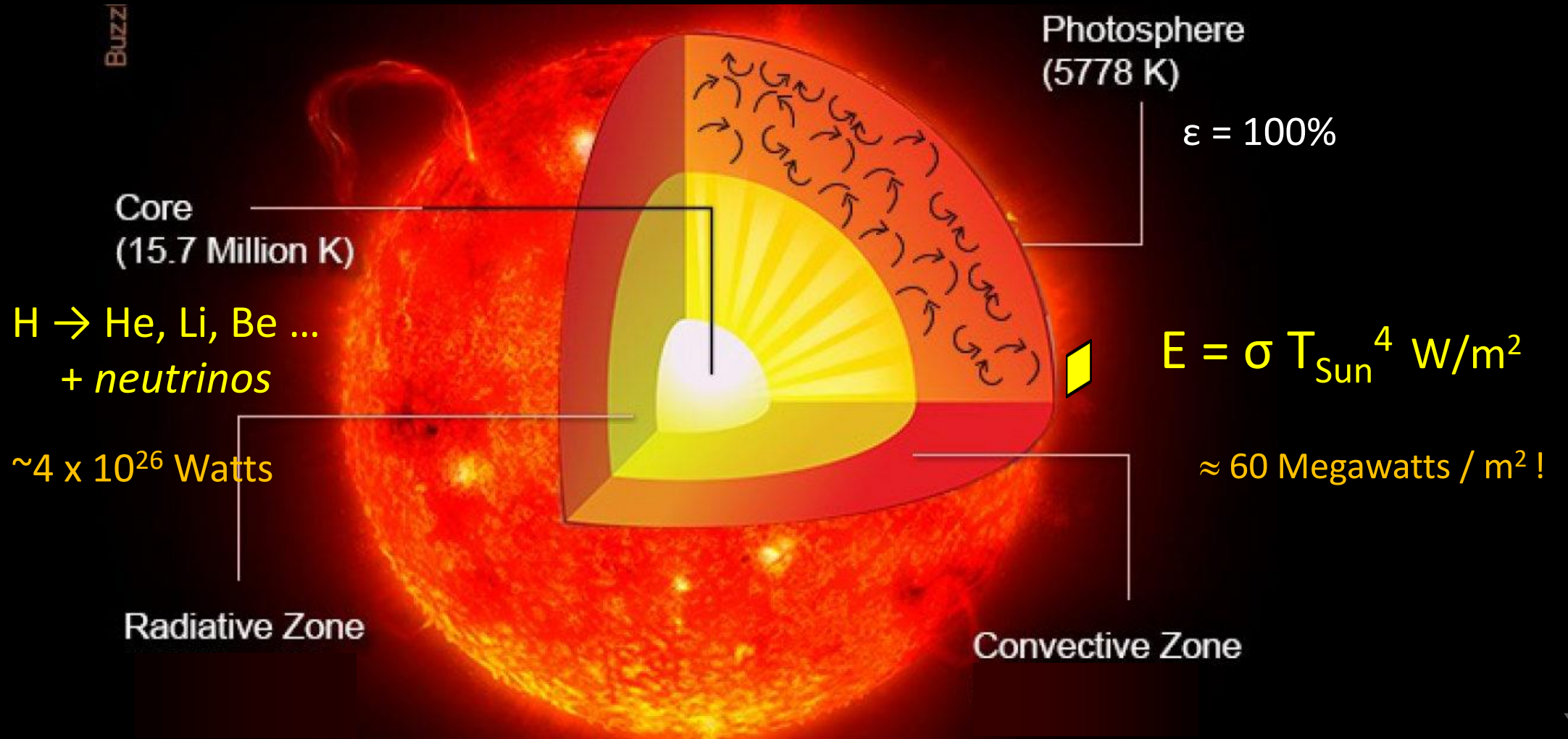
What Are the Heat Sources for Earth?

The sun is virtually the whole story...



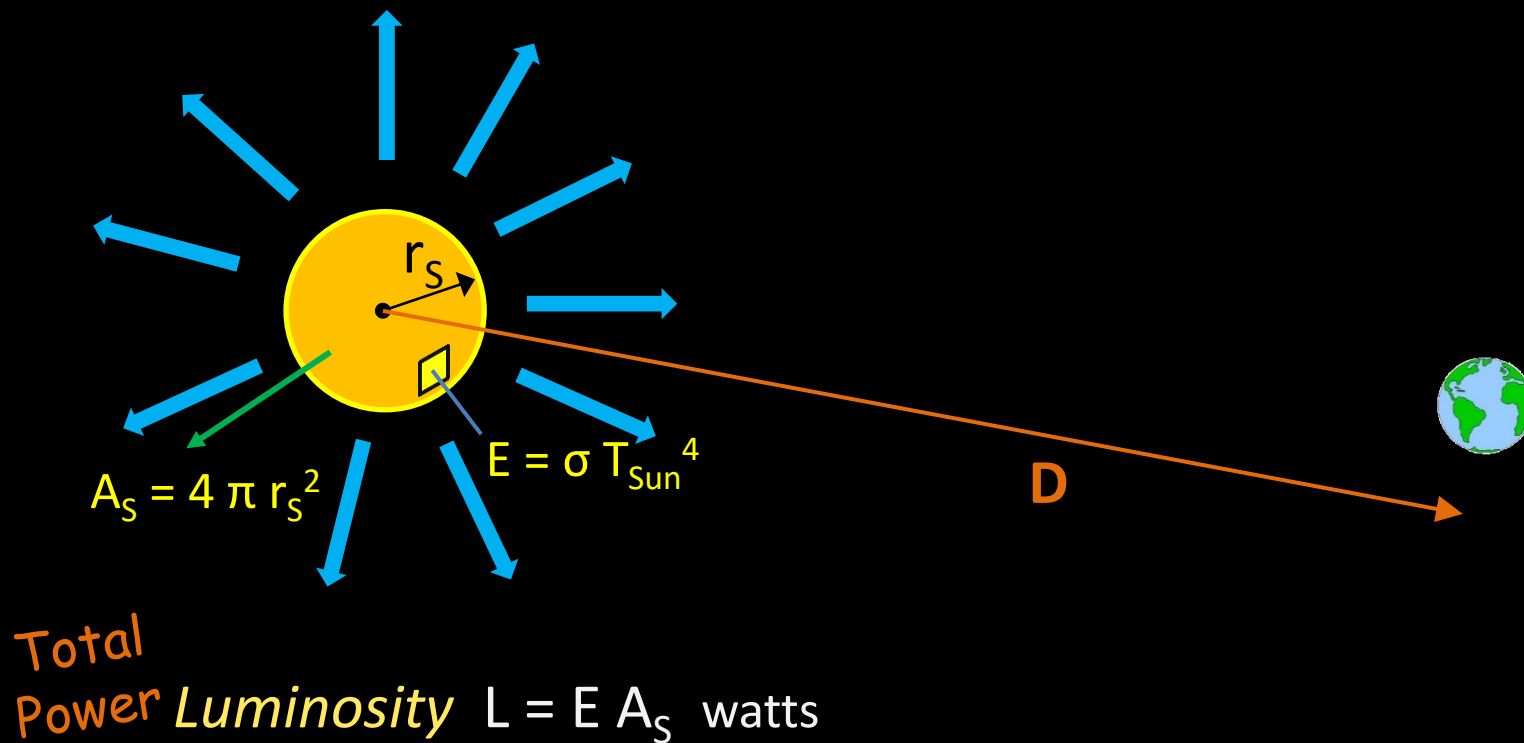
Source	Total Power	Power/sq m
	(TW)	W/m^2
Sun	120,900	237
Geothermal	45	0.088
Human Activity	≈ 17	0.033
Tides	≈ 4	0.008
Moonlight*	≈ 1.8	0.0035
Starlight	≈ 0.005	0.00001
CBE	0.0016	0.0000031
Totals	121,000	237.13

Our Sun – A fusion reactor



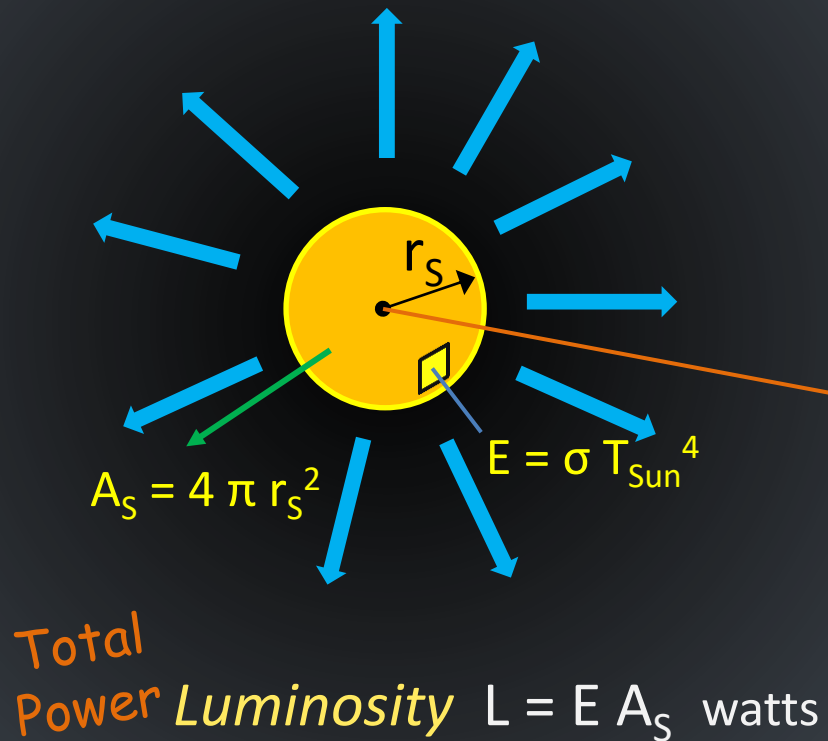
How Much Sunlight Reaches Earth?

Start by figuring out the total energy output of the sun, called its Luminosity....



How Much Sunlight Reaches Earth?

Then spread this heat out over a huge spherical shell at the distance of the Earth. Doing the accounting, get 1367 Watts per square meter near the earth...



$$A_{\text{bubble}} = 4 \pi D^2$$

Light hitting each square meter =

Irradiance

$$I = L / A_{\text{bubble}}$$

watts/m²

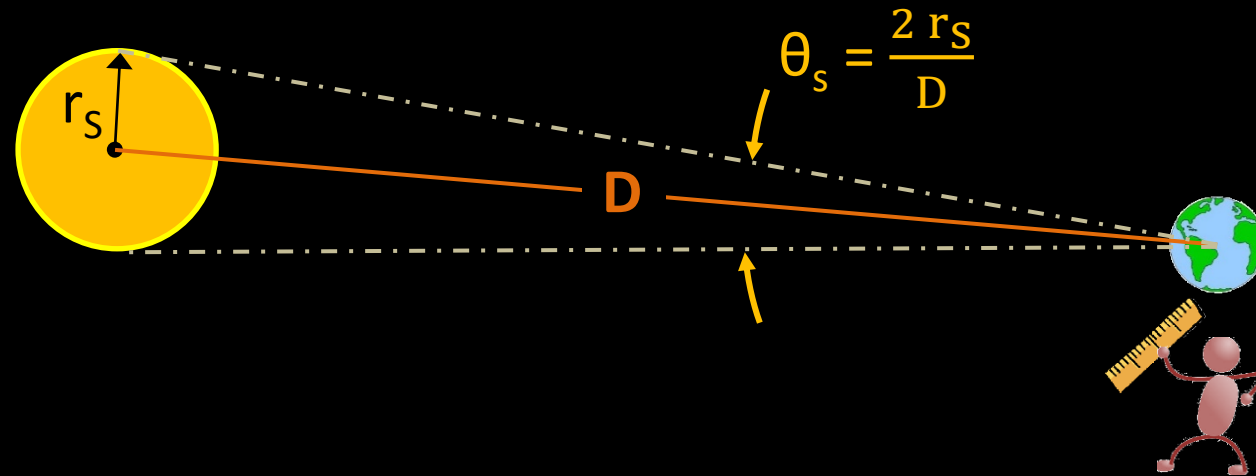
$$I = \sigma T_{\text{Sun}}^4 \left[\frac{r_s}{D} \right]^2 = 1367 \text{ W/m}^2$$

How Much Sunlight Reaches Earth?

Curiously, we don't actually have to know how big the sun is, or how far away it is. All we need to do is measure its apparent angular size, which anyone can do. It is about 0.5 degrees, same as the moon.

$$I = \sigma T_{\text{Sun}}^4 \left[\frac{r_s}{D} \right]^2$$

Light hitting each square meter =
Irradiance
at Earth
 $\cong 1367 \text{ W/m}^2$



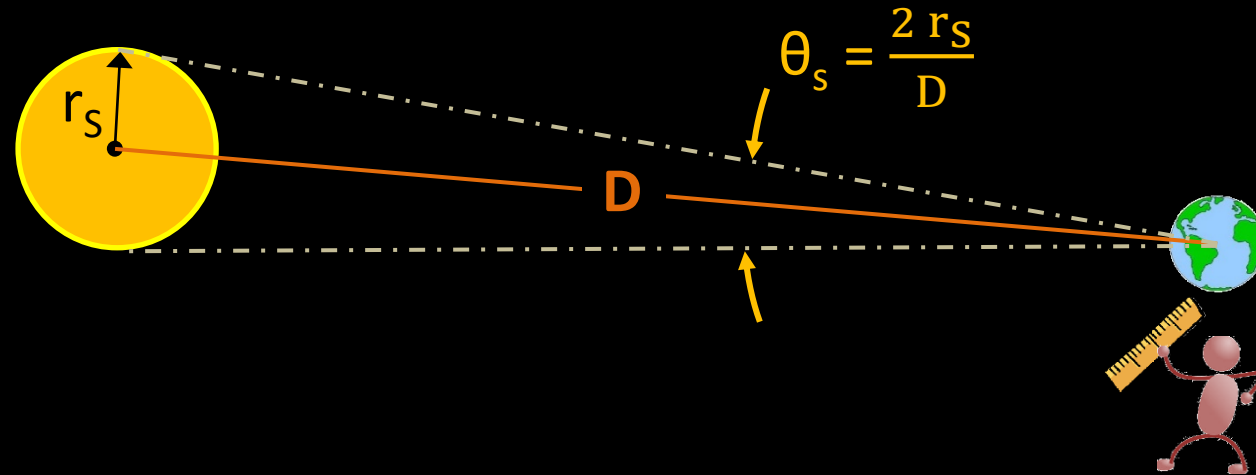
How Much Sunlight Reaches Earth?

So we can get a formula for the Irradiance of the sun at any location just in terms of the Sun's temperature and apparent angular size θ_s . We do not need to know how far away the sun is...

$$I = \sigma T_{\text{Sun}}^4 \left[\frac{r_s}{D} \right]^2$$

$$I = \sigma T_{\text{Sun}}^4 [\theta_s / 2]^2 \cong 1367 \text{ W/m}^2$$

Light hitting each square meter=
Irradiance
at Earth



Take a deep breath...

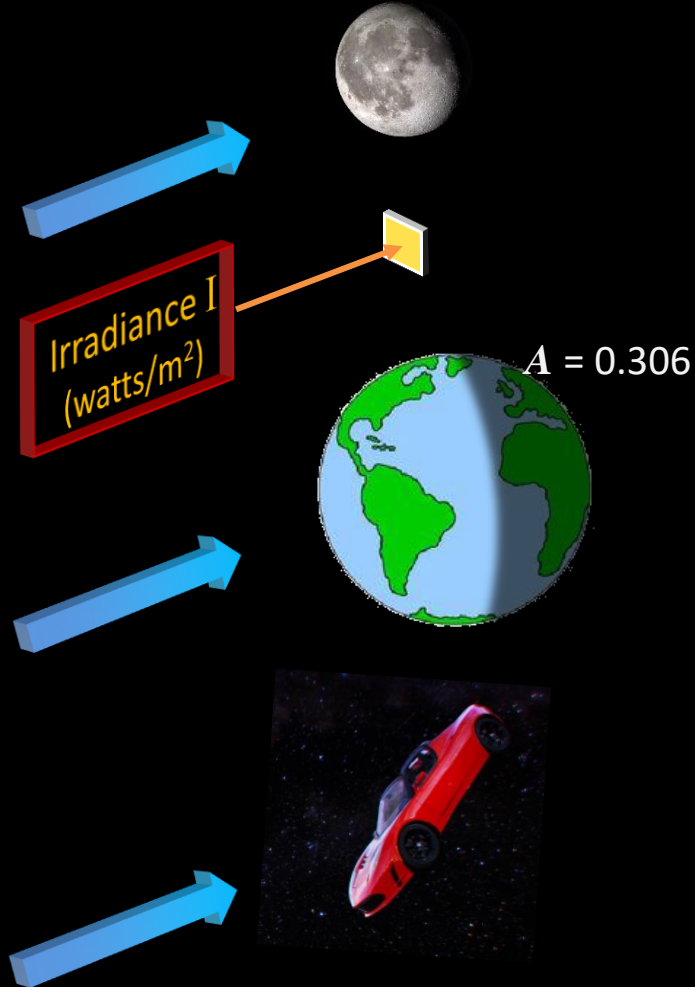


A bit complicated, but definitely not Rocket Science!

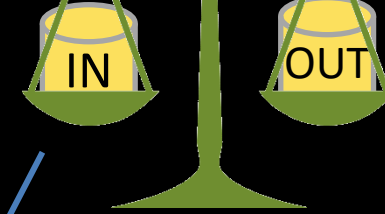


Equilibrium Effective Temperature for Objects in Space

- “Bond” Albedo A
- IR Emissivity $\epsilon \approx 1$



Radiometric Balance



For each
square meter
of surface

$$\frac{I}{4} [1 - A]$$

$$I = \sigma T_{\text{Sun}}^4 [\theta_s / 2]^2$$

Equilibrium Effective Temperature

Radiometric Balance

Albedo*:

Fraction of sunlight *reflected* by planet
(and therefore not absorbed)

For each
square meter
of surface

Irradiance
(watts/

1.0

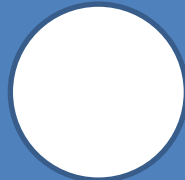
.77

.30

.25

.11

0



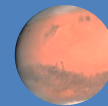
No
Absorption



Venus



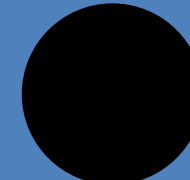
Earth



Mars



Moon

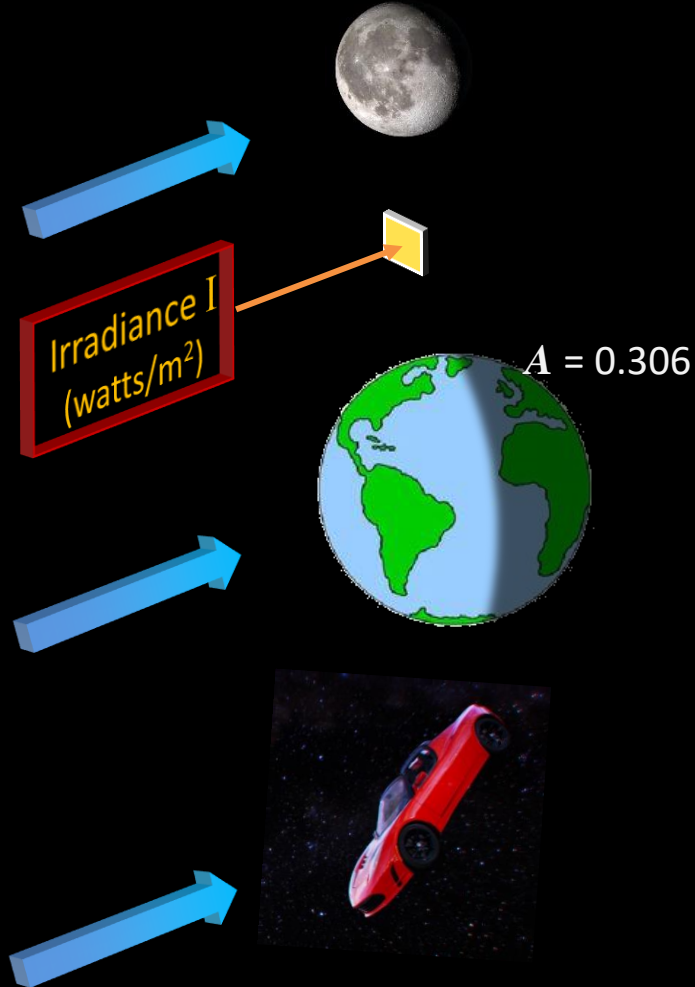


100%
Absorption

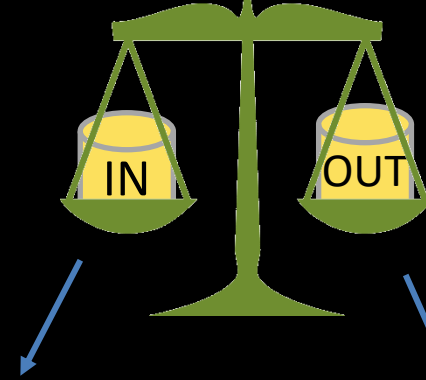
* Bond Albedo

Equilibrium Effective Temperature

- “Bond” Albedo A
- IR Emissivity $\epsilon \approx 1$



Radiometric Balance



For each square meter of surface

$$\frac{I}{4} [1 - A] = \sigma T_{eff}^4$$

$$T_{eff} = \frac{1}{2} T_{sun} \sqrt{\theta_s} \sqrt[4]{(1 - A)}$$

$$I = \sigma T_{Sun}^4 [\theta_s / 2]^2$$

So we get a fairly simple formula for figuring out the T_{EFF} of bodies knowing only the Sun's temperature, its apparent angular size, and the body's Albedo!

Equilibrium Effective Temperature

$$T_{eff} = \frac{1}{2} T_{sun} \sqrt{\theta_s} \sqrt[4]{(1 - A)}$$

“Skin Temperature”

We can easily calculate T_{EFF} for a bunch of heavenly bodies, and compare with NASA's numbers....



Bare Rocks:
(No atmosphere)

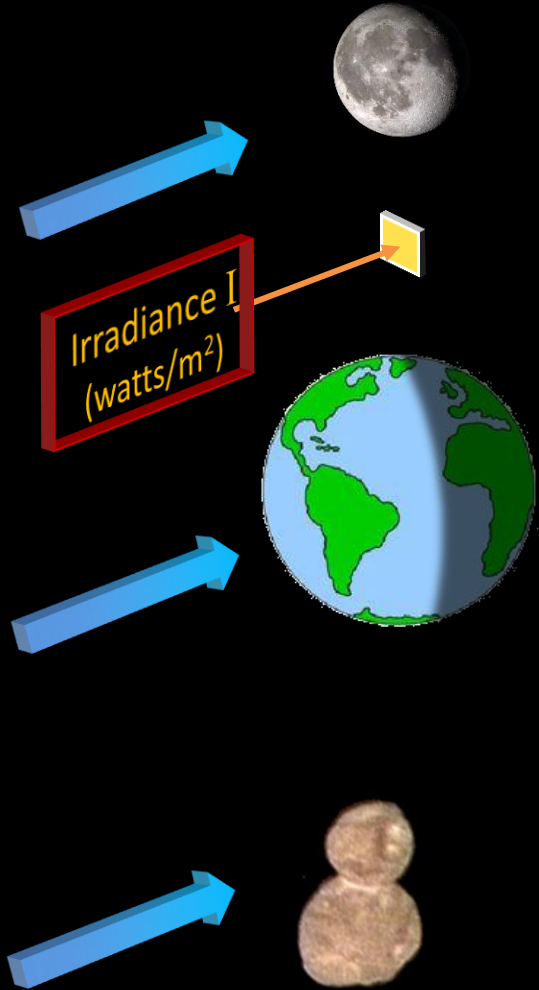
Average Surface Temperature T_{SURF}
is the same as T_{eff}



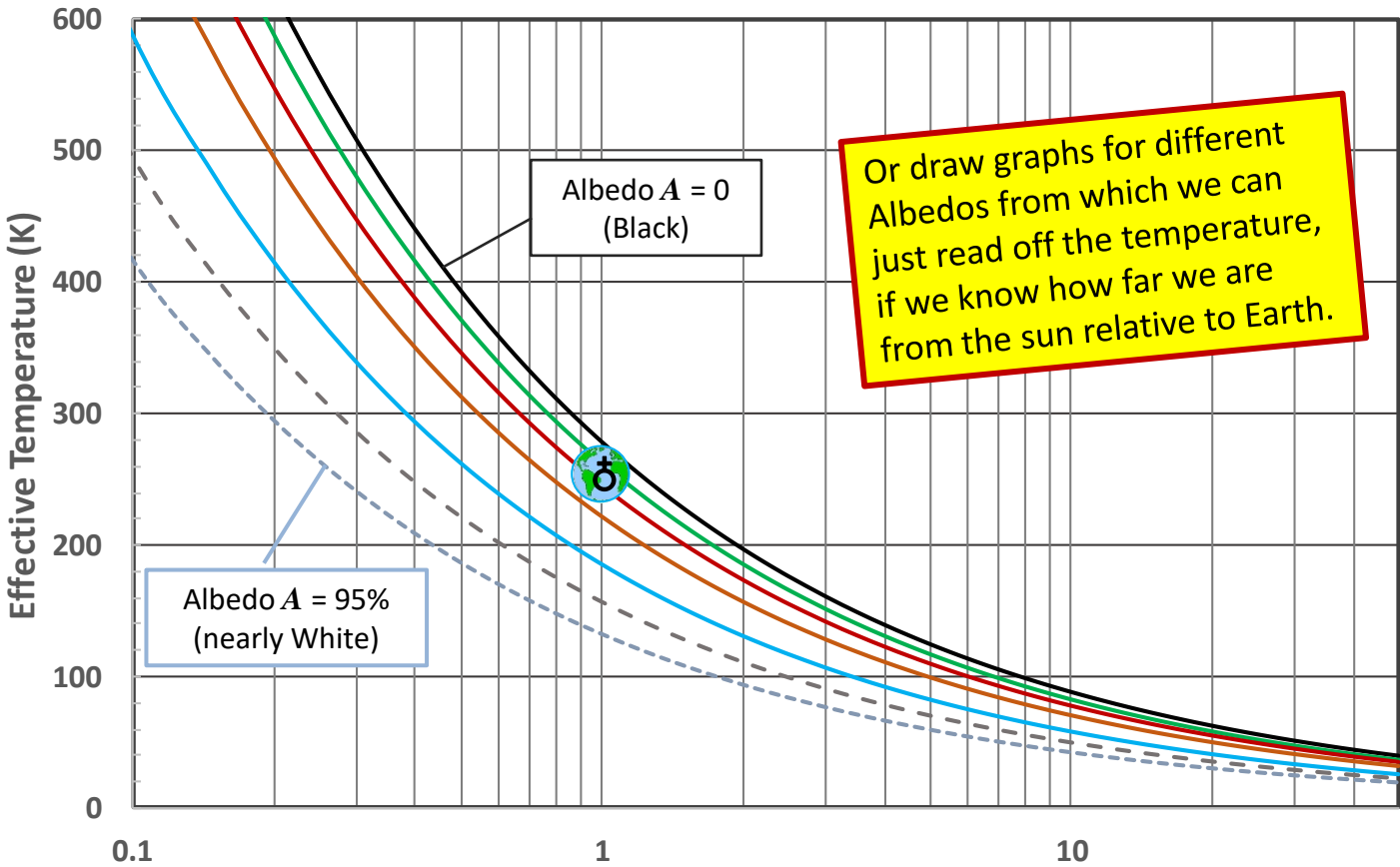
Body	T_{SUN}	θ_s Apparent Sun Dia.	D_{rel} Relative Distance	A^* Bond Albedo	T_{EFF}		
					Calculated		NASA*
Units	$^{\circ}K$	deg	A.U.		$^{\circ}K$	$^{\circ}F$	$^{\circ}K$
Mercury	5770	1.377 $^{\circ}$	0.39	0.068	439	331	439.6
Venus	5770	0.737 $^{\circ}$	0.72	0.77	227	-52	226.6
Earth	5770	0.533 $^{\circ}$	1.0	0.306	254	-2	254.0
Moon	5770	0.533 $^{\circ}$	1.0	0.11	270	27	270.4
Mars	5770	0.350 $^{\circ}$	1.52	0.25	210	-82	209.8
Uranus	5770	0.028 $^{\circ}$	19.2	0.30	58	-355	58.1
Ultima-Thule	5770	0.012 $^{\circ}$	44.5	0.09 (est)	41	-386	?

Equilibrium Effective Temperature

$$T_{eff} = \frac{1}{2} T_{sun} \sqrt{\theta_s} \sqrt[4]{(1 - A)}$$



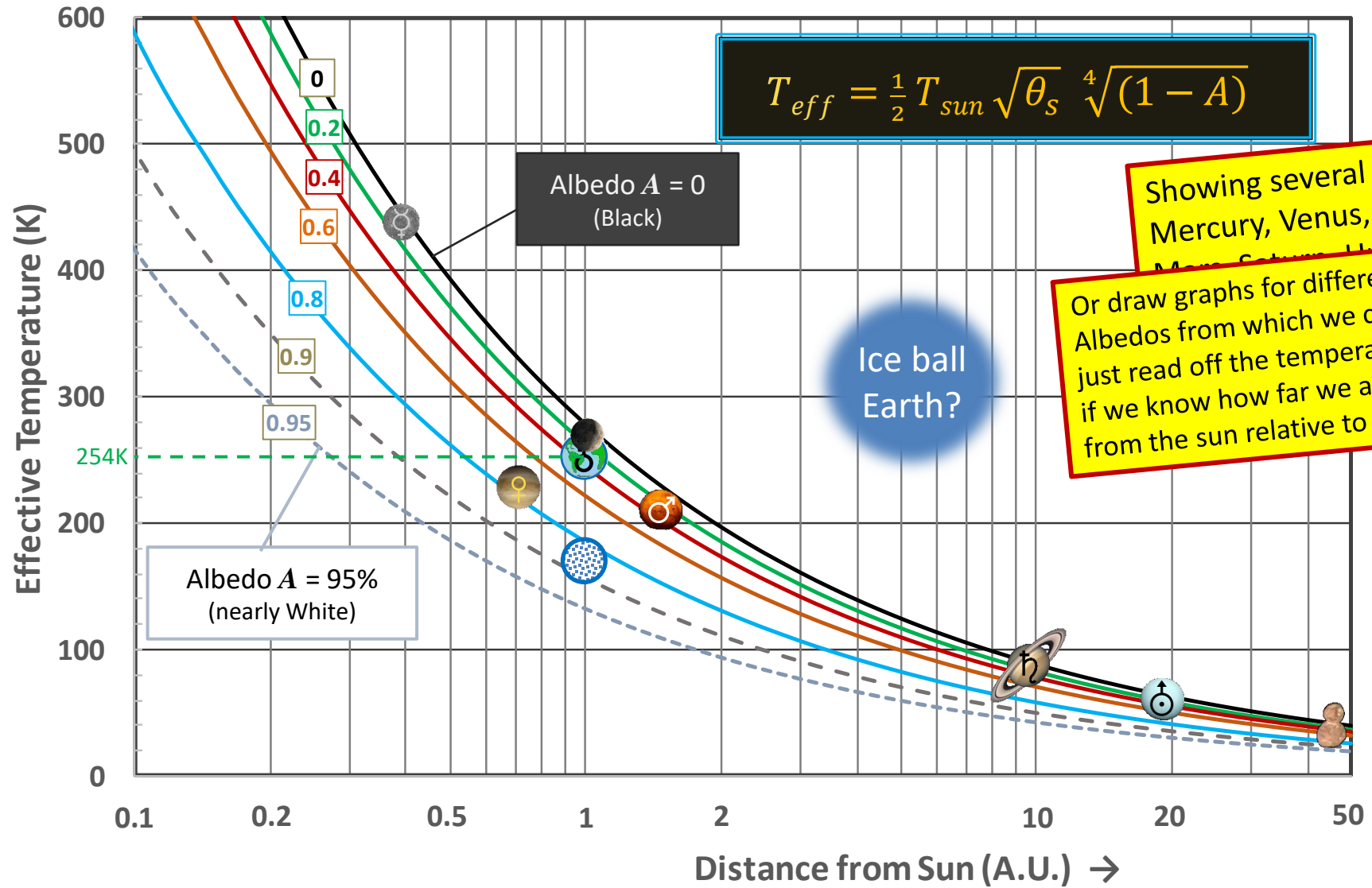
Effective Planetary Temperature vs Distance from Sun



Or draw graphs for different Albedos from which we can just read off the temperature, if we know how far we are from the sun relative to Earth.

— A=0 — A=0.2 — A=0.4 — A=0.6 — A=0.8 - - - A=0.9 - - - - A=0.95

Effective Planetary Temperature vs Distance from Sun



A Tale of Two Planets



77%

-51 °F

So now we understand the Skin Temperatures of Venus and Earth, based on their Albedos and distance from the sun.

Albedo

Effective
(*Apparent Ave.*)
Temperature

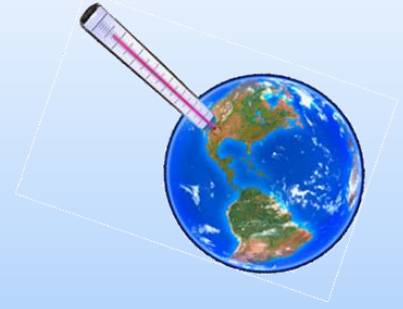


30%

-2 °F

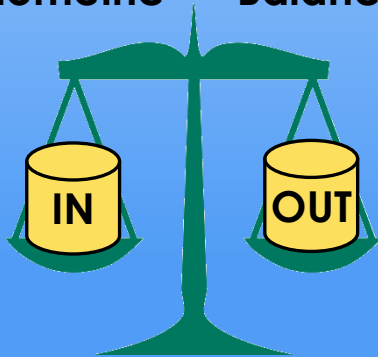


Questions about Earth's basic Effective Temperature?



Why is it about $-2\text{ }^{\circ}\text{F}$?

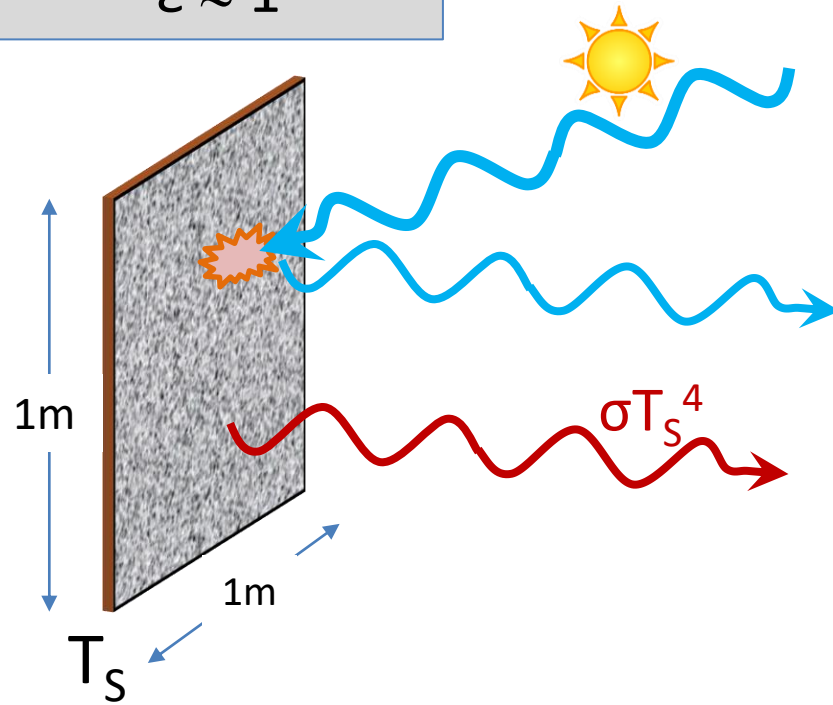
Radiometric Balance





Radiative Greenhouse Effect

Albedo $A = 0.5$
 $\epsilon \approx 1$



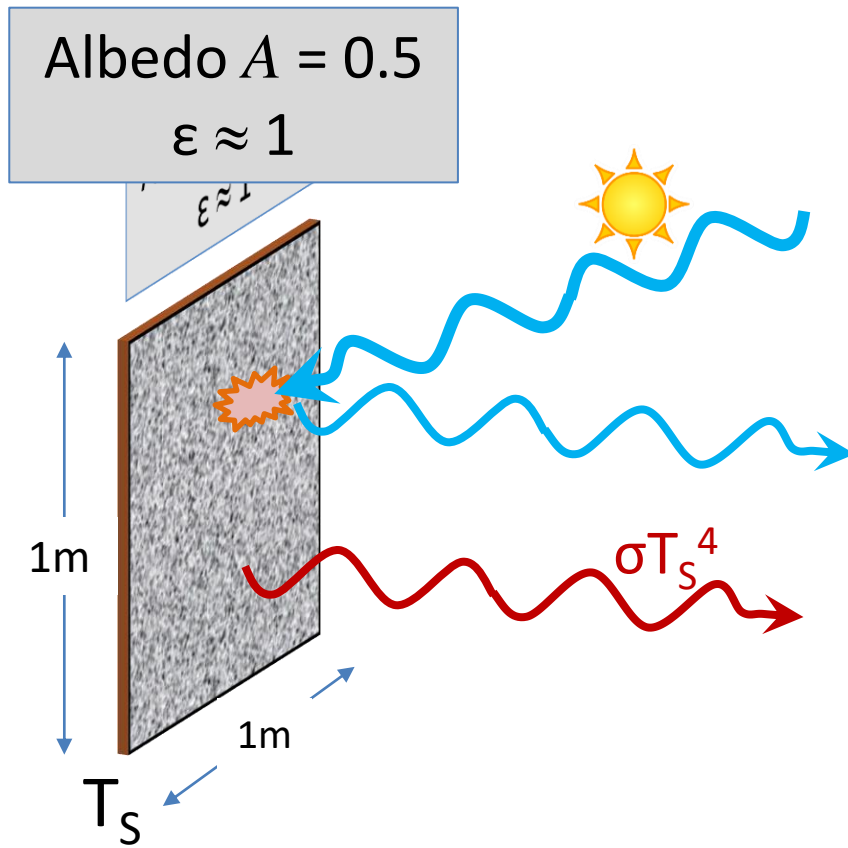
VACUUM

$$\sigma = 56.7 \text{ nW} / \text{m}^2 / \text{K}^4$$

To see how the Greenhouse effect works, imagine we have square meter of surface with Albedo 50%. 1000 W/m² of sunlight is coming in, 50% of which is absorbed as heat. Depending on its temperature, the surface also emits some amount of Blackbody IR radiation.

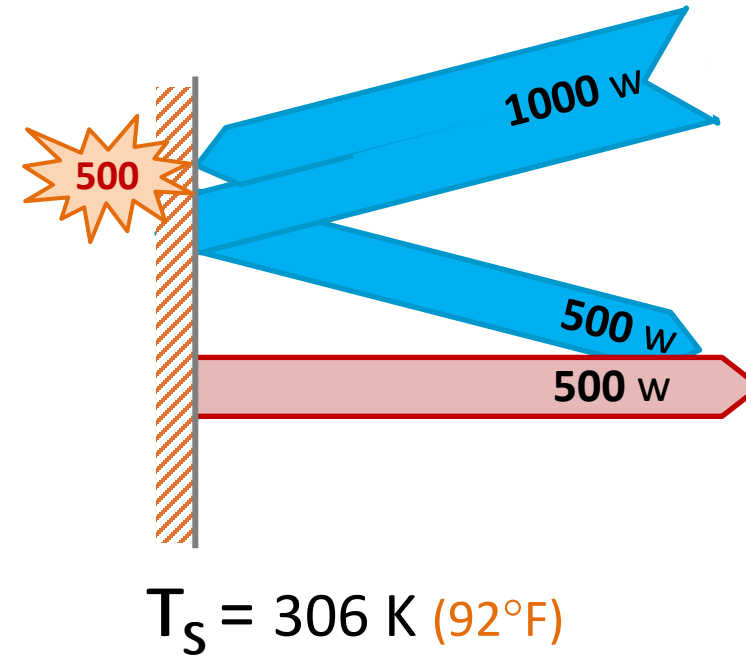


Radiative Greenhouse Effect



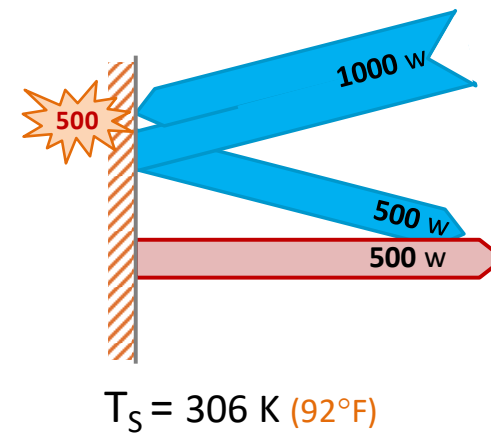
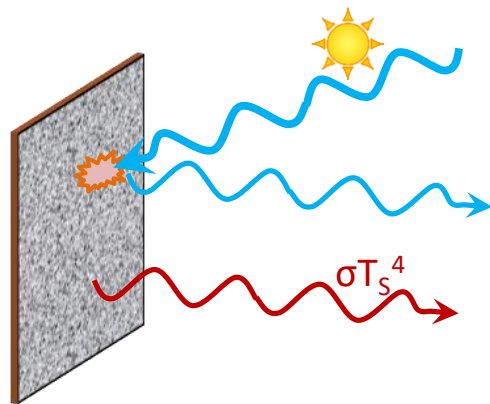
VACUUM

$$\sigma = 56.7 \text{ nW} / \text{m}^2 / \text{K}^4$$



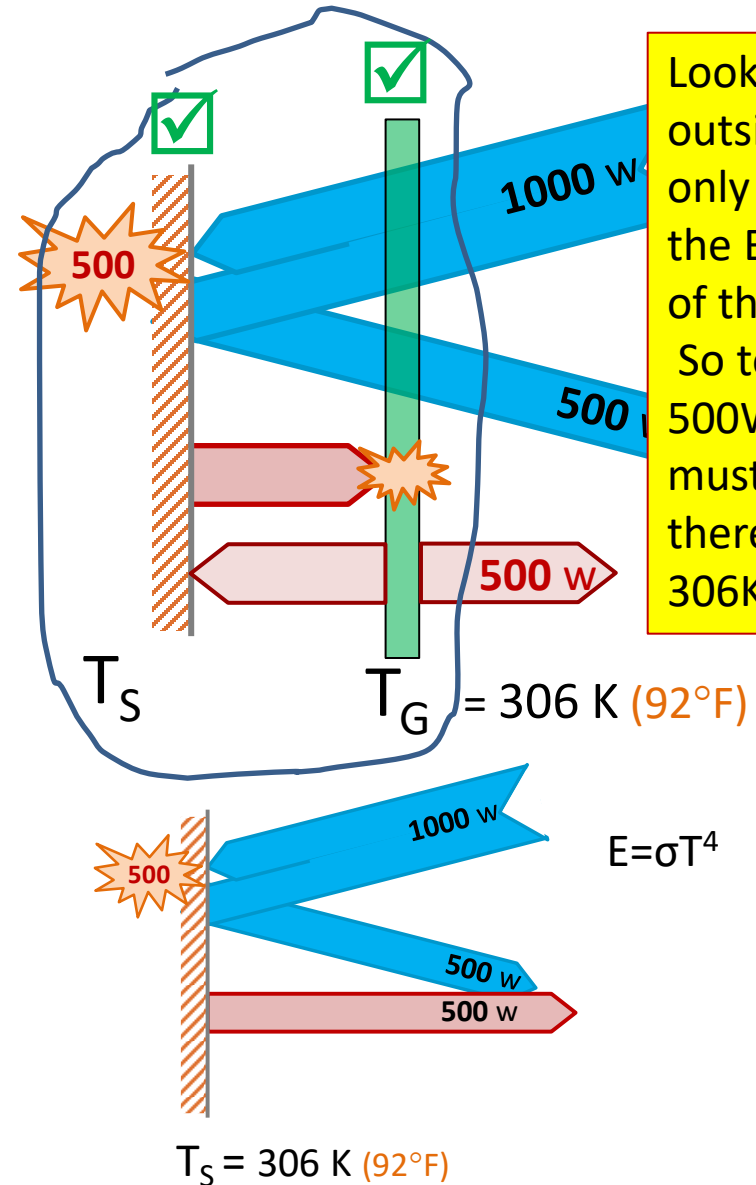
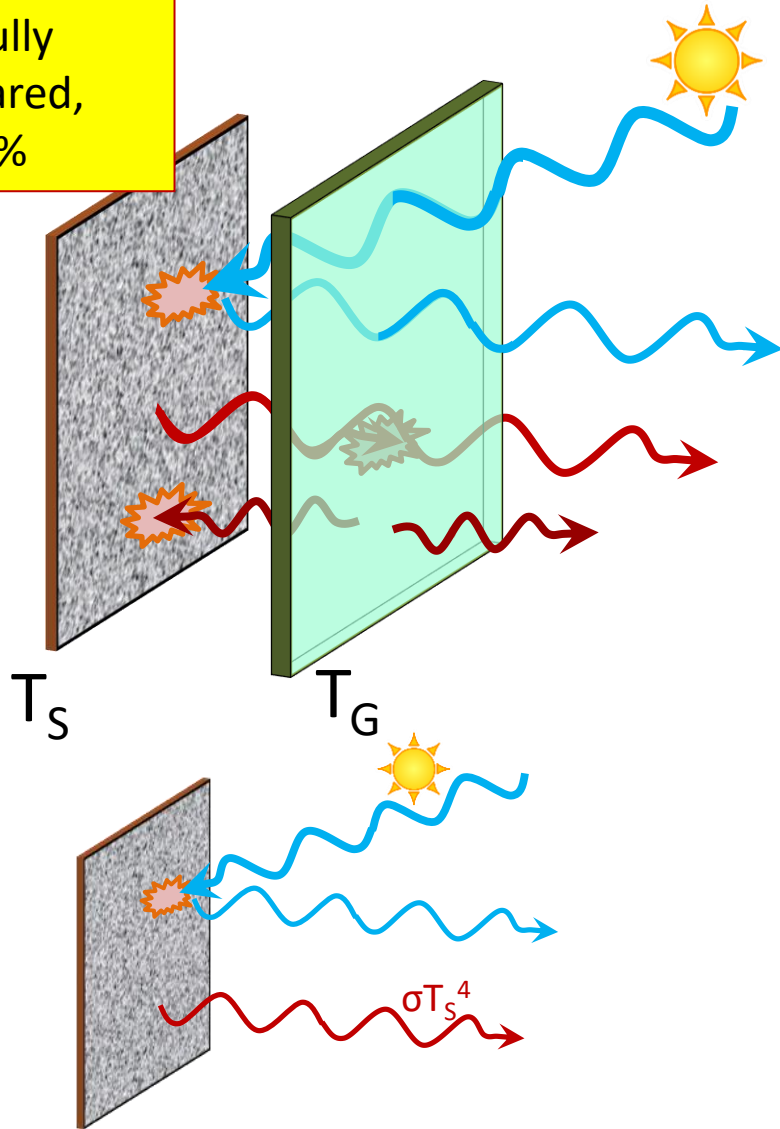
In order to achieve equilibrium, since the surface absorbs 500 watts of sunlight, it must be able to emit 500 watts of IR radiation. Thus its temperature must be $T_s = 306\text{K}$

Radiative Greenhouse Effect



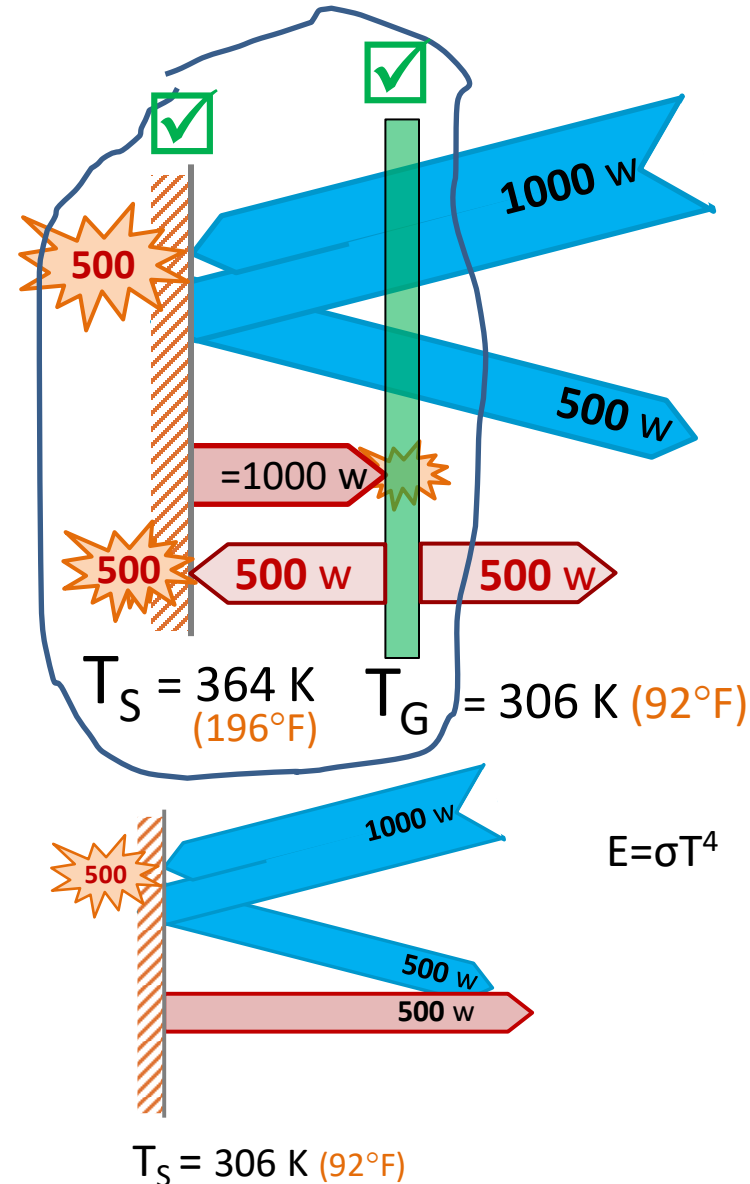
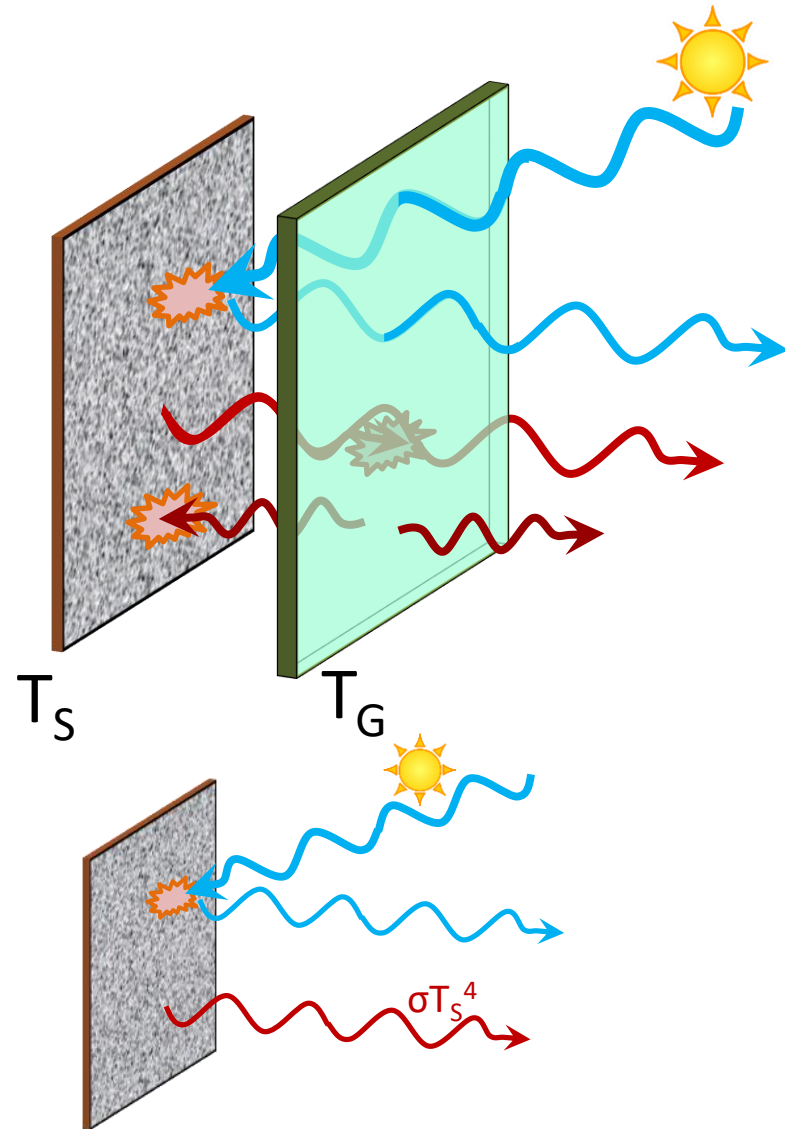
Radiative Greenhouse Effect

Now let's put a sheet of glass in front of the surface. Glass fully absorbs LW Infrared, emissivity = 100%



Looking from the outside (space), the only IR is coming from the BlackBody emission of the glass sheet. So to balance the 500W of sunlight, it must emit 500W and therefore has the same 306K temperature!

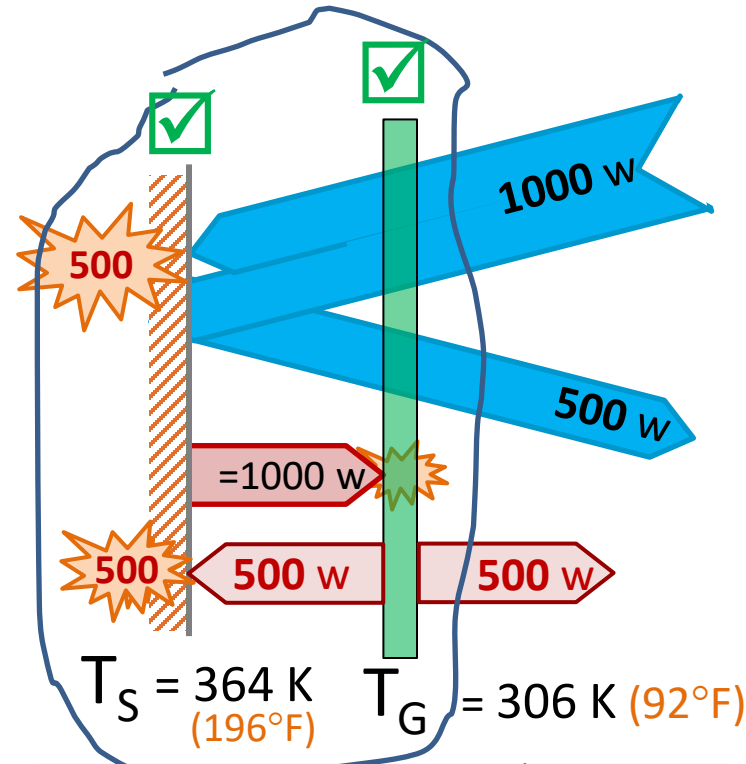
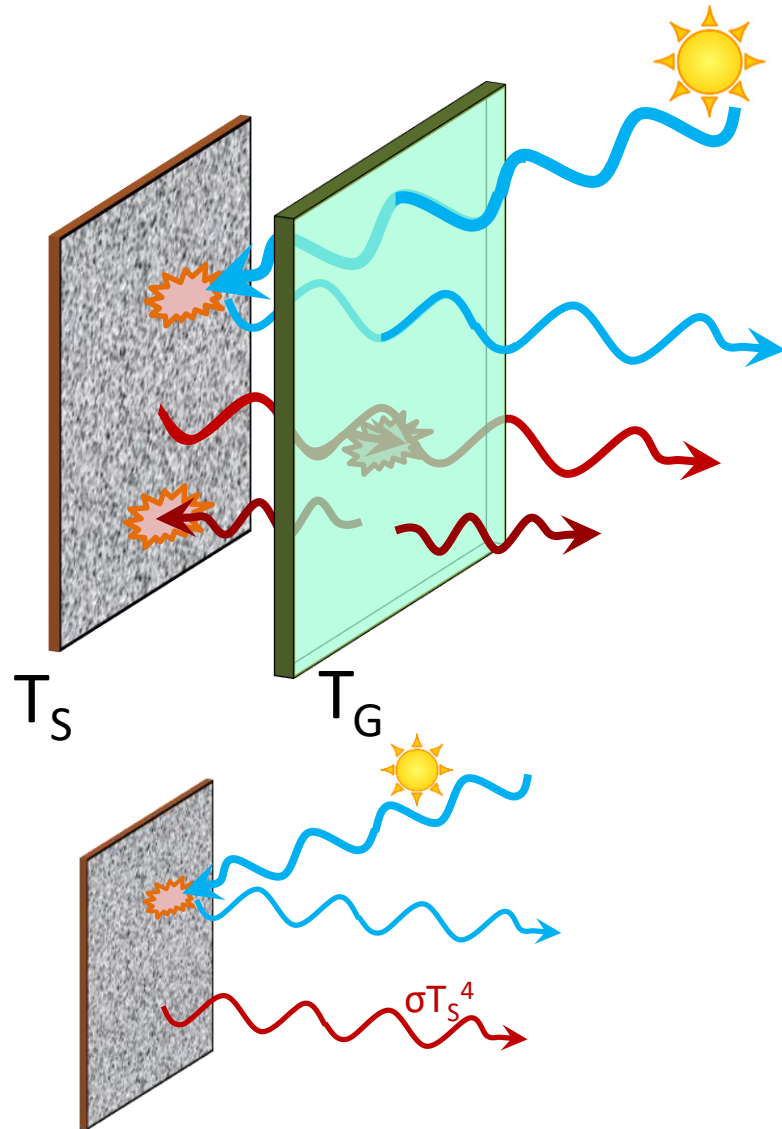
Radiative Greenhouse Effect



The glass must also emit 500W back towards the surface, where it is entirely absorbed. Thus surface is now getting 1000 watts total heat. It must then warm up so it emits 1000W, which means the temperature is 364K.



Radiative Greenhouse Effect



It turns out that the Surface is warmed up by 19%, and this is true regardless of its Albedo!

19%

Power Doubled

$E = \sigma T^4$

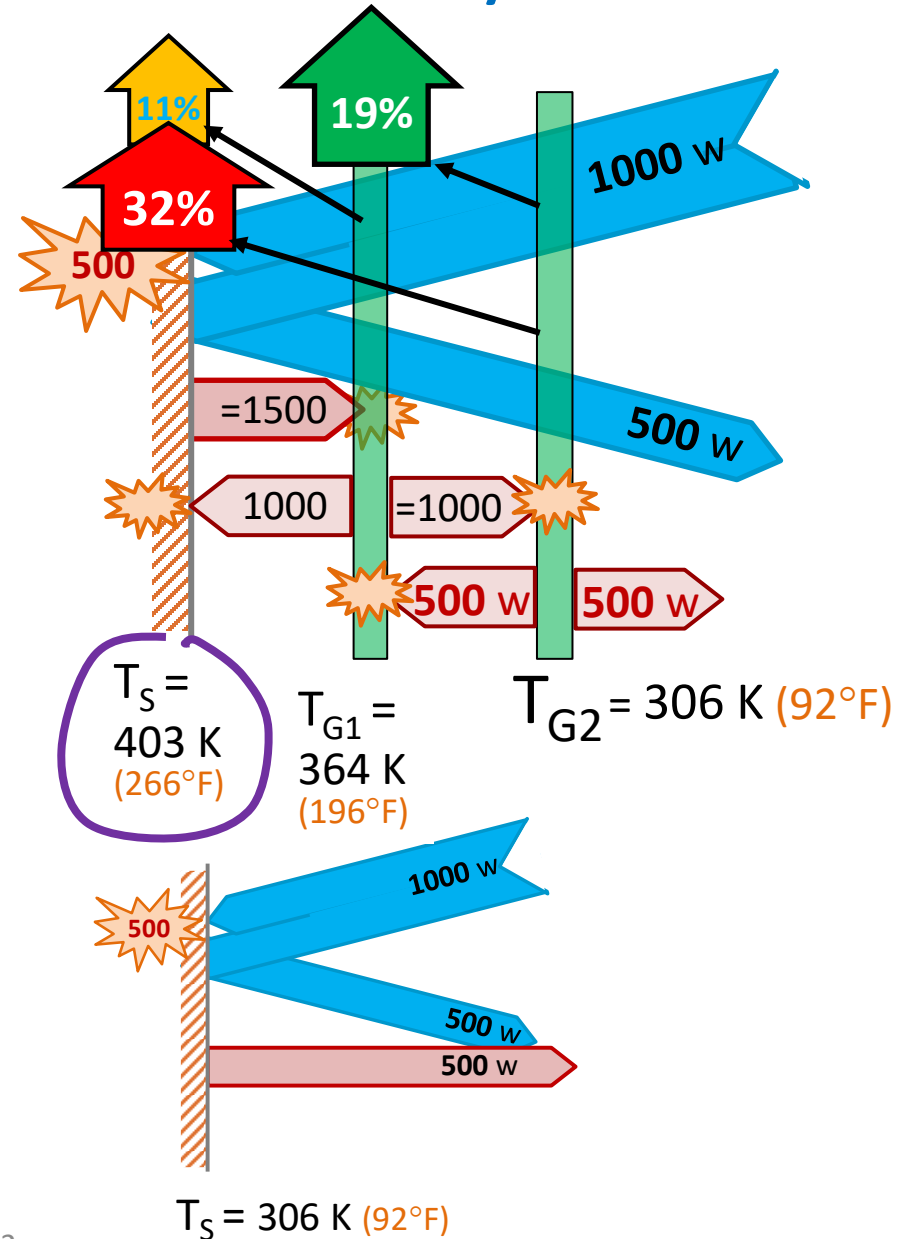
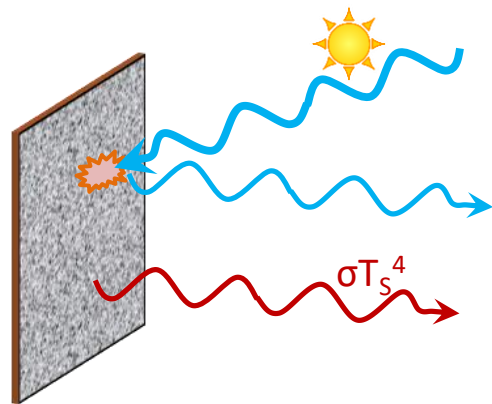
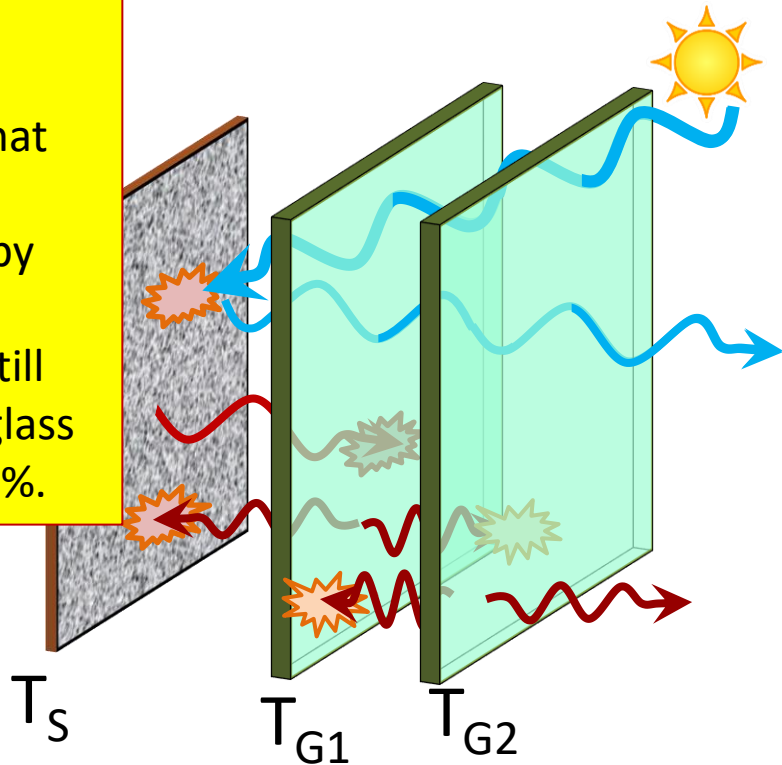
... so T^4 must **double**

... so T increases by $\sqrt[4]{2} \approx 1.19$

Regardless of A !

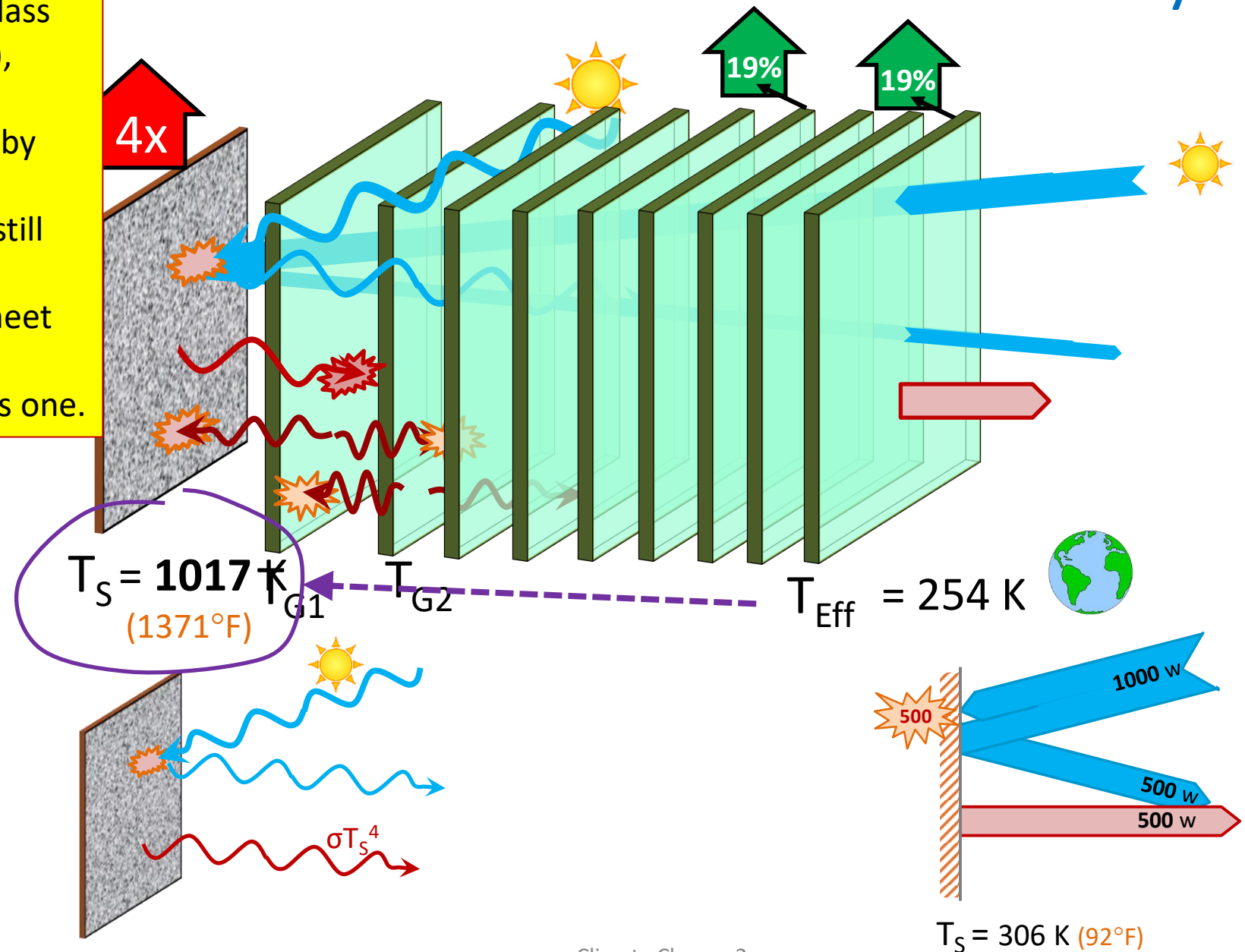
Radiative Greenhouse Effect: 2 Layers

What if we use 2 sheets of glass (still in a Vacuum!)
 Then it turns out that the Surface temperature rises by 32%!
 The outer glass is still 306K, and the 2nd glass sheet warms up 19%.

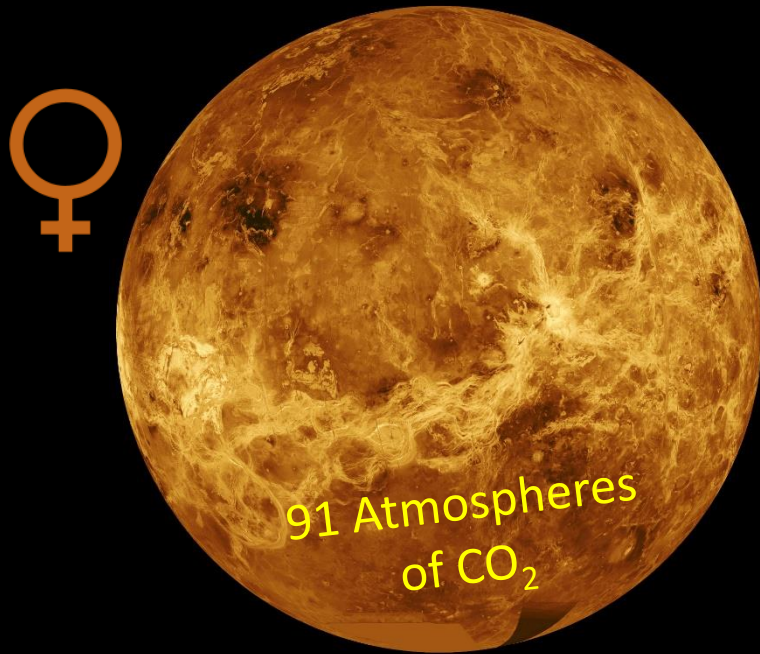


Radiative Greenhouse Effect: n = 9 Layers

With 9 sheets of glass (still in a Vacuum!), the Surface temperature rises by 400%! Super hot. The outer glass is still 306K, and each successive glass sheet warms up by 19% above the previous one.



A Tale of Two Planets: Down at the Surface

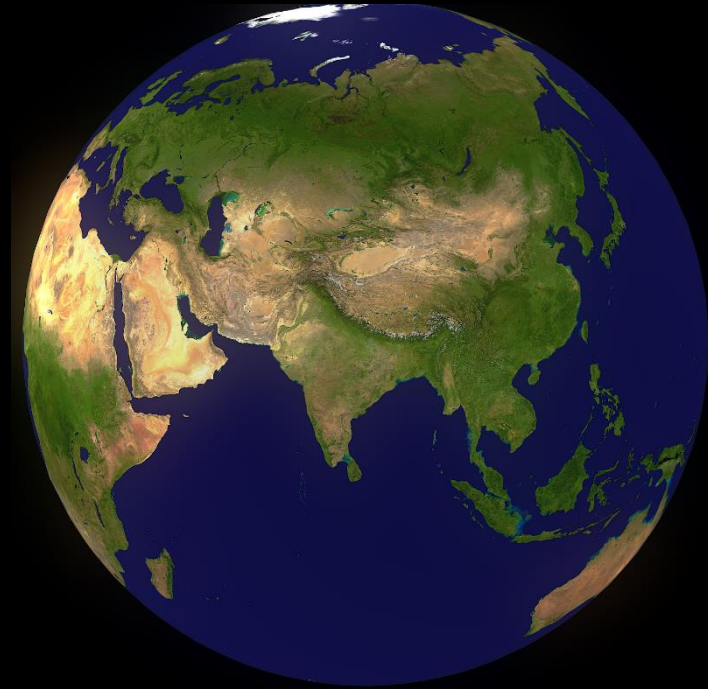


91 Atmospheres
of CO₂

867 °F

Not hard to see
how Venus
Surface gets to
867F via multi-
level Greenhouse
Effect

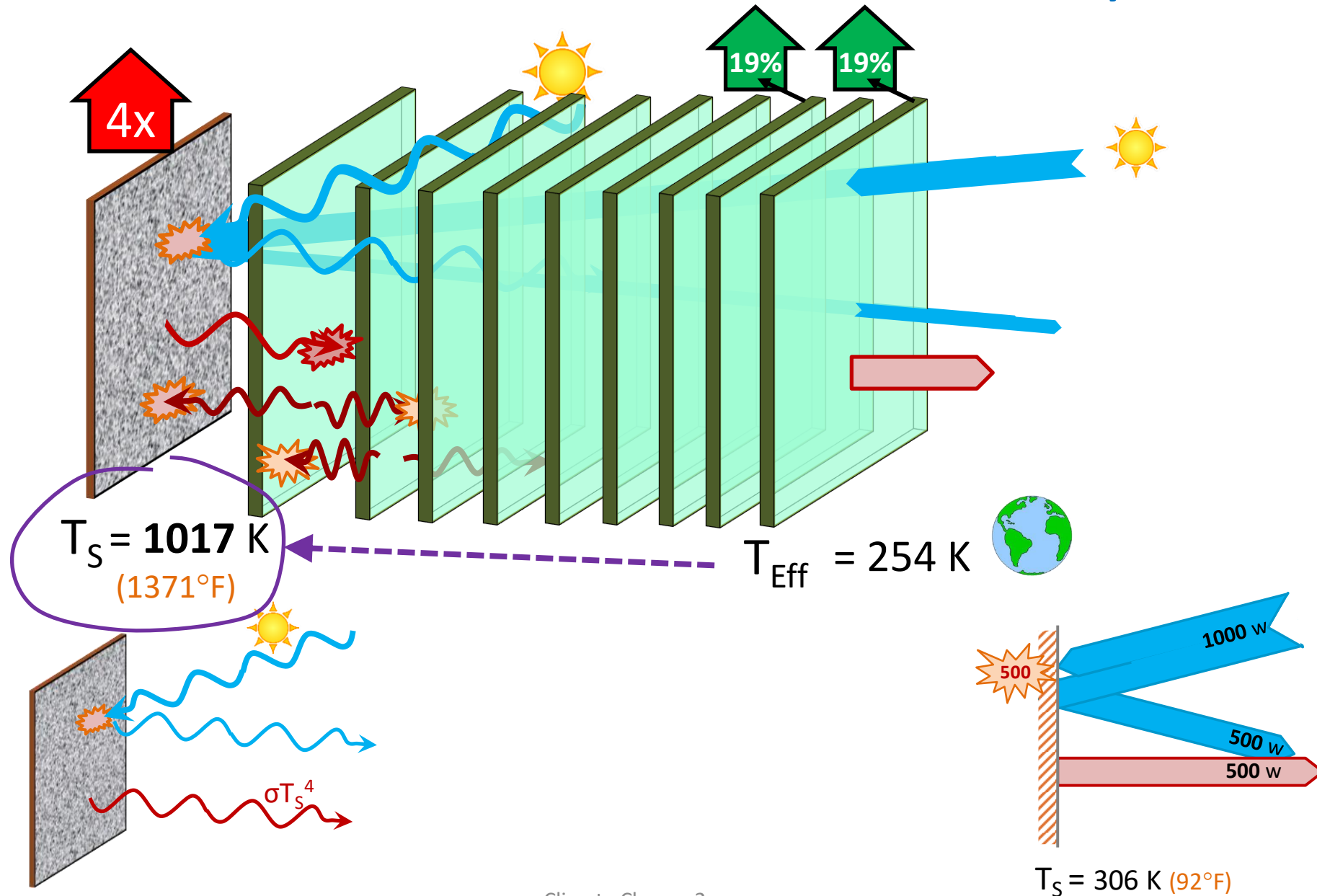
Average
Surface
Temperature



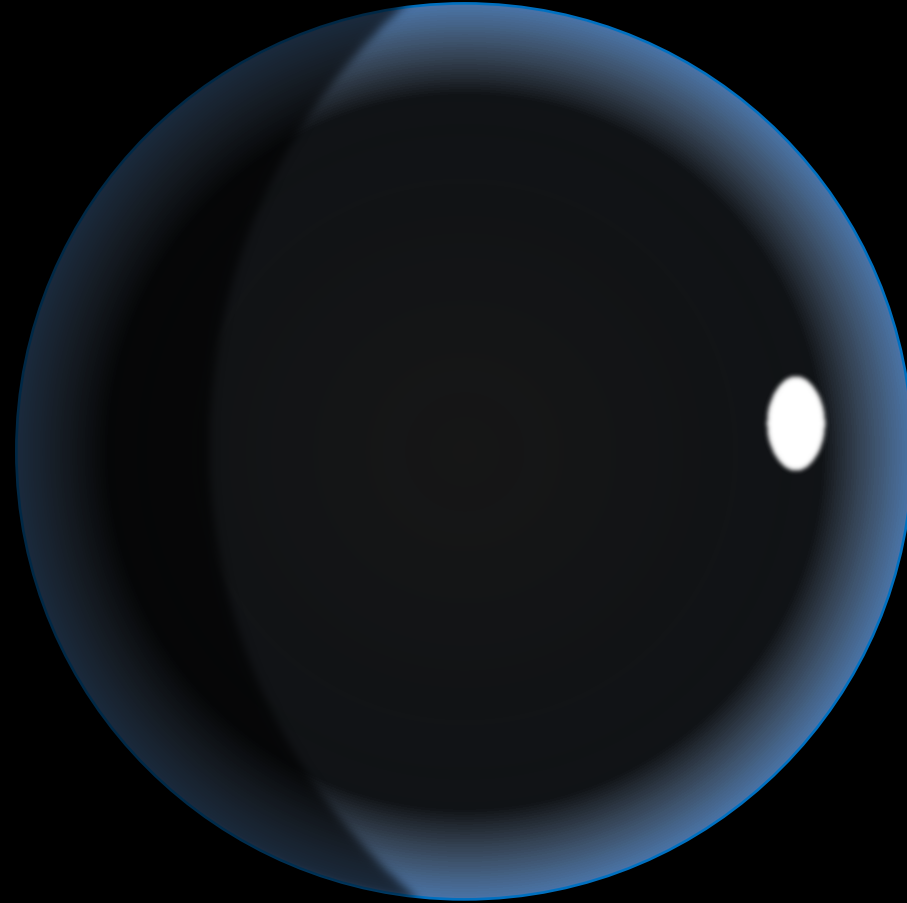
59 °F



Radiative Greenhouse Effect: n = 9 Layers



Crystal Earth?



Crystal Earth?

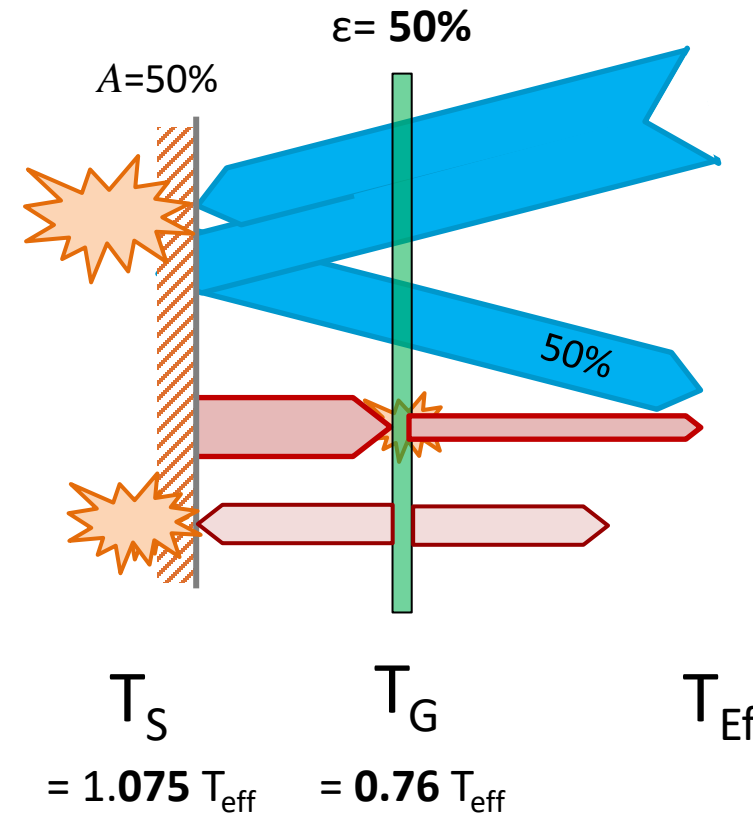
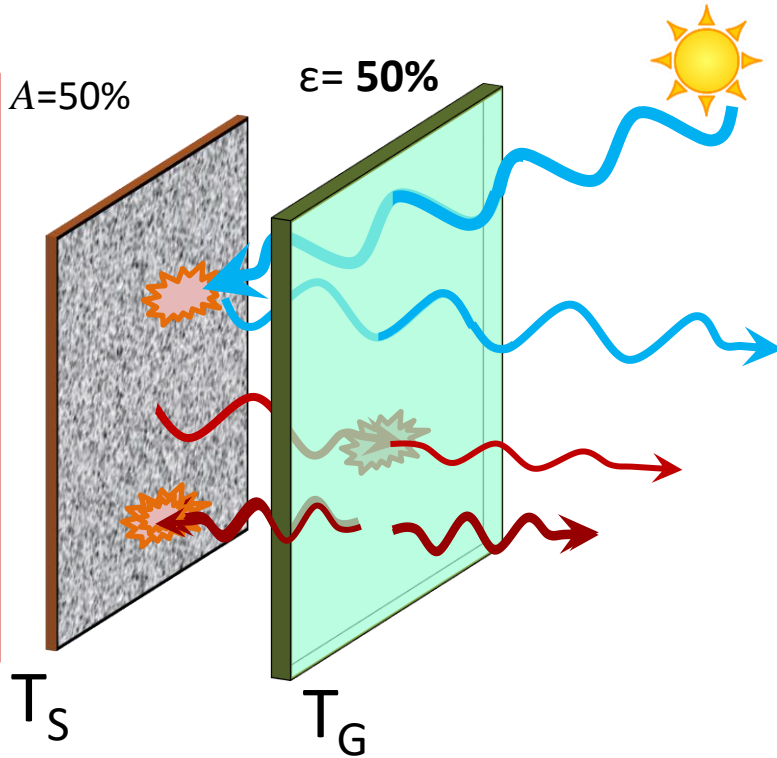


Crystal Earth?



Radiative Greenhouse Effect: Thin 'Glass'

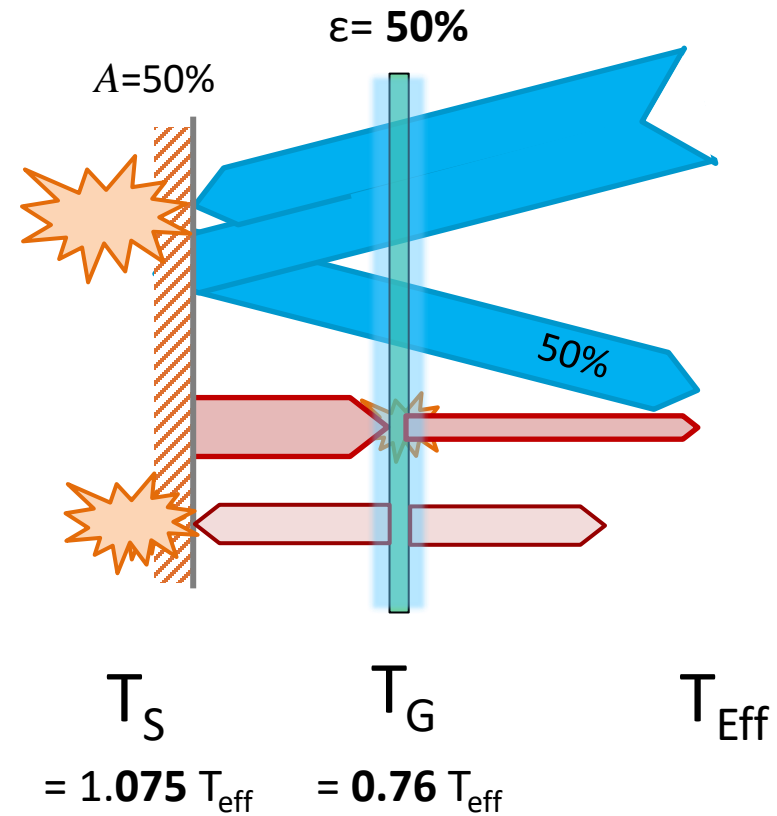
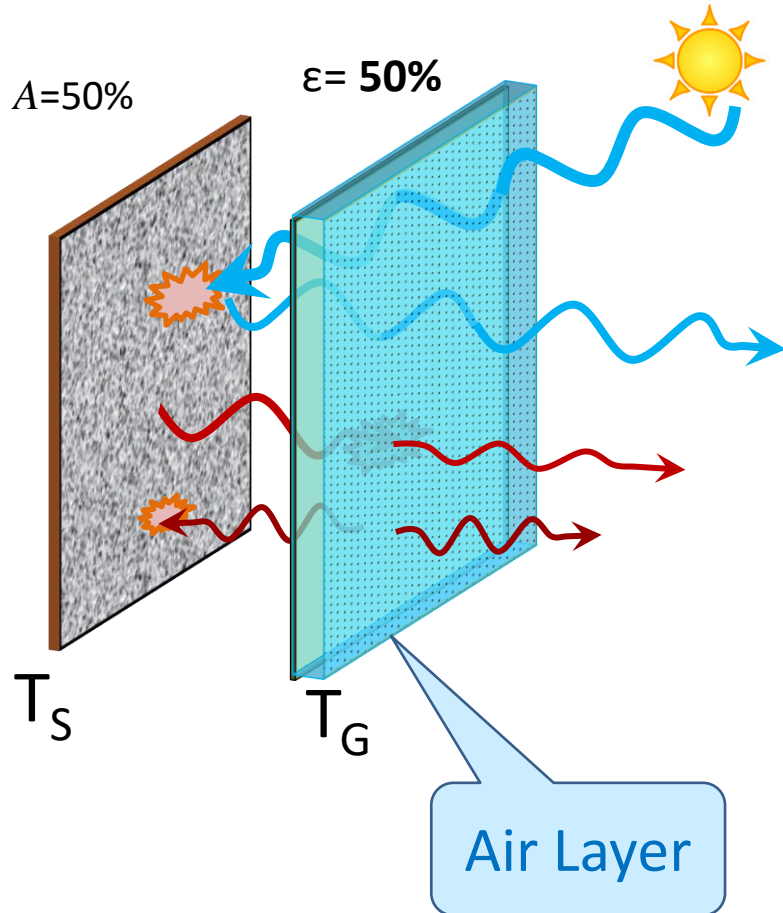
Atmospheric layers are *not* fully IR absorbing like glass! Lets look at "thin glass" which is partially absorbing, say emissivity of 50%. This means half the IR from the surface gets through it.



When we do the radiation balance logic, the needed 500W of IR radiation to space comes partly from the surface and partly from the thin glass. The surface temperature rise is only 7.5%, and the glass is very cold.

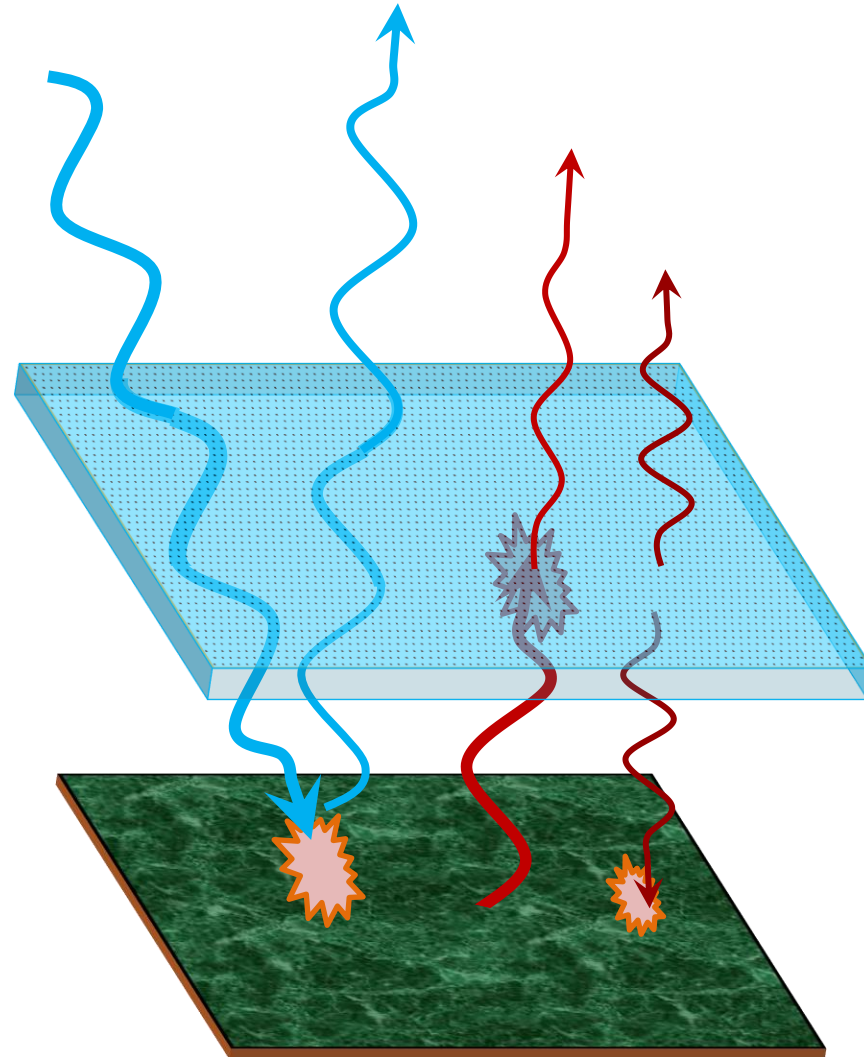
Radiative Greenhouse Effect: Thin Air?

We can imagine that the "thin glass" is actually an air layer containing some IR absorbing gas.....



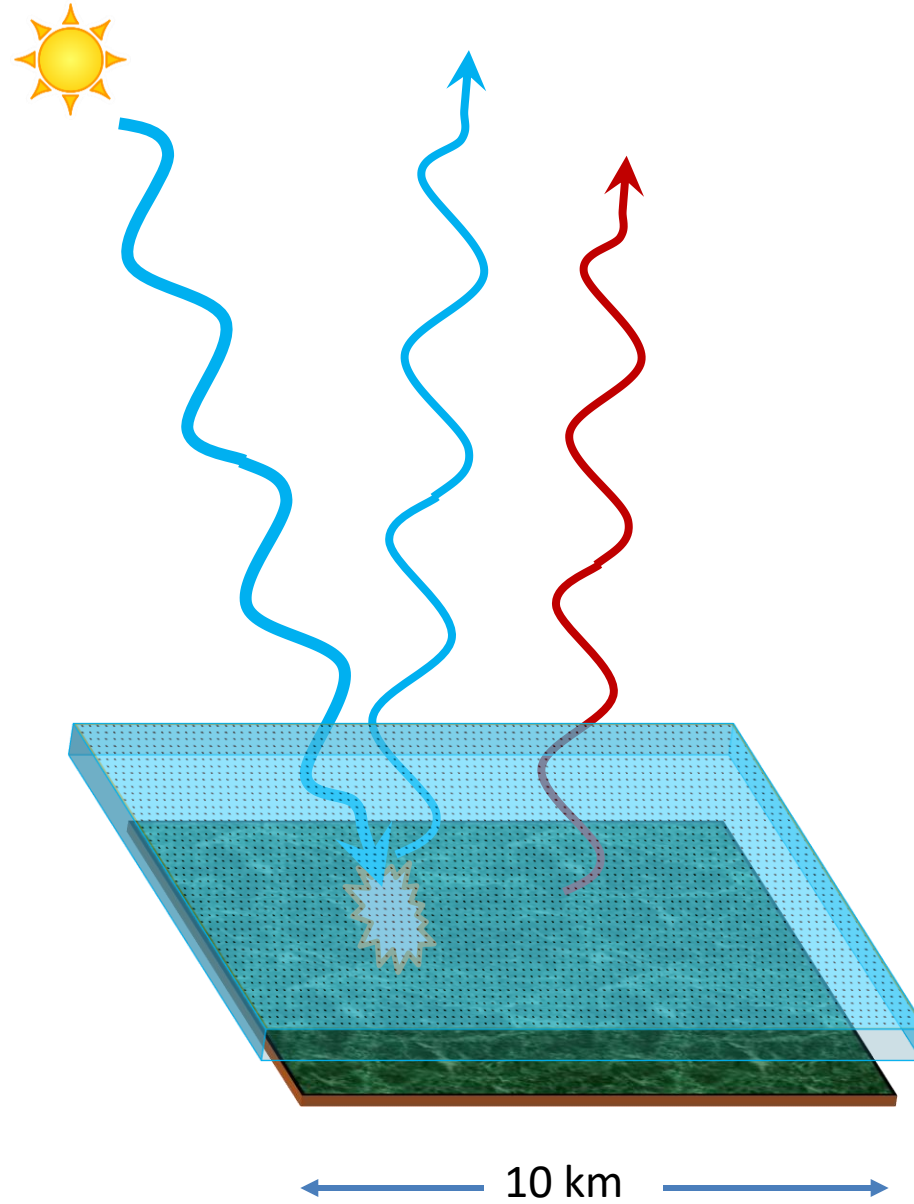
Radiative Greenhouse Effect: Thin Air?

To see how this applies to our earth, imagine turning the surface 90deg and enlarging it to 10km square. Put layers of air above it which are partially transparent to IR. Note that we are no longer in vacuum....



10 km

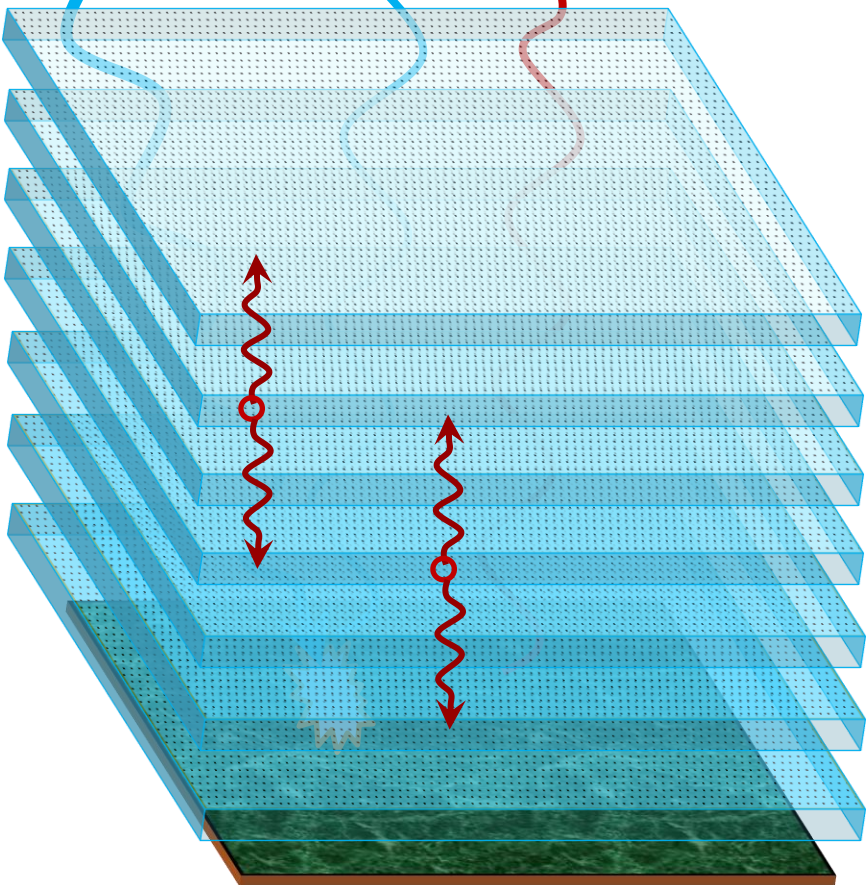
Radiative Greenhouse Effect: Thin Air?



Radiative Greenhouse Effect: Layers of Air



Ave. 254 K
(-2 °F)



Problem:
No Vacuum
between
Layers

Each air layer has a temperature T , an IR emissivity that is less than 100%, and maybe an albedo due to clouds.
The outside skin temperature remains at 254K, but the surface can be much warmer.

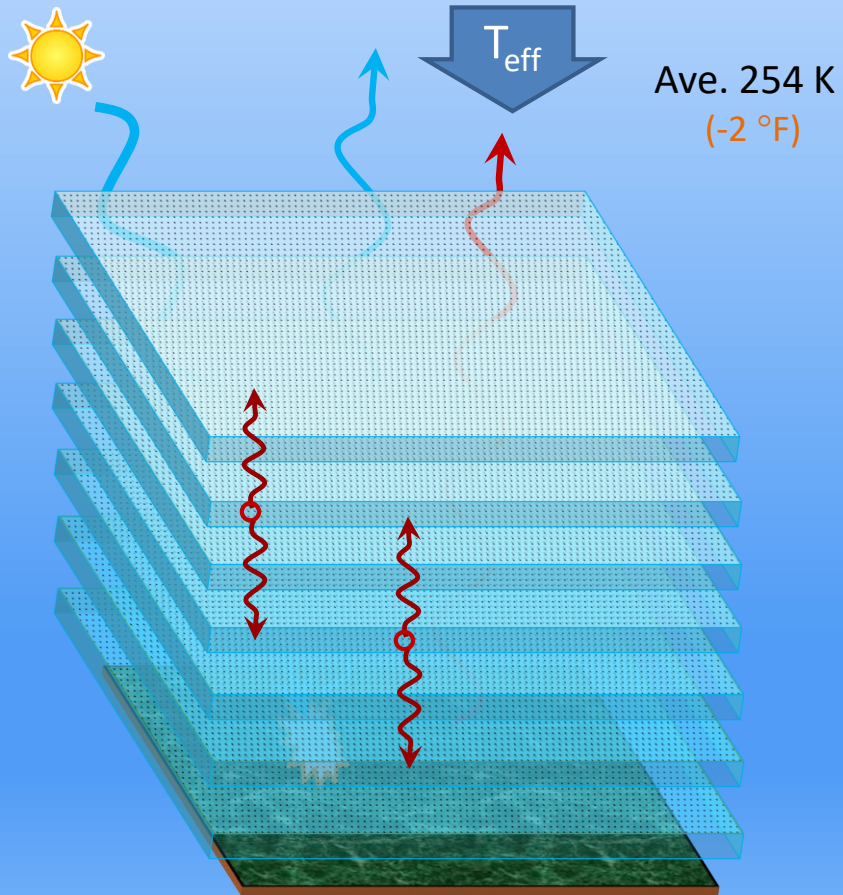
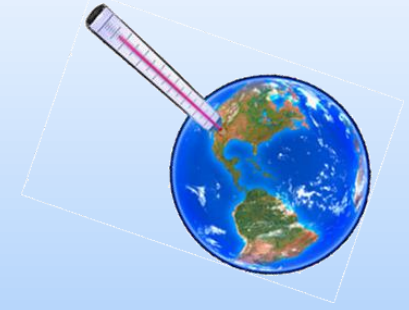
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T, \epsilon < 1, (A)$
- $T_s, \epsilon \approx 1, A_s$



10 km



Questions about Earth's basic Greenhouse Effect?



Modes of Heat Transfer

- Radiative
 - Photons



MAJOR

- Thermal Conduction
 - Molecule to molecule energy transfer



MINOR

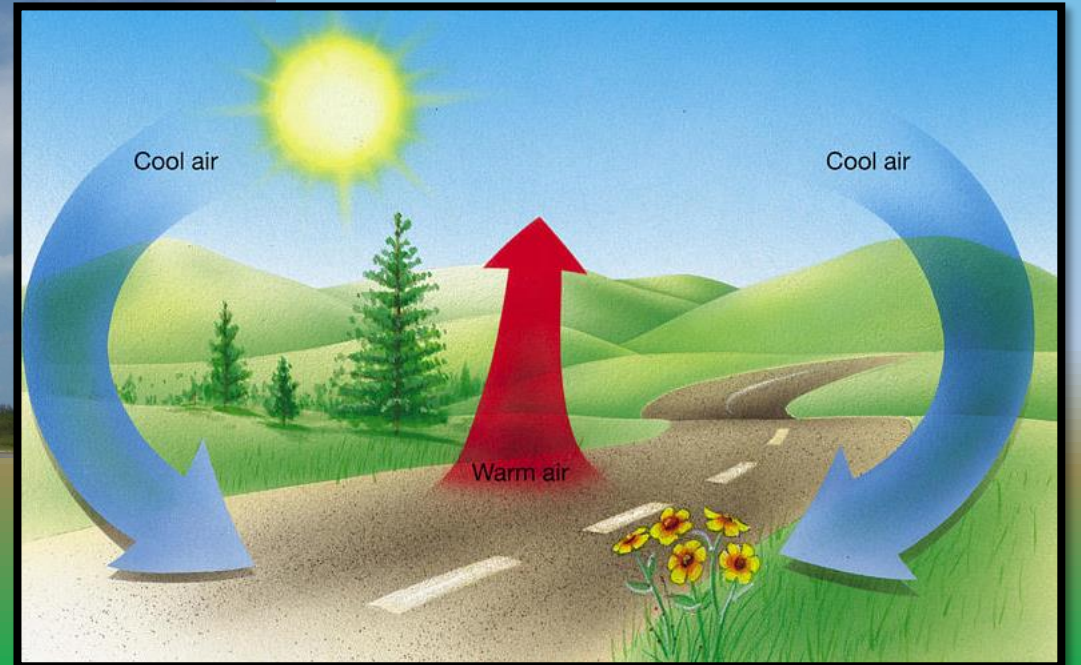
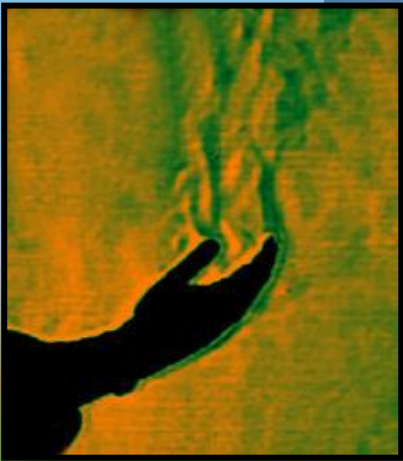
- Thermal Convection
 - Carrying heat via fluid motion
 - Often gravity-driven



MAJOR

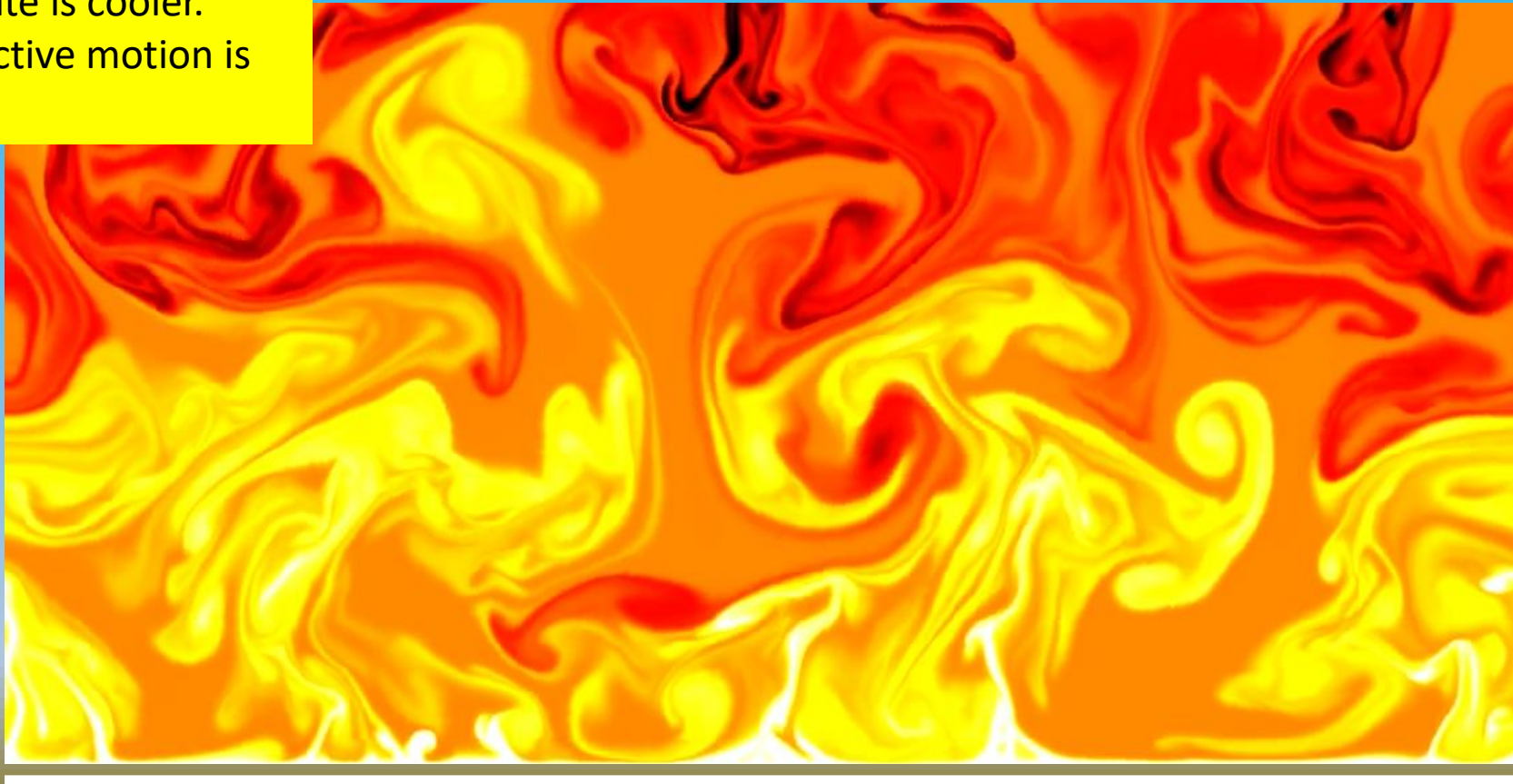


Modes of Heat Transfer: Thermal Convection by Gravity



Thermal Convection

Numerical simulation of thermal gravity convection in an unspecified fluid. The bottom plate is hot, the top plate is cooler. The convective motion is chaotic.



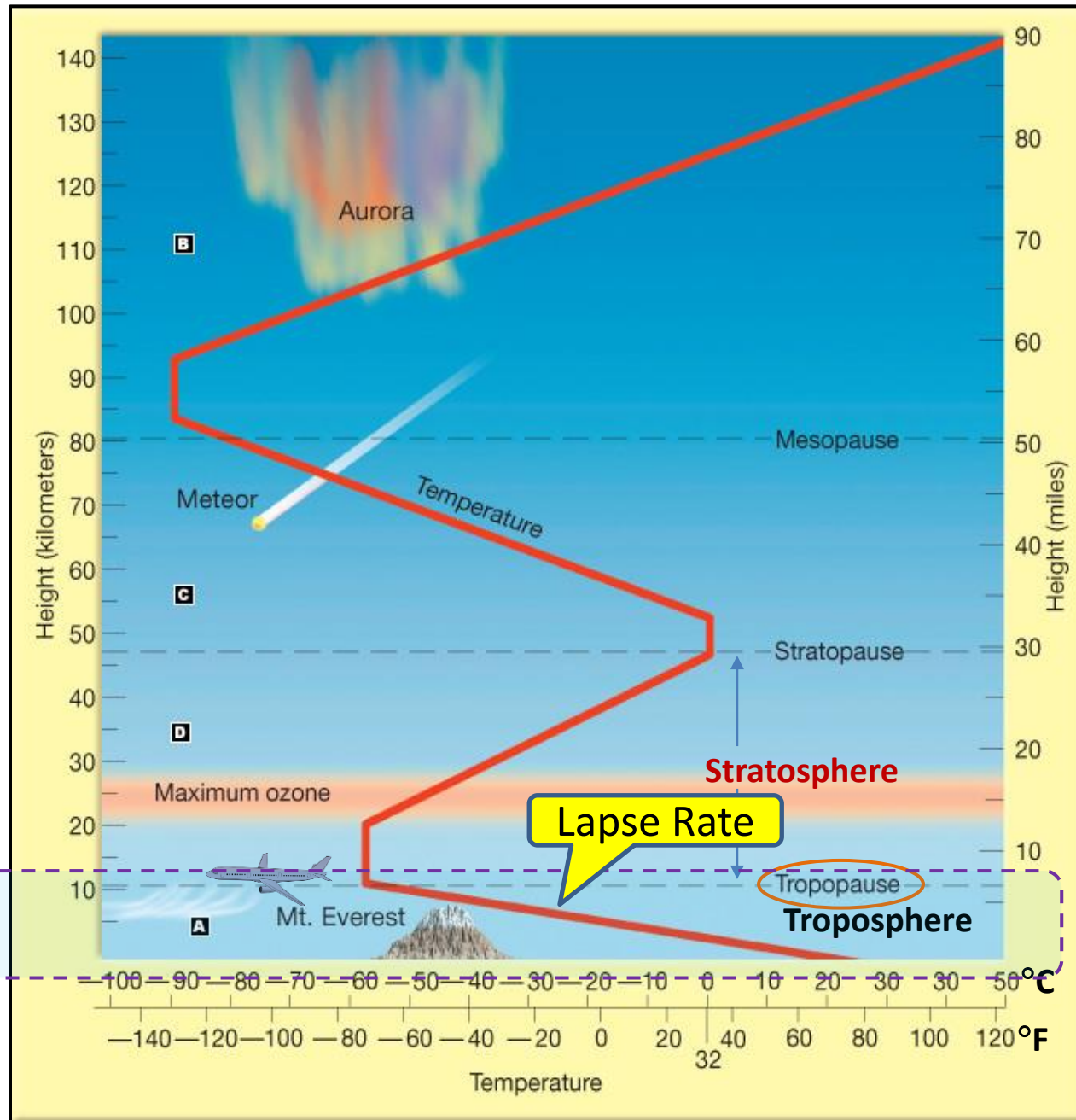
by TurbulenceTeam (2010)

Temperature Profile of the Atmosphere

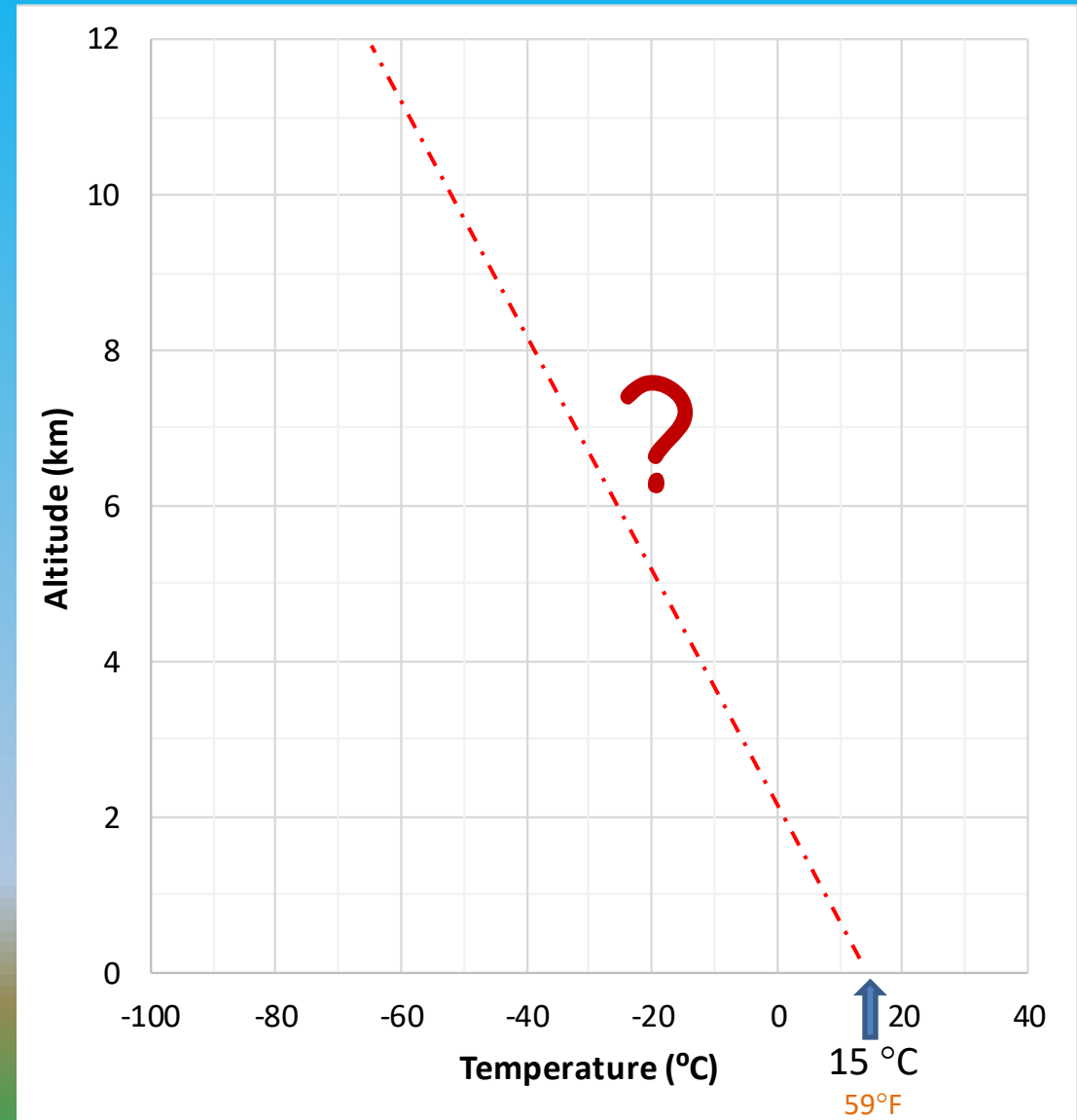
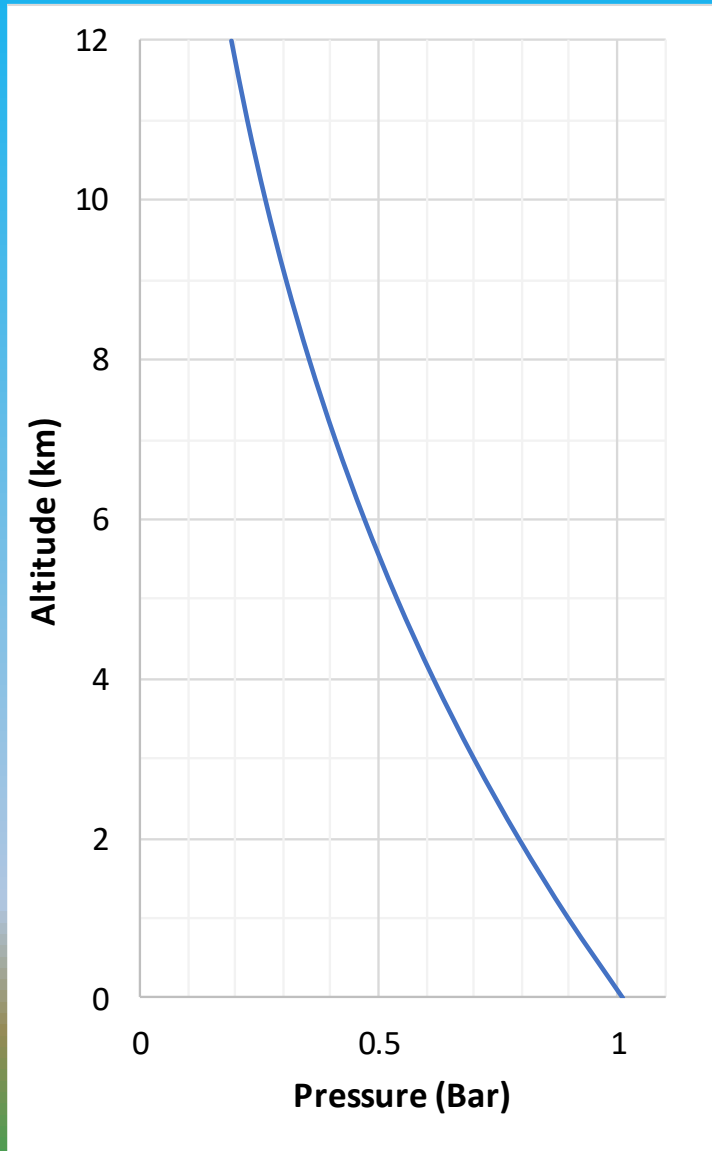
Not much air up here

Solar **UV** Absorption Reigns

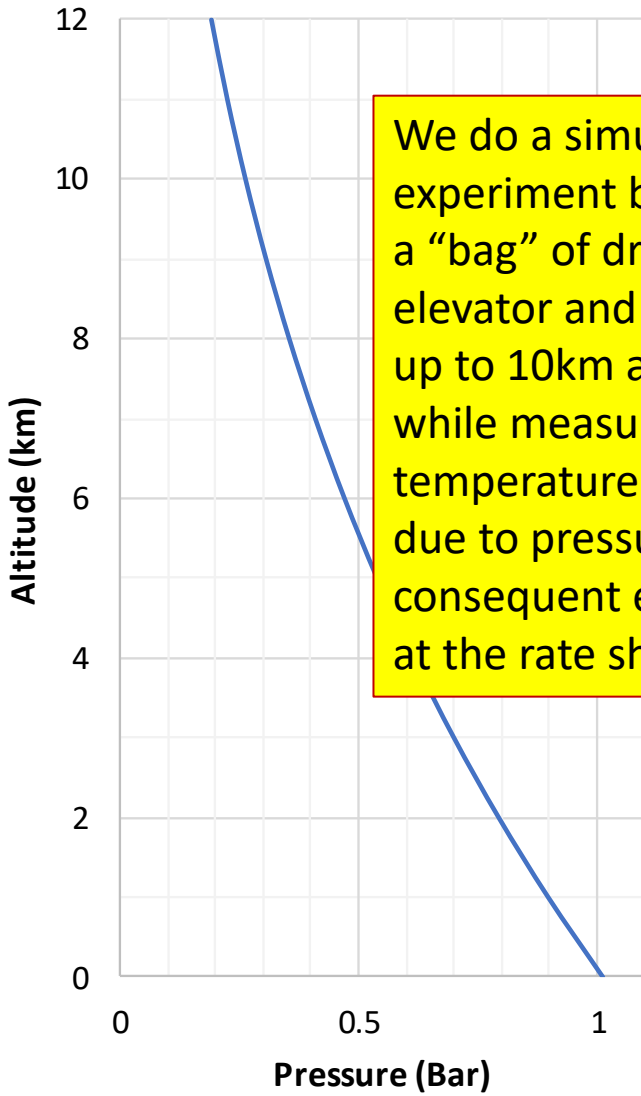
Thermal Convection Reigns



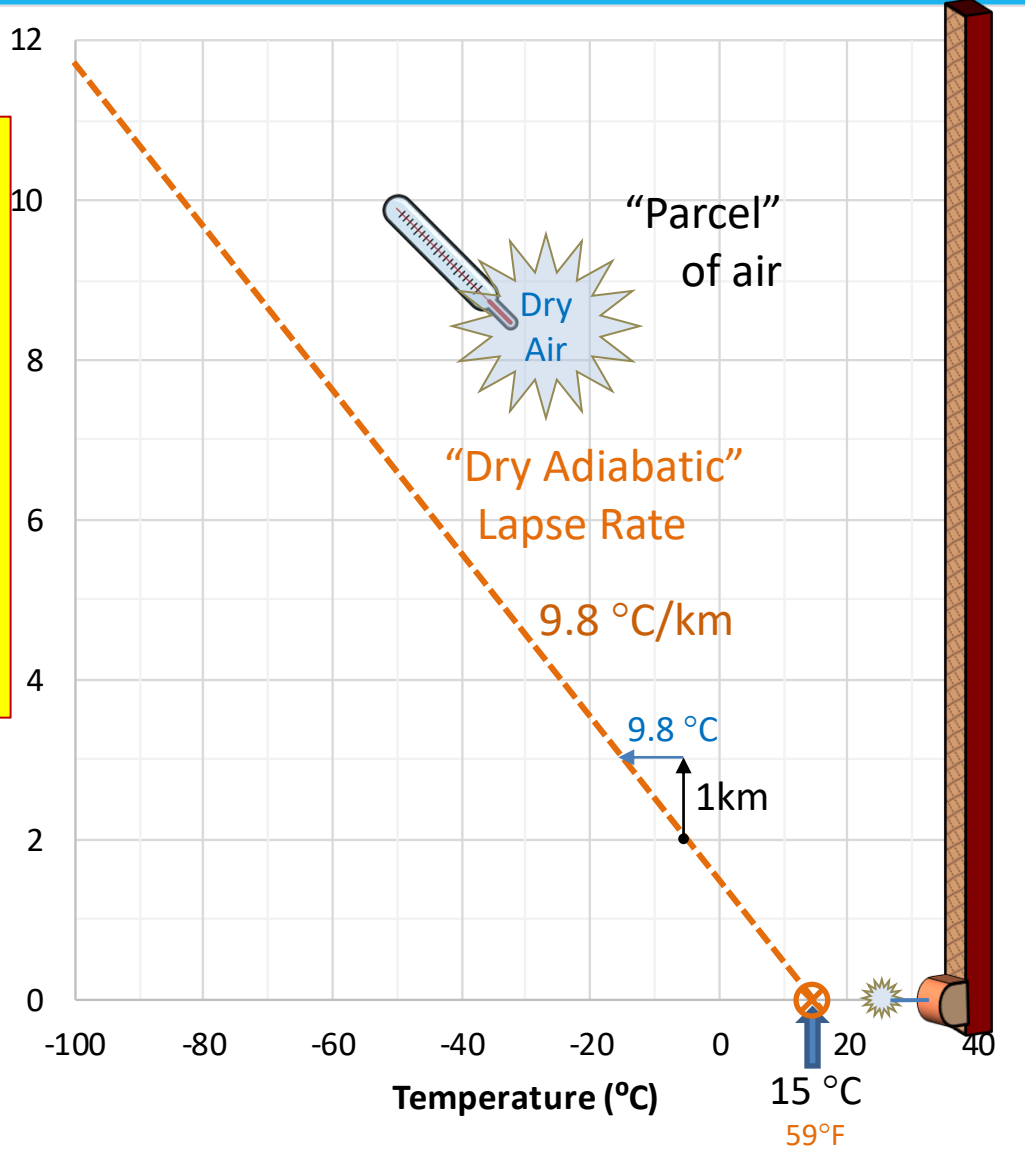
Lapse Rate in Troposphere



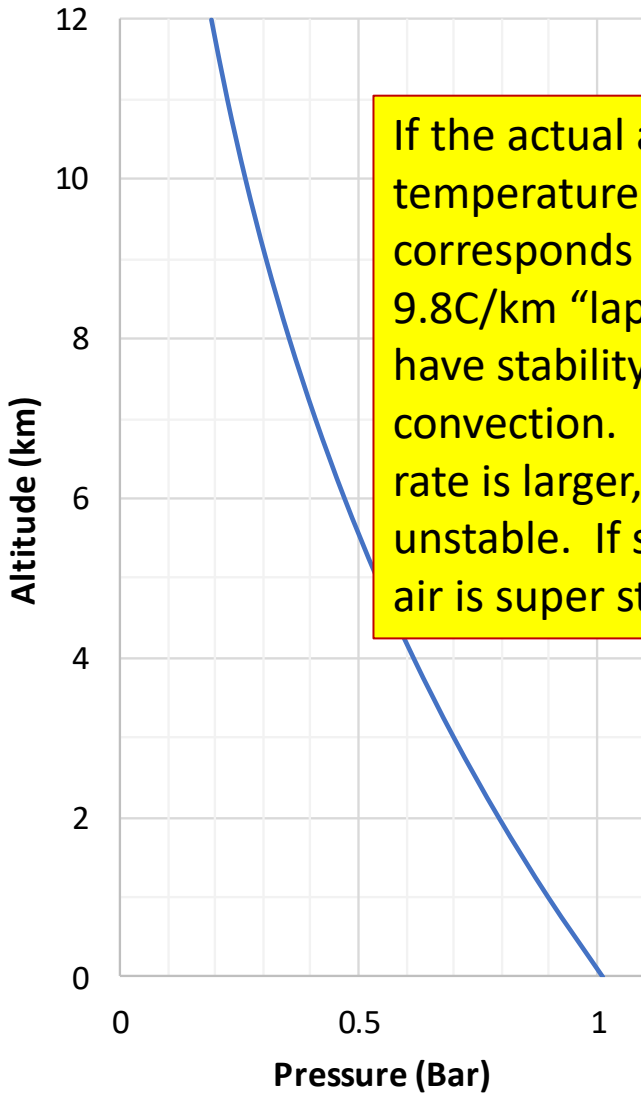
Lapse Rate in Troposphere



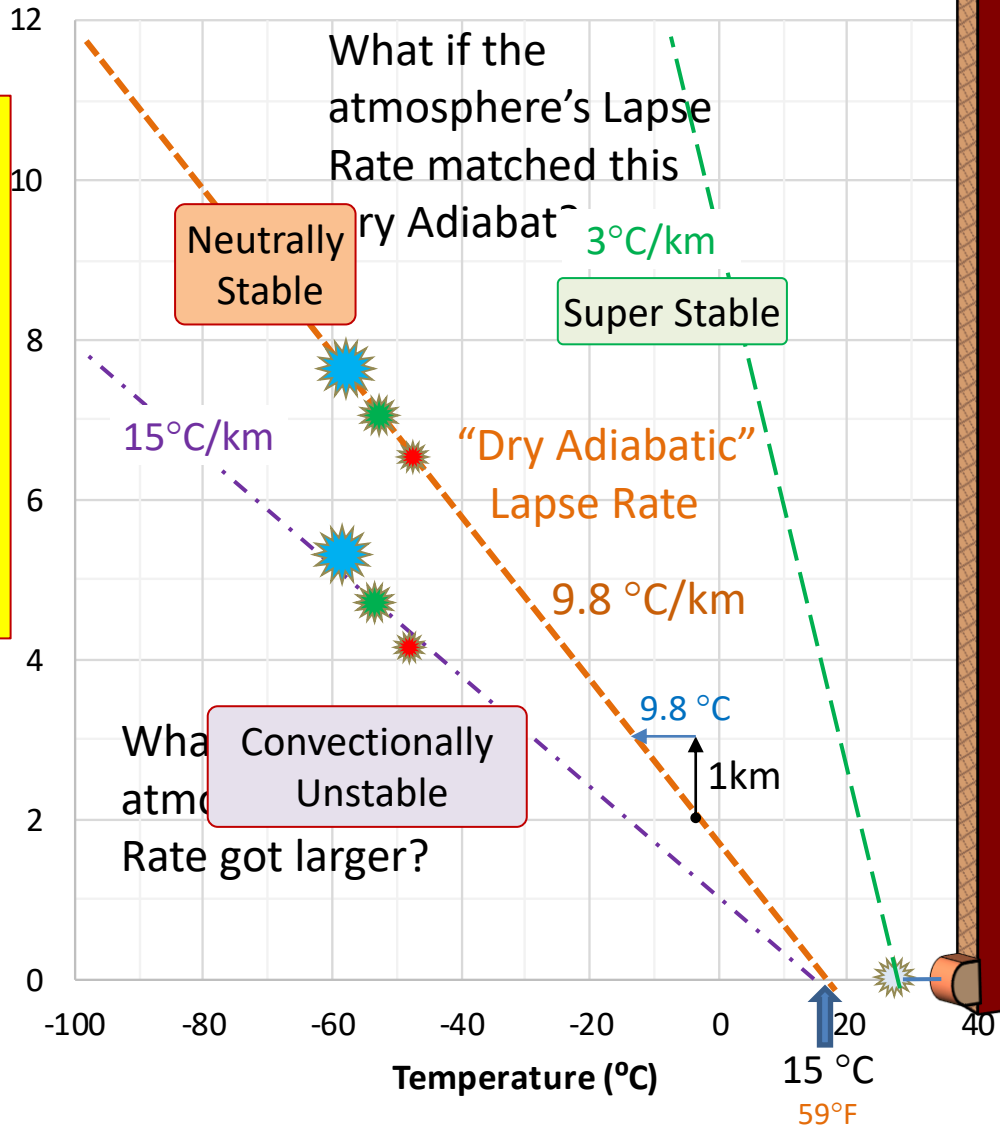
We do a simulated experiment by attaching a “bag” of dry air to an elevator and slowly lift it up to 10km altitude while measuring its temperature. It cools due to pressure drop and consequent expansion, at the rate shown.



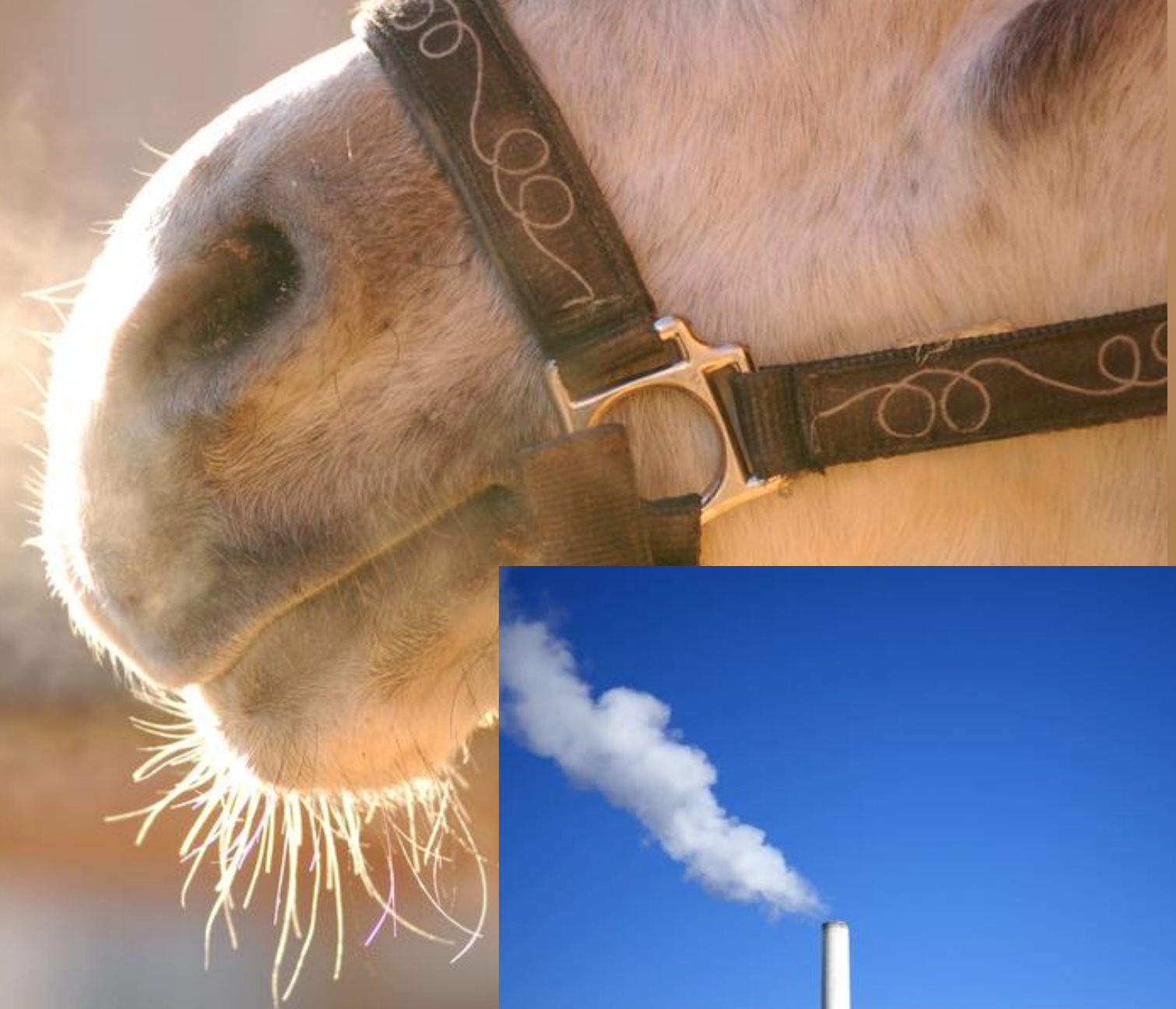
Lapse Rate in Troposphere



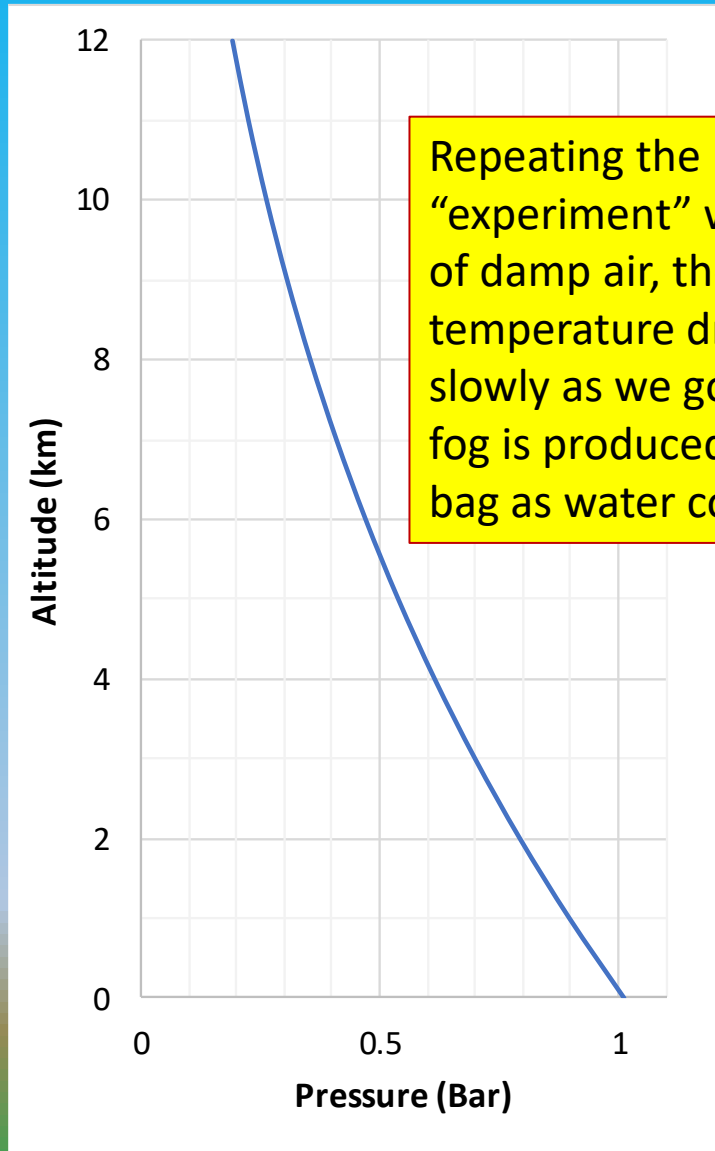
If the actual atmospheric temperature profile corresponds to this 9.8C/km "lapse rate", we have stability and no convection. If the lapse rate is larger, the air is unstable. If smaller, the air is super stable.



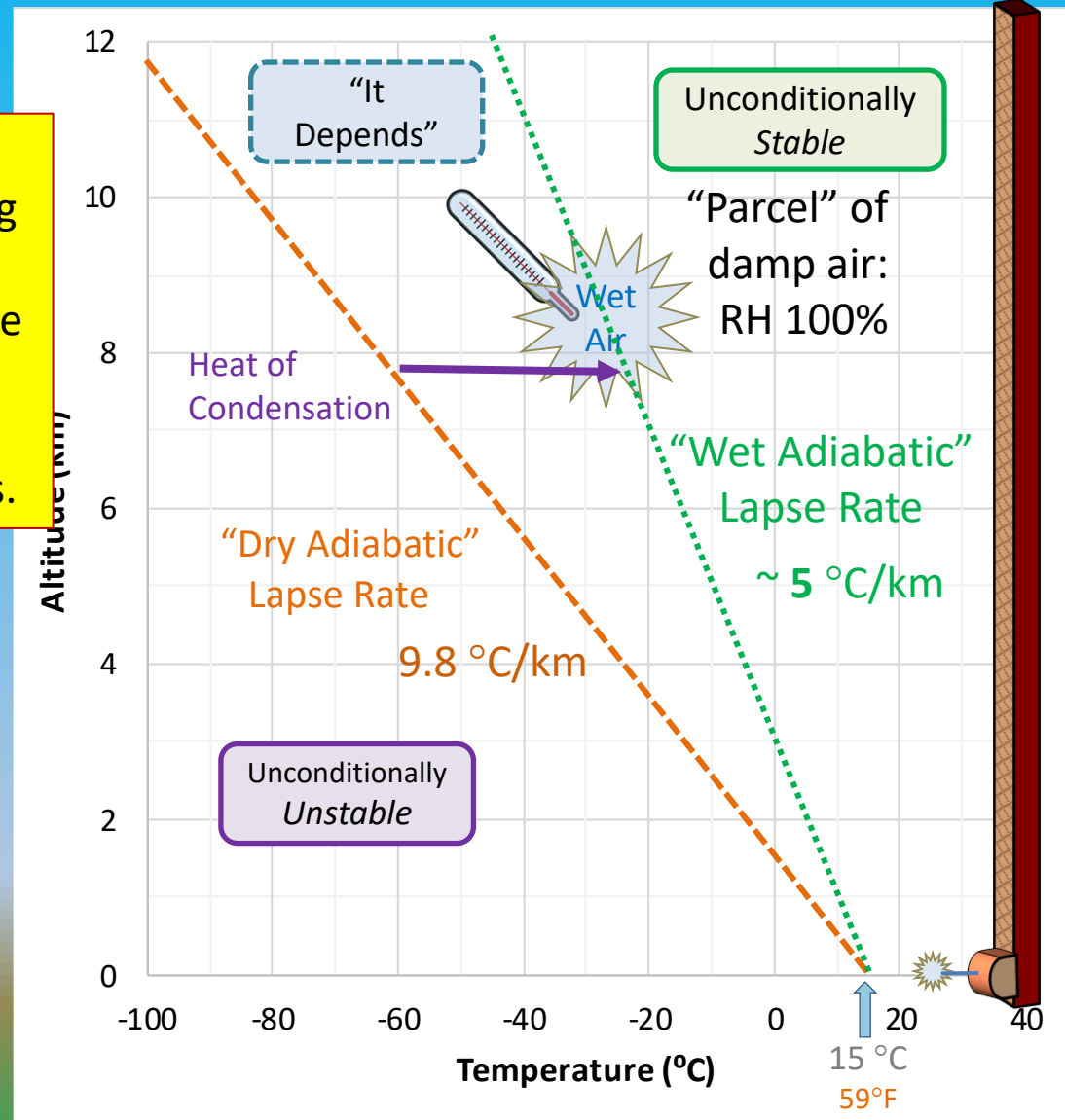
But air is wet....



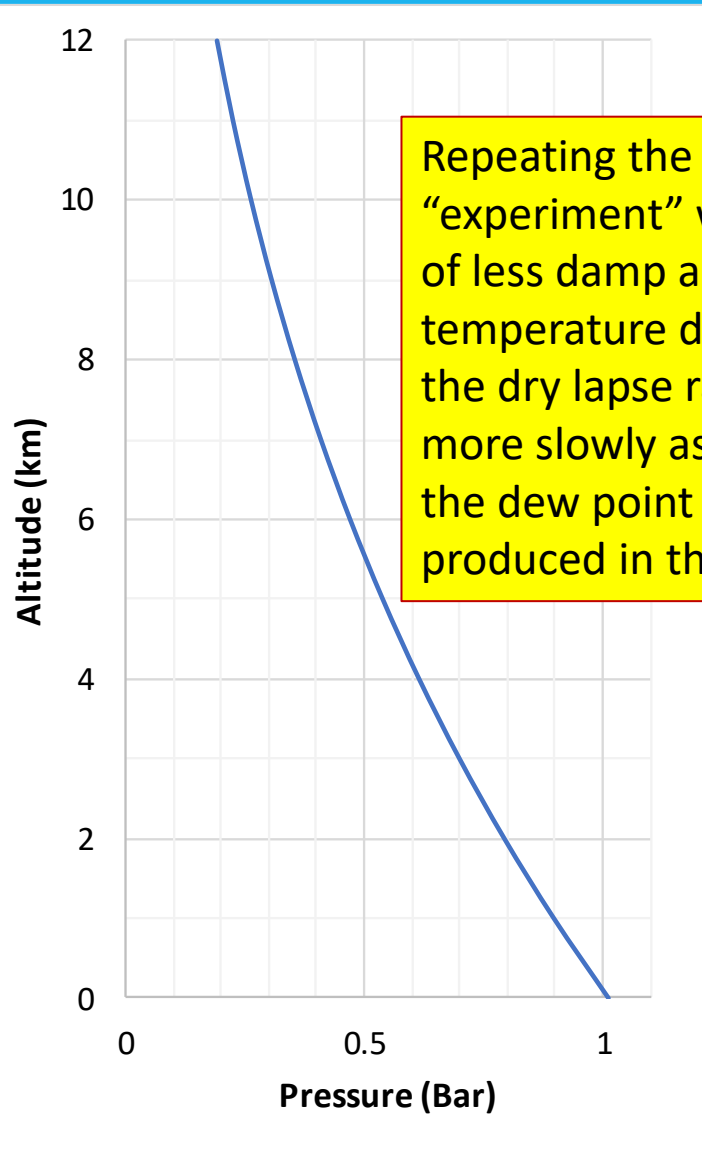
Lapse Rate in Troposphere



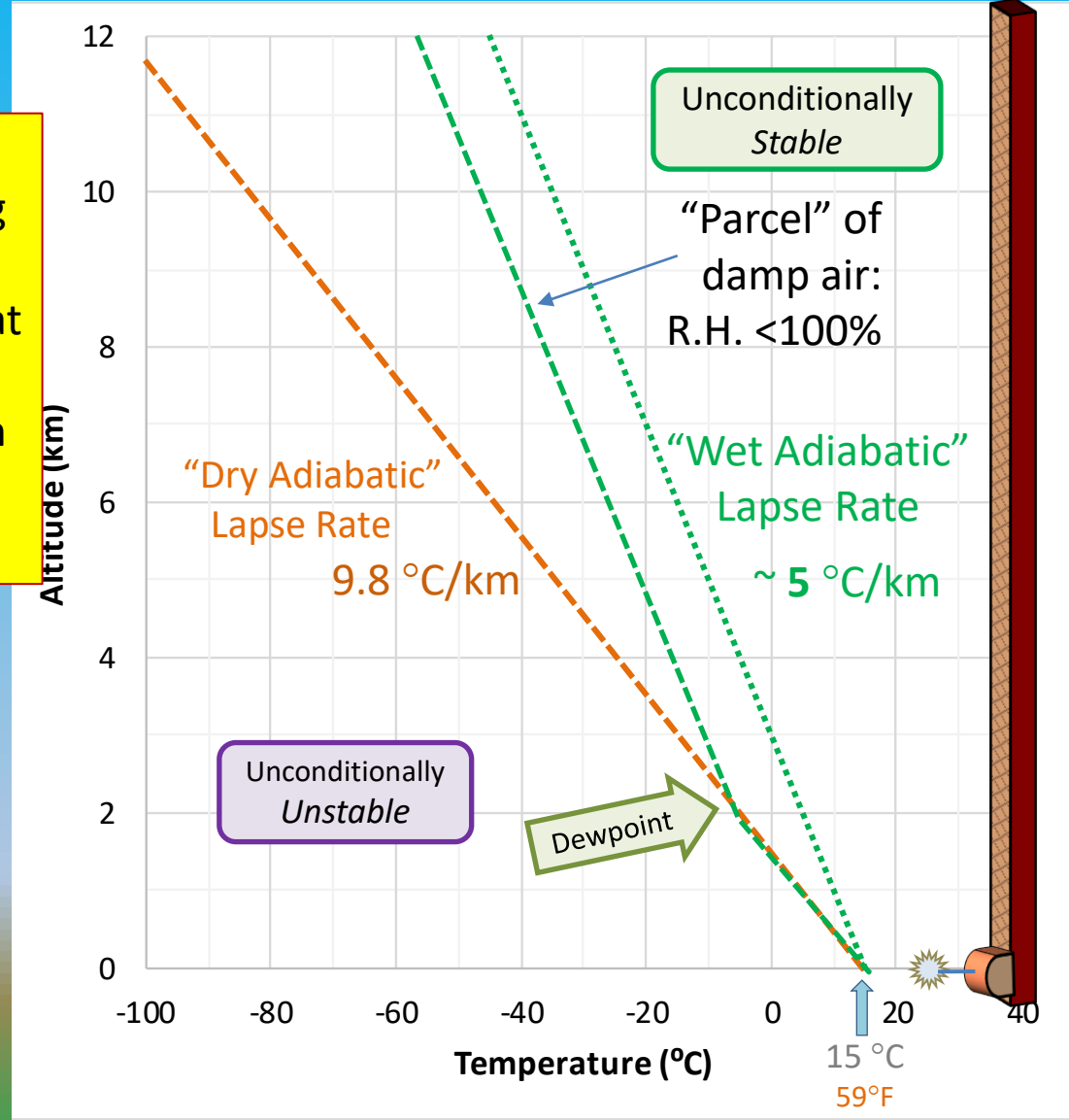
Repeating the "experiment" with a bag of damp air, the temperature drops more slowly as we go up, and fog is produced in the bag as water condenses.



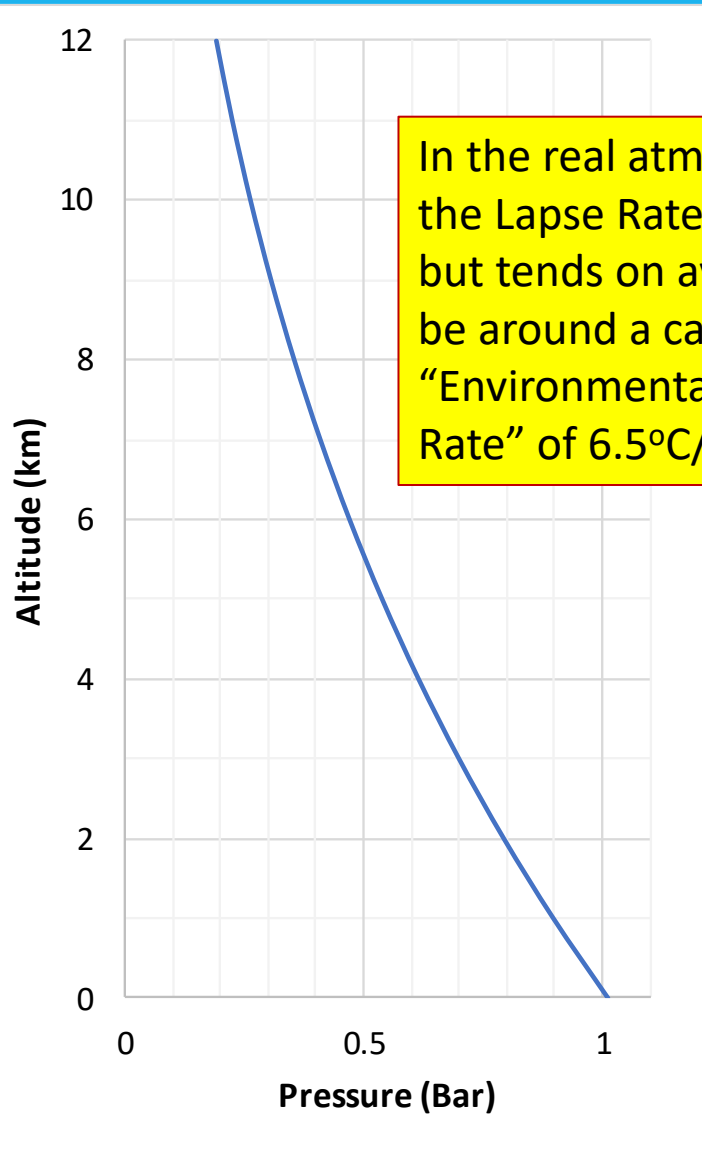
Lapse Rate in Troposphere



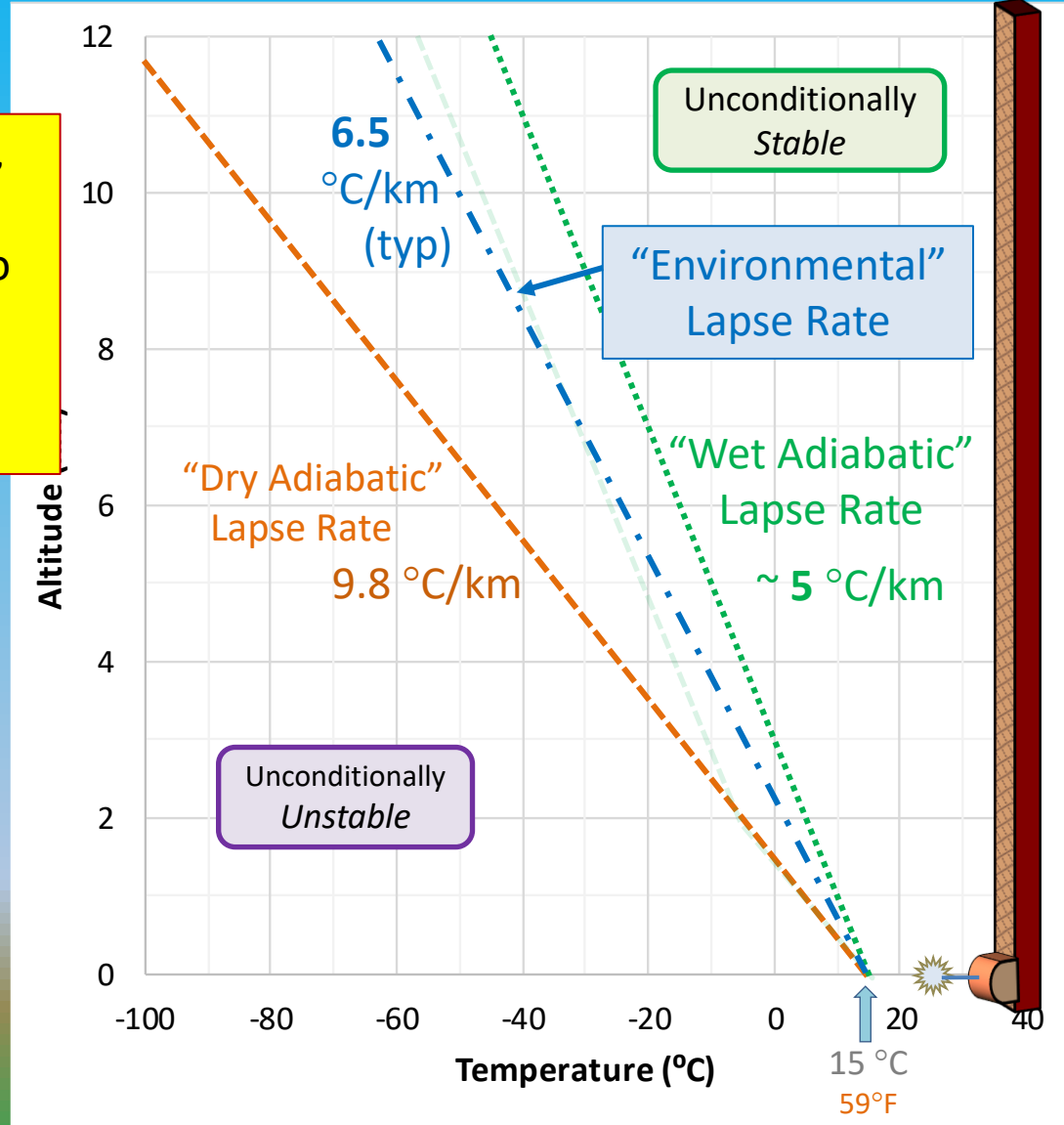
Repeating the "experiment" with a bag of less damp air, the temperature drops first at the dry lapse rate, then more slowly as we reach the dew point and fog is produced in the bag.



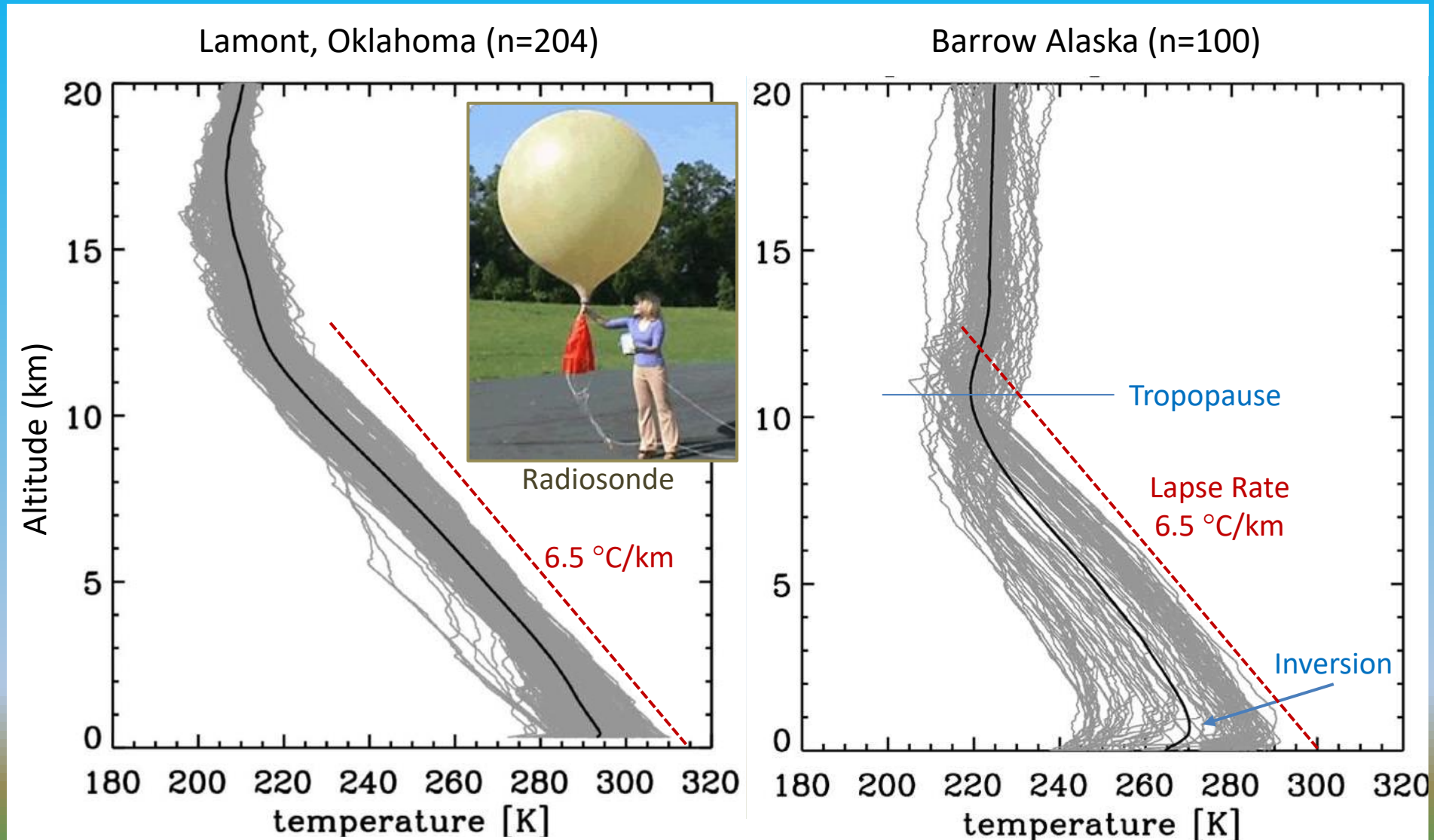
Lapse Rate in Troposphere



In the real atmosphere, the Lapse Rate varies, but tends on average to be around a canonical "Environmental Lapse Rate" of $6.5^{\circ}\text{C}/\text{km}$.



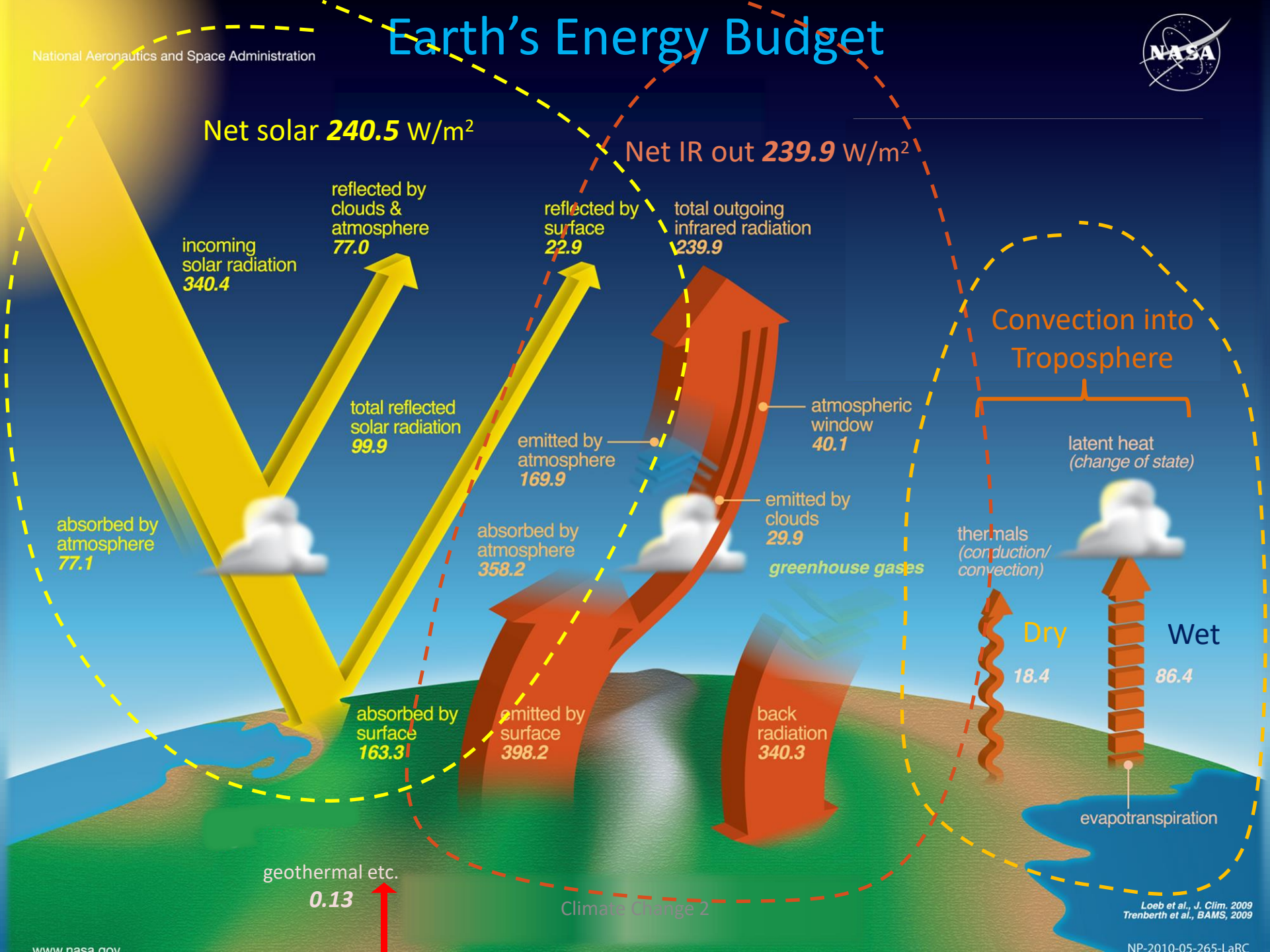
Actual Temperature Profiles in Troposphere



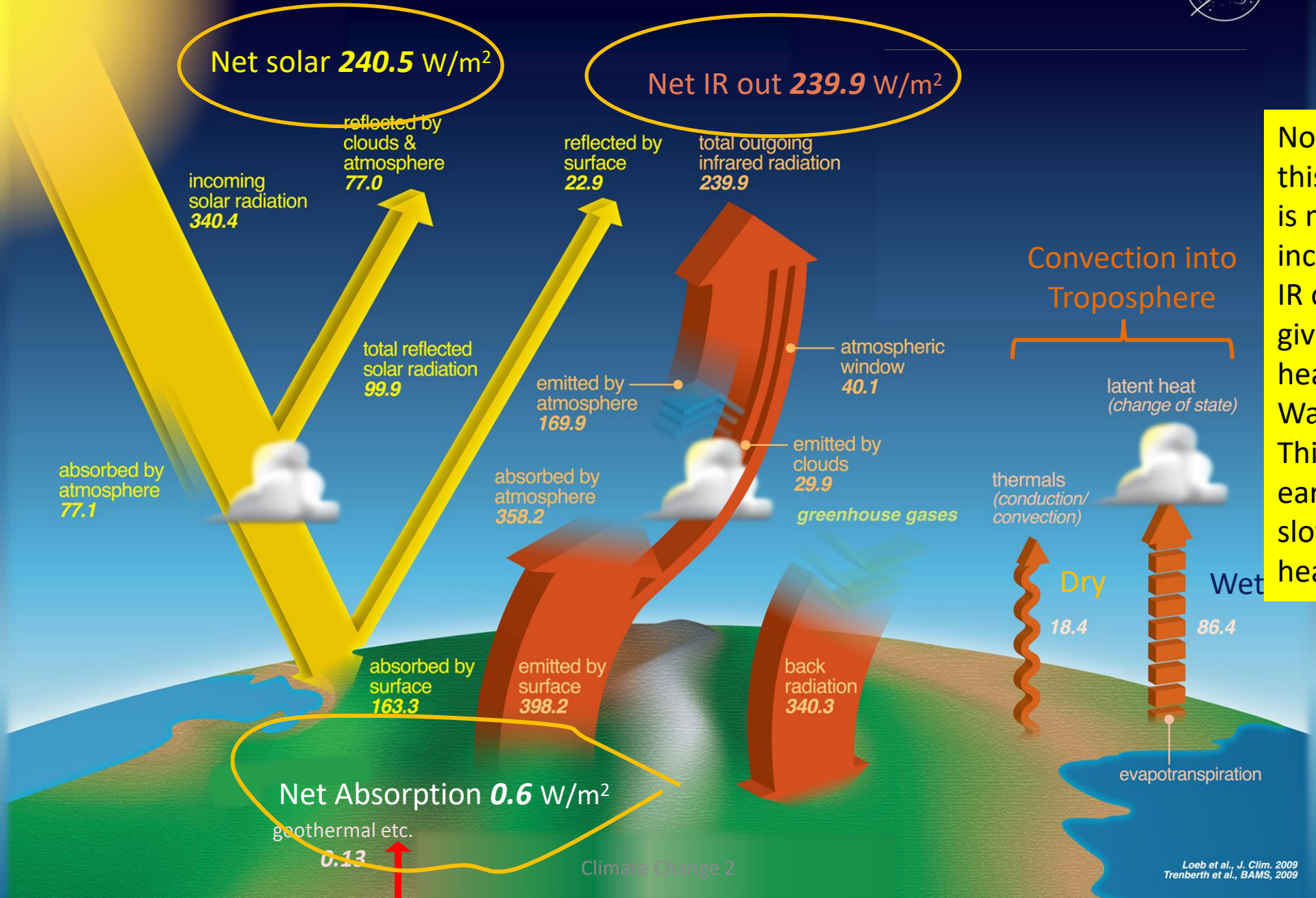
from Payne et. al., IEEE Trans. on Geoscience & Remote Sensing 46 (2008)



Earth's Energy Budget



Earth's Energy Budget

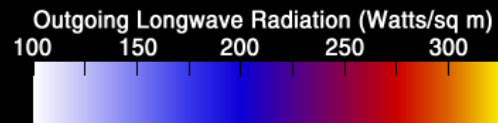
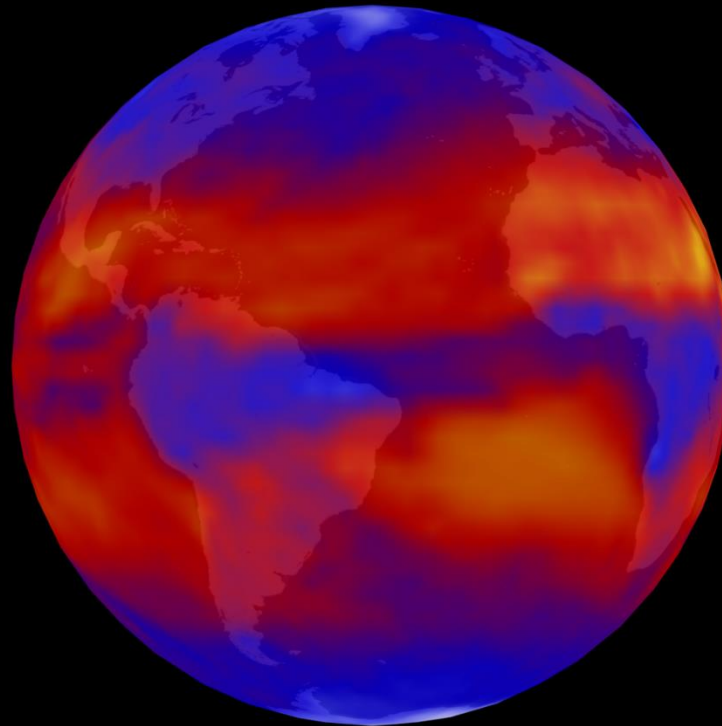


Note that in this case, there is more Solar incoming than IR outgoing, giving a net heating of 0.6 Watts/m². This means the earth must be slowing heating up.

Satellite Measurements of Outgoing Earth Radiation

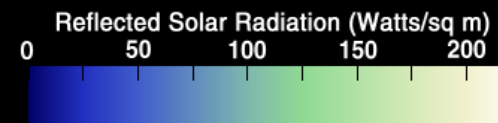
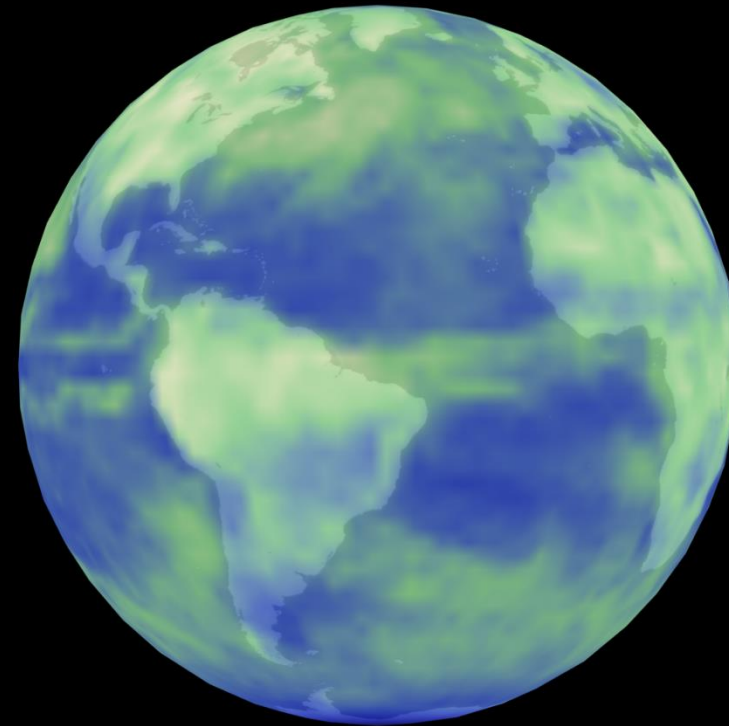
Note that both outgoing IR and reflected Solar are non-uniform and time dependent. Up to now we have been discussing averages.

LW IR Outgoing Radiation



Apr 2001

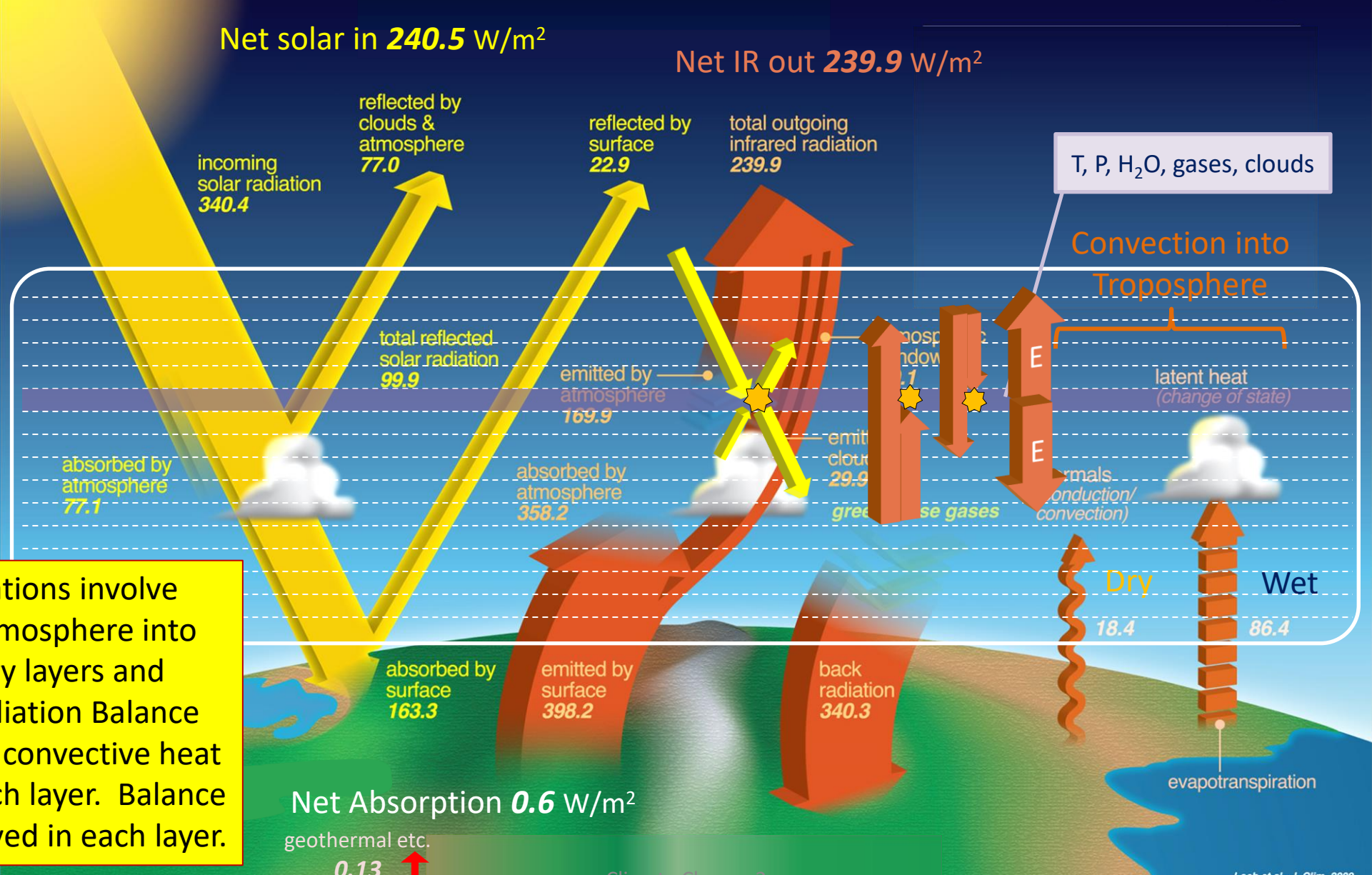
Solar Reflected Radiation



CERES Satellite Data

ceres.larc.nasa.gov/press_releases_images.php

Earth's Energy Budget



Climate calculations involve dividing the atmosphere into many imaginary layers and calculating Radiation Balance (together with convective heat transfer) in each layer. Balance must be achieved in each layer.

Earth's Energy Budget



Net solar in **240.5 W/m²**

Net IR out **239.9 W/m²**

incoming solar radiation **340.4**

reflected by clouds & atmosphere **77.0**

reflected by surface **22.9**

total outgoing infrared radiation **239.9**

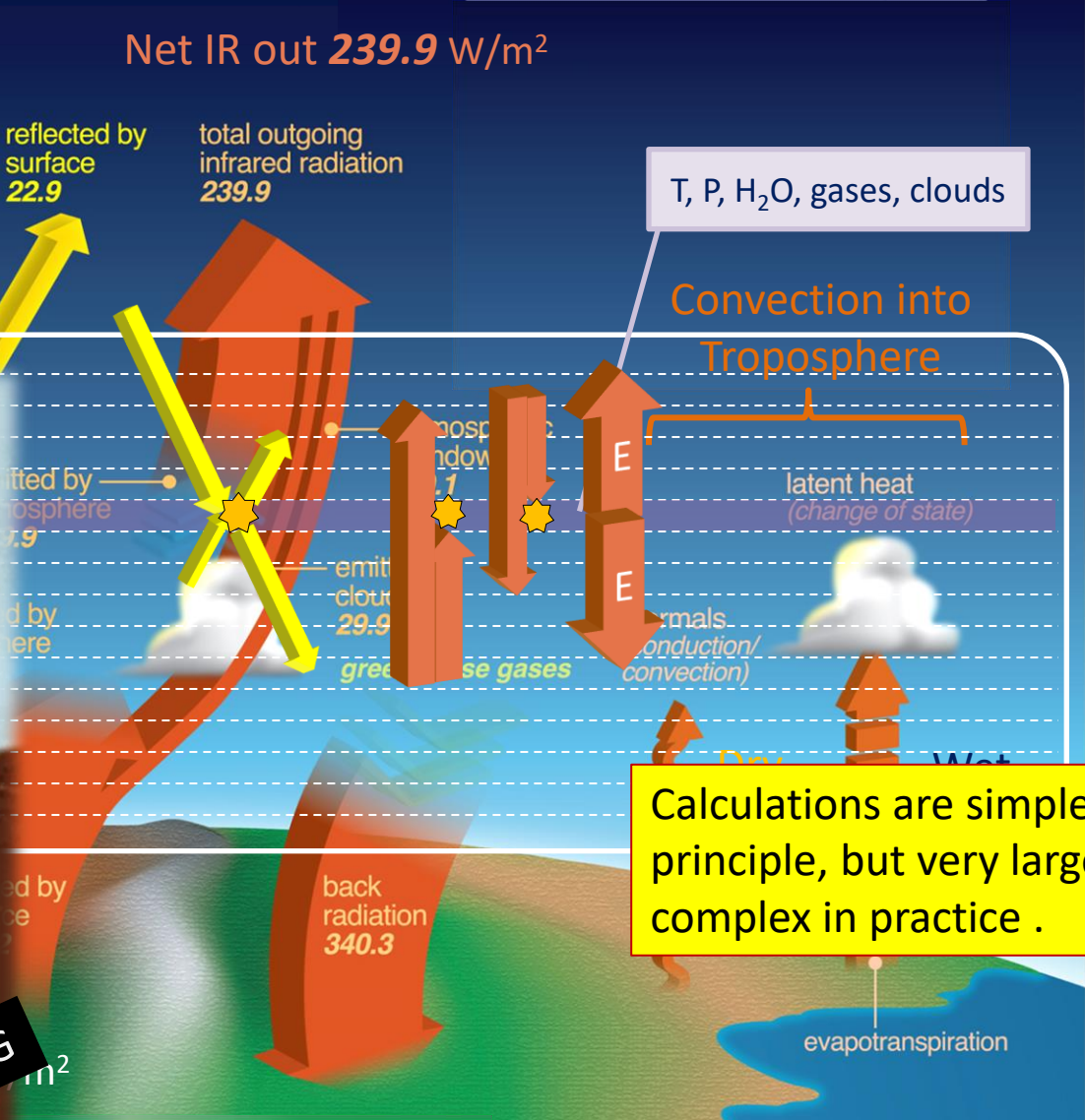
T, P, H₂O, gases, clouds

Convection into Troposphere

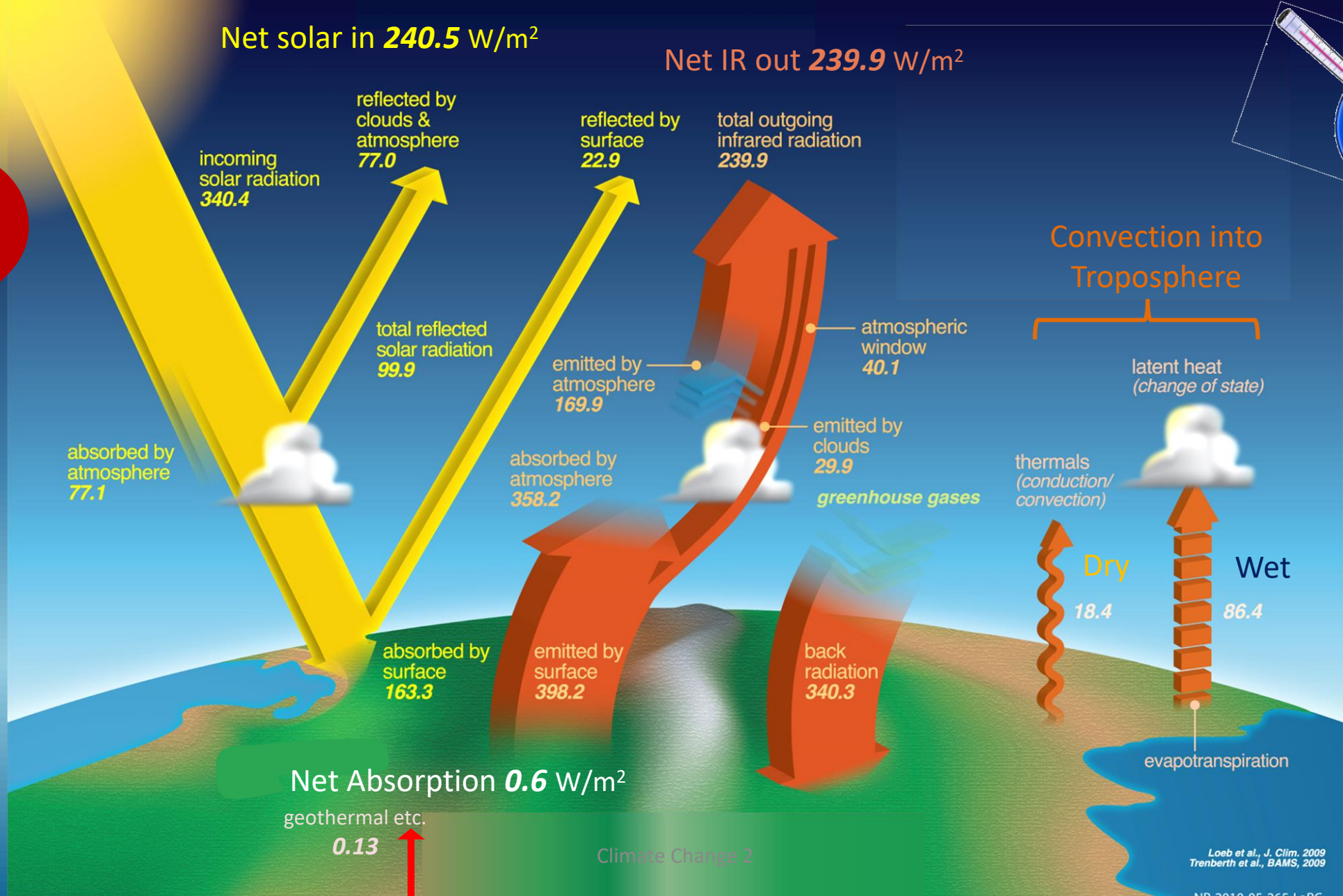
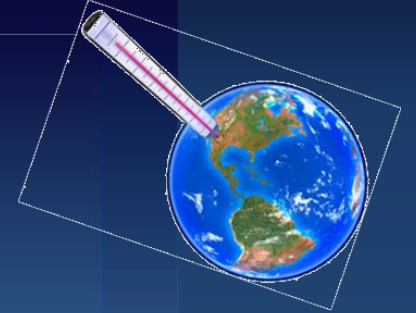
latent heat (change of state)

Calculations are simple in principle, but very large and complex in practice .

ACCOUNTING



Earth's Energy Budget



Course Outline



1. Building Blocks: Some basic concepts
2. **Our Goldilocks earth: a radiative balancing act**
3. The Atmosphere and its Gases. Modeling the climate system
4. Wild cards: the roles of clouds and aerosols
5. The Dynamic Earth System: Oceans, atmosphere, biosphere, cryosphere, people, plate tectonics
6. Natural Variability of the climate, short and long term. Ice ages
7. Carbon Dioxide, Water and other greenhouse gases: Where do they come from, where do they go, how are they regulated?
8. Future Projections: Impacts of GW and the uncertainties. Amelioration strategies.